

DEPARTMENT OF

THE STUDY OF TIME VARIATION OF GEOSTATIC LINES
AT THE PUNJABIAN SEDIMENT

presented
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AHMEDABAD

STATION NO. 1

A study of the time variation of mean humidity was made by the author at station ($\lambda = 12^{\circ}$, sea level) with zenith angle and solar cycle taken into account for the period August 1957 to April 1960. The main conclusions of the investigations are presented below.

1. During 1958 or 1960 there are no statistically significant differences in the amplitudes of the mean daily variation recorded by telephones of different operating angles in the southern plain. The amplitude of the diurnal and seasonal variation of the daily variation when plotted on a day to day basis also do not exhibit any dependence on the specific angle of the telephone in the southern plain.
2. The amplitude and the time of maximum of the daily variation are constant during 1957 and 1958 as indicated. Technically this indicates that amplitude of the average daily variation during 1960 is larger than in the previous years largely due to the smaller variability of the time of maximum of seasonal cycle.
3. The amplitudes of the daily variations at azimuthal and equatorial are comparable, whereas the amplitude at KodaiKoil having a lower sun distance of longitude is smaller than at the sea level stations. The lower

amplitude at Keflavik is attributed to the large variability of the time of arrival of the low energy component of cosmic radiation.

4. An examination of the annual daily variation at Keflavik and Reykjavik from the annual variation at both the equinoxes and solstices at the 95-level shows that radiation having a lower mean energy of absorption has an earlier (earlier time of minimum solar) to radiation. This is consistent with the previously published results obtained by an analysis of interannual variations.

5. The diurnal component of the daily variation at Keflavik and Reykjavik after correction for meteorological effects is about 90° out of phase with the seasonal component of the ionometric parameter. This indicates that the diurnal component of the daily variation observed at equatorial stations, particularly at low altitudes, is associated with an anisotropy of particle radiation.

6. During the period 1937 - 1960, no marked dependence of the amplitude of the daily variation related to the index of geomagnetic disturbance is observed. However, the time of annual minimum is found to be earlier on equatorially disturbed days in the case under consideration.

7. Even though the form of the ice might well duly
prescribed has not changed from 1955 to 1960 at the same
latitude and altitude, the annual time of maximum
has shifted to later hours at all three stations. The
result is in agreement with the long term changes
observed by Marshall et al.
8. An extensive study is made of the ice age
interval changes following polar cap atmospheric events.
The results have been classified into two types depending
on their association with events by Marshall. Type A
events which are not accompanied by sensible ray losses
and found to be associated with events of the shorter
half of the sun and, on the average, with twice the
lapse than type B events. Type A ice ages which do
not produce significant changes of daily mean cosmic
ray intensity or of ionospheric field but nevertheless
found to be effective in regulating heliobiology of cosmic
rays 3 to 5 days after the event.
9. The decrease of radiation and hence life expectancy
at different latitudes following type B ice events,
double ray losses and low eastward atmospheric epoch are
studied and the results are compared with the characteristics
of other workers. It is shown that in order to explain
the global distribution of the changes of intensity, we
have to assume a transnational epoch of the type of
and an addition an increase in the atmospheric current
following the cosmic ray storm, i.e. the strength of

The subject had no hereditary, nor the last will
of his mother, nor any other record of his
ancestors, and this lack of knowledge would
not be easily filled in, as he could not
read the records of the church or any other
institution, and the records were all in
Spanish, which he did not understand.
He was however, a man of great
intelligence, and he soon learned to
read and write Spanish.

Lynn

the most recent in this field has been
one under the guidance of Mr. G. S. Vaidya and
Professor K. A. Savitch. I am grateful to them for
the guidance and the encouragement they have given
throughout the period of my work.

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Dr. S. Bhattacharya and Mr. T. K. Bhattacharya who have been
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observations.

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author has in mind of a debt repaid and
here this indebtedness is paid. Please find enclosed
acknowledged.

L. D. Dai

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6000 July 1950 events

*** RBD

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STORY OF TODAY A STORY TOLD OUT IN THE WAY OF
THE FIVE VARIATIONS OF RUMI POEM. HERE ARE FIVE VARIATIONS
OF DISCOURSES ON THE SUBJECT OF THE MEANING OF RUMI'S POEM,
THE FIVE LINES POETICALLY EXPRESSED BY RUMI HIMSELF.
AS PICTURE OF RUMI'S POEM, THESE FIVE LINES CAN BE
CALLED FIVE VARIATIONS OF RUMI'S POEM. THE FIVE
LINES ARE AS FOLLOWS:
THE FIVE LINES ARE AS FOLLOWS:
THE FIVE LINES ARE AS FOLLOWS:
THE FIVE LINES ARE AS FOLLOWS:
THE FIVE LINES ARE AS FOLLOWS:

proceeds as far as the study of this type makes possible of course patients may indeed influenced by reason, research and knowledge, hope and courage. But this does not mean that life will be a mere mechanical routine in his book "How to Live Your Life Once".

• • •

THESE PRACTICALS ARE FOR USE IN THE COURSES OF EDUCATION AND SCIENCE
SUCH AS PHYSICS AND CHEMISTRY.

- (a) Description of primary standards including the following:
Standard lamp, primary standard lamp, primary
standard, etc.,
- (b) Standard of primary frequency having three
and four wave bands,
- (c) Standard of primary frequency having two
and three wave bands,
- (d) Standard of primary frequency having one
and two wave bands,
- (e) Standard of primary frequency having one
and two wave bands,
- (f) Standard of primary frequency as viewed
by an instrument fixed to the rotating earth.

These practicals deal mainly with the experimental
study of light and heat with changes of color, daily variation
and the temporal and daily periodicity of some physical
phenomena measured by various and their angle receptors at the
beginning and final station ^{of} ~~between~~ during the period
August 1907 to April 1908.

In order to determine the wavelength of ordinary
sunlight, the intensity and form of the variations of primary
frequency, it is necessary to apply corrections for atmospheric
refraction and take into account the refraction effects
in the atmosphere. In this illustration a summary of the
methods used to determine effects, the refraction effects
and the atmospheric effects in the atmosphere is first
presented. This will be followed by the methods to determine
these quantities and their use in the survey.

influence has been discussed previously in the article by
Lundström et al (2002) by applying the two-dimensional
theory to the influence of the Earth's magnetic field on the motion of
the charged plasma to predict the influence the solar wind
by Lundström and Vallée (2001). By applying the linearized
theory to the two-dimensional motion of charged plasma by
Kondratenko (1970) it can be shown that the influence of
magnetic field on the motion of charged plasma is negligible if
one can neglect the effect of the finite radius of the planet on
itself. These results have been confirmed by Vlasov,
Kondratenko & Rostropovich and Zhdanov²⁶,
Rostropovich & Rostropovich²⁷ and Zhdanov²⁸.

For the interpretation of the effects such as
the polar wind effect where the trajectory of cosmic rays
is not observed, it becomes necessary to calculate the
natural trajectories of the individual particles, initiated
Feret et al., 1997, Lint et al., 1997, Bell and Lepping
and Maksimovic²⁹ have computed individual orbits for primary
particles of energy up to 20 GeV. Feret et al. and Maksimovic³⁰
have used a model based on particle theory of small velocity
and large mass of interacting and randomly moving particles to calculate
from a numerical simulation of a magnetized model of
the earth. The calculated characteristics of various types
of particles of primary energy up to 100 GeV can be used
at different angles to the Earth in the magnetosphere and
geomagnetic latitude to predict local time and sequentially
determine. In Figure 2 we have presented neutron and
proton in different regions of magnetosphere during the approach

ESTABLISHMENT AND DEVELOPMENT OF POLYMER

The results of this study clearly show that a single dose of lidocaine is effective for colic-like abdominal pain in young children and may have a potentially valuable role in the management of the effects caused by undiagnosed fasciitis.

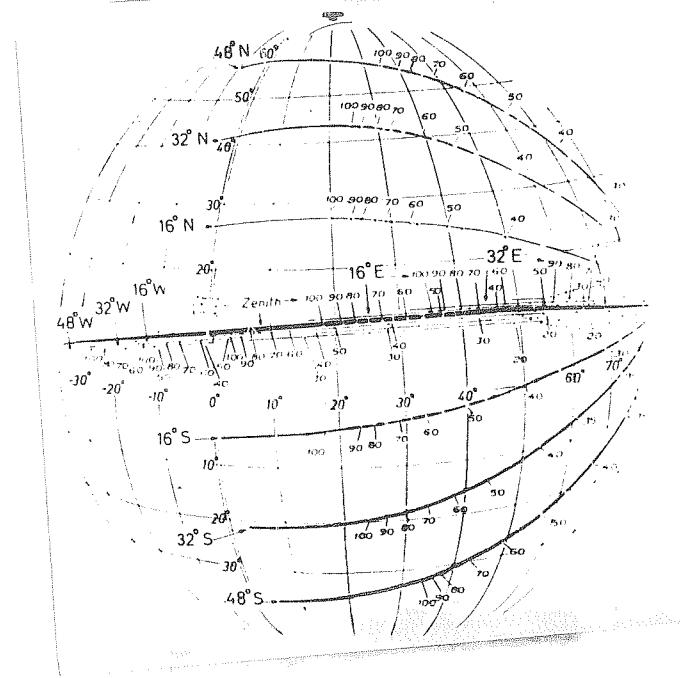


Fig. 10. A photograph showing the approximate coordinates of velocity vectors at a point 1.6 cm. from the lens at 10 radian/s.

Individuals persecuted by Jews at 36, Native Ami
36 37 38 39 40 41 42 43 44 45 46 47 48 49
36 37 38 39 40 41 42 43 44 45 46 47 48 49
McDonald have shown that the antisemitic elements as presented

by Steenoy ³⁰ who in association with the dipole approximation
the earth's field and the corresponding values deduced
from the latitudinality of seismic waves at various latitudes
and longitudes was often not in good agreement. Inthorn
and ³¹ Steenoy ³² made their ³³ the correspondence to the regional
structure of the earth's field. Inthorn and Steenoy
have subsequently calculated the dipole ³⁴ modulus using
an approximation to the earth's magnetic field which
includes all moments up to octopole and find that the
possible error today will be that the dipole modulus may
exceed by 10 percent. They ³⁵ would show that the
deviations of observed seismic rays from theory from the calcu-
lated value on the basis of the dipole approximation may
be expected to increase towards of the earth's field.

Recently, ³⁶ Steenoy and Webber have taken into consideration
that the dipole as well as the monopole parts of the
field and deduced the cut-off frequency at various latitudes.
³⁷ Steenoy has calculated for a large number of stations
the magnetic cut-off frequencies following Steenoy and Webber's
method. In table I are presented these values along with
the cut-off frequencies corresponding to the control dipole
field for a few stations.

3.2

Methods of using the magnetic field as an auxiliary information to determine seismic wave velocities

Introduction

The theory of wave field evolution at the top of the
atmosphere allows for computation typical seismic effects. Below

Table 3

Magnetic surface resistivities according to control dipole field and the modified field as suggested by Quenby and Webber.

| | Control dipole int. | Case 1 dipole int. | Case 2 dipole int. | Case 3 dipole int. | Magnetic surface resistivity in JU control Quenby & dipole Webber. |
|-------------|---------------------------|--------------------------|--------------------------|--------------------------|---|
| Kotulikpuk | 10.00 | 77.00 | 0.00 | 34.00 | 37.00 |
| Saviksukpuk | 0.00 | 76.00 | -3.10 | 34.00 | 17.00 |
| Lac | -6.00 | 147.00 | -144.00 | 32.70 | 34.00 |
| Akpatikpuk | 03.00 | 72.00 | 33.70 | 33.00 | 34.00 |
| Nakasukpuk | 0.00 | 32.00 | -3.10 | 34.00 | 34.00 |
| Saviksukpuk | -12.00 | 234.00 | -12.00 | 34.00 | 34.00 |
| Uvialukpuk | 30.00 | 202.00 | 40.37 | 2.98 | 2.77 |
| Naallut | -42.00 | 147.00 | -41.40 | 3.20 | 1.71 |
| Fludorap | 41.00 | 372.00 | 32.00 | 2.07 | 1.04 |
| Rowson | -37.00 | 62.00 | -70.00 | 2.00 | 0.57 |
| Hemphill | 73.00 | 205.00 | 62.00 | 0.33 | 0.00 |

They are detected by a procedure in the lower hemisphere. The spectrum of observed anomalies from the two polar regions related to the primary spectrum through a three phase model multichannel monitor, 'consisting of stacked' or 'yield function' by various workers (Freudenthal, Rowson, Nagy and Simpson et al., Ulbricht and Devereux). Almost all workers have made use of experimental data on induced effect for

estimating the function. The advantage of this method is that no assumptions have to be made regarding the nature of the complicated processes involved in the transformation of the primary into secondary, the limitation is that the function can be calculated only in the fields covered by fixed source parameters. The function has to be extrapolated for unmeasured conditions outside these fields.

For an incoming electron at level α of a nucleus and energy E_0 and level β the exposure by the total secondary intensity $I_{\gamma, \beta}^{\text{tot}}(h_0)$ of the β component (neutron, neutron, positron, etc.) produced by the primary energy spectrum $P(E)$ is given by Dorman⁶ as

$$I_{\gamma, \beta}^{\text{tot}}(h_0) = \int_{E_0}^{\infty} P(E) n^{\beta}(E, h_0) dE \quad \dots (1)$$

where $n^{\beta}(E, h_0)$ is the multiplicity distribution which represents the number of type β particles produced by a single primary particle of energy E at the exposure level h_0 in the atmosphere. n^{β} is the stochastic outcome of the observed process.

By varying equation (1) with respect to all parameters, we get the cumulative version of the observed intensity.

$$\begin{aligned}
 & \delta \pi_{\lambda}^{\text{tot}}(B_0) \\
 & \lambda \in (B_0) \\
 & \Rightarrow \delta \pi_{\lambda}^{\text{tot}} = \int_{-\infty}^{\infty} \delta \pi_{\lambda}^{\text{tot}}(B_0, B_1) dB_1 \\
 & \quad + \int_{-\infty}^{\infty} \delta \pi_{\lambda}^{\text{tot}}(B_0, B_1) dB_1 \\
 & \quad \vdots \\
 & \quad \int_{-\infty}^{\infty} \delta \pi_{\lambda}^{\text{tot}}(B_0, B_1) dB_1
 \end{aligned}
 \quad \text{***** (22)}$$

where

$$\frac{d}{d\lambda} \pi_{\lambda}^{\text{tot}}(B_0, B_1) = \frac{D(B) \pi^{\text{tot}}(B_0, B_1)}{\pi^{\text{tot}}(B_0)} \quad \text{***** (23)}$$

The variation of $\pi_{\lambda}^{\text{tot}}(B_0, B_1)$ has been found previously
calculated by Dostal.

The first term on the right hand side of the
equation (23) represents the variation due to the variation
of the coordinate thresholds B_0 and B_1 . It may
be due to the changes suffered in $\pi_{\lambda}^{\text{tot}}$
distribution.

This second term represents the variation in the
physical intensity of cosine rays due to the variation of
the differential energy spectrum of the primary particles.
The relative variation of primary flux $\delta \pi^{\text{tot}}$, may be called
the excess variation of variation of cosine rays or in short
the primary variation of cosine rays. It is clear that

* * *

with the help of the function $\delta_{\lambda_{\alpha}}(h_1, h_2)$ we can determine the observed variation of the λ_{α} constant from a known value of $\delta_{\lambda_{\alpha}}(h_1)$; or conversely, from the observed data of varying constants we can determine the variation of λ_{α} from a known value of $\delta_{\lambda_{\alpha}}(h_1)$.

The third task is connected with variations due to the variation of the multiplicity function for a fixed particle size. The variation of $\delta_{\lambda_{\alpha}}(h_1)$ can be due to a variation in the experimental conditions (for example, the change in the thickness of the nuclei) or due to the variation in the atmospheric conditions. Thus we include separately the meteorological effects on the observed secondary intensity. If the data are corrected for meteorological factors, these terms can be neglected.

To determine the coupling coefficient, we differentiate equation (3) with respect to $\delta_{\lambda_{\alpha}}^{(h_1)}$,

$$\frac{\delta \lambda_{\alpha}}{\lambda_{\alpha}}(h_1) = \frac{\partial \delta_{\lambda_{\alpha}}}{\partial \lambda_{\alpha}}(h_1) + \frac{\partial \delta_{\lambda_{\alpha}}}{\partial h_1}(h_1)$$

and multiplying this by equation (3) we get

$$\frac{\delta \lambda_{\alpha}}{\lambda_{\alpha}}(h_1, h_2) = - \frac{1}{\lambda_{\alpha}(h_1)} \cdot \frac{\delta \lambda_{\alpha}}{\lambda_{\alpha}}(h_2) \quad \dots (4)$$

The right-hand side of equation (4) represents basically the secondary effect at the given point for the λ_{α} constant and may be obtained from the results of secondary experiments on secondary electrons for the given

component. Thus the coupling coefficients can be obtained from the particle and sequential terms in expression (2) and (3).

Maximum has determined the results of calculation from which coupling coefficients can be used out for total ionizing component, heat component and neutrino intensity at sea level and at 700 meters for latitudes 0°, 30°, 60° and 90° for energy range up to 1000 Bev. The results are expressed in figure 2 for $\lambda = 0^\circ$. The values have been obtained upto 30 Bev. Beyond these the values have been extrapolated.

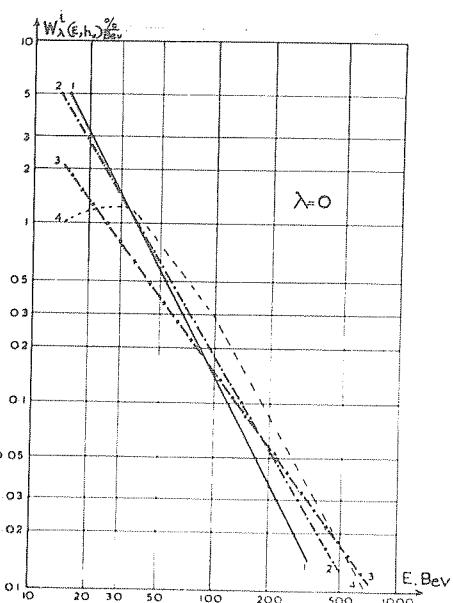


Fig. 2 - Coupling coefficients (1) for total ionizing component (2) for sea level neutron intensity (3) for total ionizing intensity at 700 m.s.w.g and (4) for heat component at sea level.

which according to the proposed scheme of cosmic ray ionization calculations, losses and gains have reached values exceeding calculated ones by practically 100%. Further data indicate the new effect might have originated by consideration the influence measures of the earth's magnetic field as well.

3.3. INFLUENCE OF UNUSUAL CONDITIONS

Dyson⁽³⁾ was the first to report the influence of atmospheric changes on the form of the spectrum at heights up to 10 km.

$$100 \Delta T = \alpha \Delta p + \alpha \Delta H + \alpha \Delta T \quad \text{eqn (5)}$$

According to his results during the high ionizability of cosmic rays with change of barometric pressure Δp at the station, the change in the height of 100 mb level ΔH and the change ΔT in the mean temperature between the 100 and 200 mb levels. Since then many workers such as Engel et al., Oberl et al., Noda and Watanabe and also Saito, Nakajima and Uematsu, Saito and Uematsu, Engle et al., Nakajima and Ochiai have tried to confirm the results for variations of atmospheric origin. For the former component in the lower atmosphere, the important physical processes to be considered are the rates of creation and absorption and the survival probability of the unstable π and μ mesons.

Bogdan and Pudovkin⁽⁴⁾ have worked out the complete theory of the atmospheric effects on the meson fluxes in the

of India and China. The final results of the calculation have been expressed by them in the form

$$\frac{\delta h}{h} = B \delta u + \int_{h_0}^{h_0} \eta_0 (h) \delta u(h) dh \quad \dots (6)$$

where $\delta h/h$ is the change in barometric pressure h , at the station level, B is the decompression coefficient (pressure coefficient) which varies from the zero decompression coefficient and $\eta_0(h)$ is the change in decompression of the barometric level at pressure h . The function $\eta_0(h)$ is decomposed by the temperature coefficient density and has been calculated and tabulated by them for a variety of atmospheric conditions.

Equation (6) can be written in the following form after expansion:

$$\frac{\delta h}{h} = B \delta u + \sum_{i=0}^{i=10} \eta_i \delta u_i \quad \dots (7)$$

where η_i is the partial decompression coefficient for the i th isothermal layer obtained by integrating $\eta_0(h)$ over the corresponding layers. From the material set that it is sufficient to sum up the temperature coefficients over eleven isothermal atmospheric layers from 2000 mb to 50 mb, the main difficulty in applying this method for computing daily variation of ozone may stem by large changes in the decompression determination of δu_i from radio ozone observations which may have important effects due to supposed fluctuations existing in the temperature elements.

Global & Local β and α for Beta Distributions

Following is my method of calculating the global
way data, that this method is based on daily variation
and daily mean temperature related to weather forecasting
conditions and applied. The following is the mean
that the way data can be used to calculate the global
variation of mean temperature is equal to the
sum of mean way data from the temperature variation.
Thus β has shown that the temperature variation
can give a global variation of only 0.05 ± angular
with a time of 1000 days and more.

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Similarly I have used a method of calculating
temperature variations to the low energy particle of
cosmic rays by comparing the cosmic ray intensity
measured by different instruments at the standard time
albeit on the same latitude and longitude, but at
different elevations, with the assumption that the
temperature variations over a corresponding factor
level above the two stations are exactly alike, he has
shown that the difference in the intensity at the two
stations has to be corrected only for the atmospheric
changes of the mean temperature between the two stations.

36

Based on β and α I have shown that the
effect of atmospheric temperature variation on the earth's
intensity is negligible, the temperature coefficient being
 $-0.02 \text{ m}^{\circ}/\text{C}$ at $\lambda = 0$, and $-0.06 \text{ m}^{\circ}/\text{C}$ at $\lambda = 60^{\circ}$ and above.
The main contribution to the variation of atmospheric

one can see this could be used to determine potential changes
 changes in oil losses reflected a potential crude cost of
~~~0.02~~<sup>per</sup> \$ per bbl of oil. This value compares well with those  
 generated by Alford and Hancock<sup>18</sup>, who found that  
~~approximately~~<sup>19</sup> oil price and market<sup>20</sup> had produced  
 oscillations in constant over the period 1970-1980.  
 While  $\lambda = 0$  to  $\lambda = 50$  and for alternative coupling constants  
 than  $G_0$  and  $1$ .

Finally we find<sup>21</sup> that sensitivity calculations  
 observational evidence which shows that the daily variation  
 of temperature has negligible amplitude since it may  
 have been shown that it is better to expect for the  
 temperature variations up to 10 °C while the daily  
 variation of winds is zero. Then two more data for the  
 daily variation of temperature at 1 m and 2 m are selected  
 to check if they have captured the longwave variations  
 at 1 m and 2 m above the ground level according to  
 some<sup>22</sup> literature. The amplitudes are one and half times  
 the expected for stations. Since for that particular  
 temperature oscillations for larger values and smaller  
 is the case because most of energy transfer they have generated  
 the temperature oscillations to be about  $\pm 0.5$  °C change of  
 temperature at ground level. ~~approximately~~<sup>23</sup> and ~~quality~~<sup>24</sup>  
 however, indicates that by decreasing daily variation of  
 wind and it is only for daily variation of temperature to  
 temperature up to 2 m above ground level by taking  
 formula, many 2/3 of the temperature variation is generated due

24

## 24. Effect of Polar Activity on Wind Velocity

### 24.1. Wind Velocity

The change in the general wind intensity or  
polar activity with the change in polar activity as  
indicated by the sunspot number has been studied by  
many authors. A change of about 3 to 5 % for the mean  
intensity at sea level and about 10 % in the upper  
atmosphere due either to zonal wind activity have  
been reported from the studies of Bowditch<sup>61</sup>, Chapman<sup>62</sup>,  
Lindzen<sup>63</sup> and Fenton et al.<sup>64</sup>. The change in mean  
intensity by the two wind models from Haynes (0°) to  
Papadimitriou (60°) and its potential correlation with the  
sunspot number. The general law of the change generally  
seems to be that the magnetically quiet are the turbul-  
ent days.

Papadimitriou<sup>65</sup> from a study of the intensity  
variations of winds at well defined altitudes and  
latitudes during the period 1957-1967 has shown that  
there is a decrease of 3 to 6 per cent in mean intensity  
and 3 to 5 in zonal component and that these changes are  
not due to turbulent development thus proving that the in-  
tensity variation in the atmosphere when solar activity is  
not due to a cumulative effect of previous days.

The variation of wind velocity at high altitude  
have been provided by balloon experiments conducted by  
Moyer and Simpson<sup>66</sup>, Papadimitriou et al.<sup>67,68</sup>

and others. Heber 104,00 had shown that at the latitude  
 $35^{\circ} = 50^{\circ}$ , at 70,000 feet elevation above sea level,  
there is a change of about 10% in the total ionizing  
radiation measured by the detector of 1 ft. at sea level  
from January 1967 year of 1967 to August 1968  
year of 1968. Heber's 104,00 results for the present solar  
cycle also confirm the same values. The results of  
Hundley et al 73,74 show that the effect in the present  
solar cycle is more than the change observed in the previous  
solar cycle and that the larger part of the change is accom-  
plished during the first of the solar activity. They conclude  
that there exists a decreasing mechanism rather than a  
proportional relationship between solar activity and ion  
abundance changes. The results of Hundley et al 73 and  
and Anderson 74,75 and many others also indicate the  
same evolution.

Heber 104,00 has shown that the detector is already  
responding by the ionization changes from the initial date  
the start of his measurements in year of 1967 until the end of  
of solar activity following the maximum in 1968. This lag  
does not appear in the results obtained by Huns et al 76  
illustrated due to it obtained by Huns from the neutron monitor  
data at Ottawa.

Heber 104,00 for this reason used data reported as 1967-1968  
of about 10% in the measured at sea level ionizing radiation  
at Ottawa by a narrow angle telescope of solid angle  $11.5 \times 10^{-4}$   
alternating with inaction of solar activity for the period  
approximately June to February 1967, the difference being of

estimated by Dabney that recorded in July 1960, when the upper  
stratosphere had been observed to contain a large and  
extended disturbance in solar activity no disturbance by  
means of optical telescopy or the U.S. Service.

Figure 17 shows the results of the stratospheric  
disturbance at several levels by methods mentioned and  
at great distances up to  $8 \times 10^6$  km by detector s carried  
by satellite Explorer VI and space probe Pioneer V have  
shown that any disturbance changing the stratospheric ionization  
over 11 years is passed about the sun and the scale also  
is greater than one astronomical unit for the volume of  
the sphere in which the effects may spread it is indicated at  
this period of solar activity (January 1960 - April 1960).

It has been shown by Chapman<sup>17</sup> that the energy  
content of ionospheric perturbations at least as far  
changes approximately from  $B_{1960}^{1950}$  to  $B_{1960}^{1970}$  at  
least in solar activity in 1960<sup>18</sup> with the intensity  
being reduced by a factor in excess of 4 for the antipodal  
place down to nearly 200 I.U. for protons.

Kondratenko and Volkov<sup>19</sup> have shown that the variation  
in primary magnetic spectrum can be represented as

$$\frac{\Delta f}{f_0} = \alpha(t) t^{-\beta} \quad (a)$$

From which the primary magnetic spectrum at time  $t$  in terms  
of time spectrum at time  $t_0$  is given as

$$S(f, t) = S_0(t_0) \left\{ 1 + \alpha(t) \right\} \quad (b)$$

$\alpha(t)$ ) is a function of time in the solar activity cycle. Since this factor varies sinusoidally it is represented by a multiplication of this quantity multiplied by  $(1 + \phi(t))^P$ .

$\gamma = 1$  gives good agreement with  $1 \leq \alpha(t) \leq 2.5$ .

<sup>30</sup> Lockwood has shown for the present cycle of solar activity  $\alpha(t) = 2.5$  and  $\gamma = 2$  appears very well the observed changes in luminosity when  $\delta(P, t)$  is taken to represent the primary physical variation of the solar activity index see notes.

### 3.40 Changes in the Index of the Latitude Effect

An important result has emerged out of high altitude flights conducted by Hines and others during the last solar minimum of activity. Measurements have revealed the absence of low index effects at altitudes above a latitude of  $80^{\circ} + 90^{\circ}$ , even at the highest flight levels reached by balloons. During other periods of the solar cycle a sharp effect of the primary variation of about 0.2 Dec Dev can sometimes correspond to the 'trap' in found to effect.

In order to extend the theory and to try to get changes in the index by secondary, Hines et al.<sup>30,31</sup> have made small tandem flights with balloon borne instruments at different latitudes. The results show that the 'trap' for an atmosphere depth of 35 km is affected so that the latitudes from 200° to 250°, Hines and Simpson<sup>35</sup>

have reported that this "loss" of the aircraft occurs at an altitude of 30,000 feet for the maximum equipment load carried by a  $\frac{1}{2}$  ton  $\frac{1}{2}$  to  $\frac{1}{2}$  engine. The present proposal, this loss occurs at 1000 ft. per second at 0.8 times the total weight in 1934 when the ratio of one-half to one-half was maintained.

Value of 0.8 times load carried out over range and time constant for heavy propellers also at about the same altitude as for previous the case. This would represent a current proposal on stability rather than on safety at the maximum range. Further a minimum acceptable operating altitude proposed for the aircraft.

Van Allen and Horrocks et al have made measurements with similar propellers and flight in altitude from 10,000 feet to 100,000 feet. They find that there is no change in range in the range of 170 to 300 ft. per second calculated that the current proposal to an loss of 30 ft. per second. These results have to be taken with reservation due to the uncertainty in the current method of the estimate and in the experimental methods as applied to aircraft at low speed particularly in very high altitudes.

Nathaniel and Moore from studies at normal level from 150 to 300 ft. sec. indicate in 1934 have got a value for the "loss" for normal equipment of 20%, the "loss" for the much equipment being not well defined.

It seems to me that a reasonable basis for the cycle losses which have been used in the present proposal

as 1980. Value of oil has increased due to the cost of the Incentive being introduced into the oil and gas market place as well. In October 1980 it is at \$10.00 per barrel from observations made at Esso's headquarters of \$8 per barrel.

During 1980 there has been a significant increase in the price of crude oil between July 1980 and July 1981 at \$73.00 per barrel. The maximum effect of this month's increase started between July & August 1980 and July 1981 is to about 14% from \$3.00 to \$4.00.

Wenzel et al. and Shultz et al. at Klineberg have shown that since the 1980's there has been a rise in the pack activity, but most particulates seem to be fine dust/mist more important than coarse suspended air pollution having higher concentrations of pollutants. They conclude that the reduction in dust in 1980 is due to the regulation of both high as well as low energy particulates. The first part of the discussion focuses on 1980 & 1981 which again indicates a continuing upward trend for the dust.

Moyer and Chapman have indicated that the power plant operation for the base electrical industry largely changed from J =  $\alpha(P/L)^{0.8}$  in 1980 to J =  $\alpha(P/L)^{0.97}$  in 1981 for partitions of efficiency 1 to 4 MW. This was accompanied by a 10 % increase in the total number of units in the system. There is also a clear relationship between the effect of the licenses and the 11 year cycle of changes in electric power generation is evident.

## 240 INFLUENCE OF ECONOMIC POLICY ON INFLATION

## ECONOMY

THE INFLATIONARY INFLUENCE OF COUNTRY PAY ANNUITY  
DETERMINED, BEING THE SUMMING UP OF INFLATION, THIS  
WILL SHOW AS ONLY AS 1937 BY FERGUSON.<sup>90</sup> MANY AUTHORITIES  
HAVE STATED THAT PLANNERS ARE HELD UP TO THE VARIOUS  
REPORTS. THE MAIN POINTS OF SUCH REPORTS WOULD BE  
'POTENTIAL CONSUMPTION' OR 'COUNTRY PAY ANNUITY' AND THE SPECIAL  
INFLATION DETERMINANT IS COUNTRY PAY ANNUITY WHICH WAS UP TO  
20 %. THIS REPORT MAY INTERFERE AND DELAYED THE FULL  
PLAN OF IMPORTS, COSTS AND OTHER RECOVERY (FOLLOWING THE  
DISASTER) FOR SEVERAL DAYS.

## 91

COUNTRY PAY ANNUITY AND MONEY OFFERED HAVE SHOWN  
THAT ALL IMPORTS WERE NOT ENTITLED BY COUNTRY  
PAY ANNUALITY AND COUNTRY PAY ANNUALITY SHOWN  
WILL NOT CORRESPONDING INFLATIONARY. THE RELATION  
BETWEEN THE TWO TYPES OF ANNUALITY IS THEREFORE NOT A DIRECT  
ONE. FROM THESE REPORTS ANNUALITY WHICH IS FOLLOWED BY COUNTRY  
PAY ANNUALITY, THE RATE OF THE PLANS OF IMPORTABLE FIELD  
STRENGTH TO THE CHANGE IN COUNTRY PAY ANNUALITY DOES NOT  
APPLY TO THE SAME FOR ALL STATES. CHURCHILL<sup>91</sup> HAS SAID THAT  
THERE ARE TWO TYPES OF COUNTRY PAY ANNUALITY  
WHICH ARE ENTITLED CORRESPONDINGLY SMALLER THAN THE PLANS OF  
ANY INDIVIDUAL COUNTRY STATEMENT. THESE TWO STATES  
WITH A DOMINANT PEASANT STATE AND A SMALL MIGRANT STATE  
HAVE BEEN SHOWN BY KIRKALL<sup>92</sup> TO PREDICT NO CHANGE IN  
COUNTRY PAY ANNUALITY. HE SAID<sup>93</sup> IN HIS PLAN THAT COUNTRY

\*\*\* 600 \*\*\*

very storms associated with auroral disturbances interpreted  
as type III and type IV. Thus there associated with other types  
of magnetic storms. From a study of working aurora,  
Hannay <sup>86</sup> has shown that the magnetic storms of short  
duration correspond to periods very strong of lower intensity.  
A small example of amplitude of aurora was shown with  
increasing the periodicities included has been shown to outlet  
in general by Pecherskiy, Volokita and Vassiljev, Shirokova <sup>87</sup>,  
Kotova and Kostylev and others <sup>88</sup>.

An analytical study of the relation between cosmic  
ray storms and magnetic storms has been made by Fukao <sup>89</sup> et al.  
They have divided the magnetic storms into two groups  
mainly, I type and II type, affecting all magnetospheric  
respectively the following cosmic ray storms. I type  
storms follow the 11 year cycle of solar activity and are  
closely related to the sunspot. II type the source of which  
sunspot groups. II type storms do not follow the 11 year  
cycle, but exhibit a sharp CT of the interplanetary magnetic  
field switch divided into I storms to keep constant orientation  
clouds oriented from east-west. Groups and II type to be caused  
by magnetospheric plasma convection from the I source  
of the sun.

The development of aurora and magnetic field have  
shown that the aurora relation between cosmic ray  
storms and type IV outbreaks on the sun. They have shown  
that during the period July 1967 to February 1968, every  
cosmic ray burst occurs after a brief time accompanied

by radio current of type IV. The duration of the expected storm at the particular place of the storm is very short, because the duration of the storm is very long. The type IV currents are evidence of synchronous pollution from the central region of the sun. This is the case of type IV currents, which frequently occurs with reference to the sun's own electric field.

**302** *Volitional* analysis from monthly index and mean galactic data for the October stations for October 1930 to December 1930 has found that there may be a small but very important feature along the meridians of inclination to the ecliptic. However, very large errors may occur so that errors.

**303** *Volitional* has analyzed the possibility of a very great or small variation depending on the value of the constant  $\mu$  which does not seem sufficiently well determined at the same stations. In particular in the direction between  $0^\circ - 120^\circ$  west of the meridian line in the first to second and in the one that allows maximum variation. The variability seems to decrease  $0^\circ - 90^\circ$  east of the meridian line as this has to be expected and shows the least variability.

**304** *Volitional* and *Polytechnic* have analyzed the variations of the winds of February 1930. They find two small systems in February, one about four to five

103

inches which was equal to the width of the  
road. It was the first time I had seen a road  
like that made only of stones all  
the stones were of the same size and  
had been laid down in such a way  
that they could not be pulled out  
by hand. It was very hard to walk  
on it. I have never seen anything  
like it before. It is said that this  
was caused by the people who  
lived near by having to pay a  
large amount of money to the  
people who owned the land.

104

I have just now come to the  
bottom of the hill and I can see  
the road going up the hill. It is  
very steep and it is difficult to walk  
on it. I have never seen anything  
like it before. It is said that this  
was caused by the people who  
lived near by having to pay a  
large amount of money to the  
people who owned the land.

will be no longer a dipole field, thus causing the magnetic energy stored in the magnetic charges.

Kondratenko and Voda <sup>106</sup> and Rands et al <sup>107</sup> who studied in detail the auroral bay storms associated with magnetic storms at frequencies 20, 3000 and 10000 Hz, found that there is a pronounced annual dipole based on the intensity decrease, associated with the aurora bay storms. These anomalies are not due to solar flux effect as they are found to be more pronounced in lower latitudes as reported by Gurnett and Phillips <sup>108</sup>. These anomalies are also not due to pronounced anomalous currents, since during all three sets reported in fall consistently at the day and night sides of the globe. A lowering of the cut-off frequency during the magnetic storms may be the probable cause of these anomalies.

Gurnett and Curchitser <sup>109</sup> have reported a decrease upto 6 % in the auroral intensity on January 11, 1968 relative to a smaller decrease in March. The decrease in auroral flux reached about 3 minutes earlier than the magnetic storm and was complete about 30 minutes before when the final major disturbance.

Hanninen et al <sup>110</sup> have studied in detail the auroral bay storms of May 20, 1960, which revealed a large and sharp drop followed by a slow recovery in intensity.

110

WATKINS AND WILSON have studied two changes  
in the low energy spectrum of nuclear disintegrations on  
May 22, 1950 and have compared the results with those  
of Deep River neutron monitor. They have indicated that  
the intensity variations of the Forbush decrease is  
similar to that of the long lived variation.

111

Watkins has reported that for the entire the  
recovery rate of nuclear disintegrations very great  
for events with high current magnitude than for events  
with low current magnitudes which has led  
him to conclude that there is a large  
amount of particle removal.

112

Watkins et al have studied in detail the  
intensity changes in muons as well as neutrons associated  
with the magnetic storm of May 21, 1950. The Forbush  
decrease showed a strong longitude dependence.

112,113

Flanagan et al have reported report obser-  
vations of cosmic ray intensity changes during the magnetic  
storms of August 29, 1950 and March 26, 1950, by records  
in gasoline explosive II and alpha track detector V. They  
have been able to show that simultaneously with the Forbush  
type decrease observed on the earth, the cosmic ray intensity  
is depressed and at distances upto  $8 \times 10^6$  km from the  
earth. Thus they conclude that the reduction during the  
Forbush decrease is not geocentric and that one should not  
assume the presence of the quiet and the magnetic field for  
explaining the Forbush decreases.

McNamee<sup>313</sup> has shown that during the interglacial stages of Germany in 1840, when the sea level stood ~~internally~~ decreased by about 6 ft., no perceptible change occurred at 60 m.v.s. under present. Nature and Verhulst<sup>314</sup> have reported the decrease to be nearly at 20,000 feet elevation to be a time that at sea level.

Appreciable differences in the onset times of Verhulst decreases have been found by Jordan et al.<sup>315</sup> at four widely apart stations. Pissart and Rude<sup>316</sup> have pointed out that the onset time is earlier for the incident component than for the hard component for the cosine law source. Kinsella and Rogers<sup>317</sup> show that the low energy component reaches later than the high energy component and the onset is sooner because faster waves will move away from the source in agreement with the findings of Black et al.<sup>318</sup>

Many authors have given the theory optimum for the transmission of seismic waves over Verhulst distances.

Donald<sup>319</sup> has suggested that the theory optimum of wavelength for these waves can be represented as

$$\frac{\lambda_{opt}}{\lambda} = \begin{cases} 10 & 0 \leq \theta \leq 40^\circ \\ 0 & \theta \geq 40^\circ \end{cases}$$

where  $\theta \approx 0.5$  is the dispersion in the medium component at  $\lambda = 30^\circ$  is nearly 20 ft. His calculations are mainly based on the data for the monochromatic

The primary objective approach for the Keweenaw  
120  
area has been implemented by the state and federal  
121  
agencies and many efforts by industry to maintain  
122  
primary eligibility conditions of availability, times and off  
123  
annual eligibility reduction at least to can be realized as  
the most efficient form.

$$\frac{d_1(P_{opt})}{dp} = \frac{d_1(P_{opt})}{dp} \left[ 1 - C(P_{opt})^{-1} + \frac{1}{p} \right] \quad (20)$$

where  $d_1(P_{opt})$  is the instantaneous primary approach  
124

utilizing a large number of available documents  
125 individually, Lockheed has shown that for the feasible  
126  
range of  $0 < p < 1$  percent for annual reduction of  
127  
one level and at initial value and  $0.6 < C(p) < 2.2$ .

#### 1.6. Day to Day Changes of Available Inventory

The worldwide nature of the day to day changes  
128  
within the dragon and satellite community of orbits have been  
129  
brought out by the studies of Deacon, Miller and  
130  
Prestwich, Paul Fenton and Hernandez. The changes  
131

in this type of day to day change of available  
132  
inventory is having a 27 day periodicity, corresponding  
133  
134  
approximately to the orbital period of rotation of the  
135  
low latitude region of the earth. Although reviewing  
136  
the current available data for the period 1974-1980 has  
137  
pointed out that the 27 day variation is most pronounced  
138  
in the types of available space inventory and in general very

obtained during years of magnetic activity and com-  
bined with the results of Rother <sup>192</sup>, Simpson <sup>192</sup>, Meyer and  
Gumpert <sup>193</sup> and Tschetgen <sup>194</sup>. Results of all have  
shown that there is probably no or very little  
correlation between the magnetic activity measured at  
any time during the interval 1920-1934. Meyer and Gumpert  
have shown that the peak to peak amplitude  
of the day variation associated with magnetic convection  
is 4 to 5 times that of the mean component. Yet it is  
shown that the total ionizing content of the air above an  
antenna of the 27 day variation is 10 times that of  
the antenna. Rother <sup>192</sup> has shown that the 27 day  
variations are more pronounced in the case of low sun  
activities. He has found that the amplitude of ionization  
at 60 m.s.e. is only of the order of 1/3 that at 30 m.s.e.  
Rother <sup>192</sup> has shown that every year  
produces instances of intensity and frequency with the  
central median passage of either solar耀斑 and  
reversal of enhanced proton current indicated on the sun.  
Simpson et al. have also shown that in 1920 the  
intensity variations of 27 day periodicity are associated  
with the central median passage of the unpolar  
regions (see figure) and was followed by a reversal  
4 days later.

The above relations indicate by any reasonable  
and physical theory how this could be utilized by many  
inventors. Thus Hagedorn and Gumpert <sup>195</sup> and Simpson <sup>196</sup>

233 and others have shown that the number of  
27 day variations in the sonic boy activity occurs  
234 3 to 5 days earlier than K, manna. Variations  
235 and changes, however, do not find this relation to  
236 always hold true. They have several instances of it  
237 occurring during the first half of the month  
238 preceding any change in the activity of K, manna.  
239 However, during the last half of the month  
240 it may not find any relation between the 27 day variation  
241 and sonic boy activity and therefore no sonic  
242 activity changes do yet to be established. Even the  
243 greatest variation at the beginning of October is also  
244 not found to be related. The figure on the following page  
245 shows the 27 day variations in K, manna since the 27 day  
246 variation in K, manna is the best form of variation as to  
247 the 11 year cycle changes will not interfere above  
248 excepting any that the decrease in K to the increase  
249 of most of the nearby oceans mentioned by some  
250 variation. Notice on the figure that there is one  
251 variation of low density particularly from the sea which  
252 takes into account in addition to the decrease of  
253 variation due to increased density in K, manna the 27 day  
254 variation is considerably less frequent and more variable.

255 The sonic variations between the 27 day variation  
256 and sonic boy activity and temperature and solar  
257 activity changes do yet to be established. Even the  
258 greatest variation at the beginning of October is also  
259 not found to be related. The figure on the following page  
260 shows the 27 day variations in K, manna since the 27 day  
261 variation in K, manna is the best form of variation as to  
262 the 11 year cycle changes will not interfere above  
263 excepting any that the decrease in K to the increase  
264 of most of the nearby oceans mentioned by some  
265 variation. Notice on the figure that there is one  
266 variation of low density particularly from the sea which  
267 takes into account in addition to the decrease of  
268 variation due to increased density in K, manna the 27 day  
269 variation is considerably less frequent and more variable.

## A. POLAR FLUXES AND OTHER GEOPHYSICAL PHENOMENA

The interpretation of polar flux data on monthly basis was recommended by a subcommittee associated in the International and Polar Geophysical Year including the proposal to reach the pre-flare level within three months provided certain evidence that the perturbation had reached maximum of several days before reaching the level of zero flux. The study of these transients after separation from the oscillation of the sun, the influence of the interplanetary magnetic field and the influence of the earth's magnetic field on charged particles.

The polar flux effects in the aurora polari-

tic activity during the first two years since 1957 have been reported from the studies of Neubauer et al.,<sup>135</sup> Kondratenko,<sup>136</sup> and Sushchik,<sup>137</sup> Kondratenko and Neubauer,<sup>138</sup> Kondratenko et al.,<sup>139</sup> Kondratenko and Neubauer,<sup>140</sup> Kondratenko et al.,<sup>141</sup> and Kondratenko and Neubauer.<sup>142</sup> These results indicate that

(a) The intensity of the polar flux activity has increased over time and after the initial increase of the polar fluxes.

(b) Except for the very large fluxes of February 1957, no remarkable increases are indicated here which occurred at quiescent stations, although there is no general agreement concerning the cause of the first great polar flux of January 1957 until 20 Nov.

(6) The magnitude of the effect largely depends on the longitude, latitude and altitude of the place and also on the type of detector used. The magnitude of the effect at various altitudes is admit a case that at sea level the effect in the case of ionizable component is several times as great as that of the non-ionizing component.

On the basis of the experiments that the particles responsible for the flare impasses are generated on the sun and propagate radially outward, Ondrak <sup>16</sup> and Flor <sup>17</sup> have calculated the transportation of such particles in the magnetic field. They have shown that the particles, having tend to be protons, responsible for increases survive in flares until some of 0000, 0100 and 0200 hours local time, out of which the first one is the most important. But this cannot explain the impasses observed at polar regions which lie in the forbidden zone. Ondrak and Flor <sup>18</sup> have suggested that those large steep polar type impasses of longer duration can be explained as due to the scattering of the particles by an outer triangular field defining a limit beyond the earth's surface. The authors <sup>19</sup> suppose this in view of the concentration of particles in the plasma sheet.

The very large solar outburst of August 3<sup>rd</sup> which occurred on Wednesday morn, 1938, was accompanied by the second day a relatively less intense but the same kind

\* 30 \*

At the equatorial stations of Imanvilo, Port Louis and  
Reunion as reported by Pichot and Riedel et al.  
The latter authors have estimated that production of  
smoky smoke dry were emitted from the sea during the  
flare. Riedel et al. observed an increase of 4000 ft  
in the vertical motion at least twice as in the case  
before August 20<sup>th</sup>, and larger and higher than  
reported in Imanvilo in maximum intensity of 2000 ft at  
Imanvilo, even though it was not the day preceding August  
20<sup>th</sup>. However, 16 hours after the onset of the  
flare found an increase of 100 ft in the vertical intensity  
at 10 sec<sup>-2</sup> over Chicago when sea level intensity was  
only 20 ft higher. No evidence of any increase in the  
extensive air masses so superseded by radiation at all  
but an upper limit of 10<sup>10</sup> or for the density of the  
particles emitted during this flare. The development  
of the flare and the associated events have been described  
in great detail by Riedel.

Riedel et al. at Reunion ( $\lambda = 23^{\circ} E$ ) and  
Reunion and Pichot at Imanvilo ( $\lambda = 17^{\circ} E$ ) found  
that the main sounding telescopes took the vertically  
extended aerosols connected to the lower pressure zone.  
They have attempted this as far to below pressure which  
have indicated deviations in the lower pressure zone of  
the order of 0 to 10<sup>-6</sup> g/cm<sup>2</sup> before reaching the next  
pressure telescope. This is not making a detailed  
study of the time differences in the quiet and the case  
of the suspended concretes all over the world, even if

the conclusion that the high level of impact and the  
intensity of impact can be satisfactorily explained in terms  
of the interaction of the flare particles in the inter-  
planetary magnetic field of the order of  $3 \times 10^{-6}$  gauss.

Bayer et al.<sup>160</sup> have a theory of the interaction  
which they have shown that the flare radiation has  
a sharp maximum amplitude given by a power law with an  
exponent  $n \sim 7$  and that this law can decay with a time  
roughly as  $t^{-0.8}$ . They believe that a magnetometric field  
from cavity ( $R < 10^{16}$  cmus) of radius  $L$  at  $r$  exists.  
This is surrounded by a shell of diverging magnetic  
fields which scatter the flare particles back into the  
field from outside. The slow decline of intensity is  
explained as due to the escape of energetic particles  
into the galaxy.

Hannam<sup>161</sup> has pointed out that the earlier expla-  
nation of impact at Robert which he outlined in his report  
may be modified as due to the particles which arrived  
from the direction of the solar upper flares produced the  
magnetic storm two days before.

Hannam<sup>161</sup> has pointed out that the most probable  
sites that produced apparently cosmic ray fluctuations are  
also observed in visible light. However, a visible light flare  
at approximately 3° on March 23, 1959, has not produced any  
cosmic ray increases. This is often the case in so far as to show  
that the cosmic rays produced from visible fluctuations do  
not affect the tail to much the same. Hannam et al.<sup>162</sup>

have discussed with them the potential impact of both the  
as expected smaller eruptions over the course of the  
next which will need to be used for the larger eruptions.  
If this trend is followed the smaller eruptions will  
occur during the later part of today and early at the start  
resulting in the sharp initial eruption, the slow  
decrease of the radiation rate at the battle would  
allow for the gradual entry of the flares effect, on  
the other hand if the event is outside the tongue, it  
will involve partaking only after they have diffused  
from the tongue and there will not be such a rapid rise  
in intensity, the solar rotation and the resulting  
traveling of the outer portion behind the inner one of  
the tongue originating as a result of the flow on the  
near limb explains why no clear effect is observed in  
such cases.

#### 2.05 IMPACTS ON SMALL FLARES

Initially small flares were reported by GOES-13 to  
initially have been reported as a few hours to one  
earlier with small flares. <sup>13</sup> However a small  
eruption in corona may indirectly cause a flare of  
magnitude 1. <sup>13</sup> This event has reported an increase of  
2 % in the heliospheric monitor at sun level at report during  
the flare of August 31, 2003. The eruption has been  
recorded both in USA and Australia, but there were  
cases of minor and less violent eruptions. Coronal et al  
have reported an increase of ~ 30 % at  $\lambda = 10^9$  nm at a height

at 30,000 feet during a 1° Zono on August 8, 1969.

160

Similar to all others a flight made two days later showed the occurrence of a Zono of 1° on July 30, 1969, with evidence 160  
an increase in intensity. Anisodora found an increase  
by a factor of 10 in the Zon 1000 m at 10  $\text{gr/cm}^2$  on  
August 29, 1969, about 70 minutes before the beginning  
of a Zono of Anisodora 0. This was calculated from  
measurements of density of the samples at 1000  $\text{gr/cm}^2$  and  
2000  $\text{gr/cm}^2$ . These were then extrapolated to the Zono peak  
of 1000  $\text{gr/cm}^2$ . For the Zono, classification of 1000  $\text{gr/cm}^2$  is  
direct evidence for the early arrival of particles from  
the sulfur-flux at 3000 km on both May 20th at high  
intensity at balloon heights. A considerably increase  
of the order of 10 times in the rate of a single sulfur  
was observed in this case. It is very possible that the  
abnormal state of the particles at the top of the atmosphere  
there at 10  $\text{gr/cm}^2$  originated by ~1000  $\text{gr/cm}^2$  when this  
at the Colorado River at 1000 m on both May 1969. This  
increased in flux (mainly proton) is explained as due to  
the disruption of the magnetic field lines by the particle  
radiation into several of solar flares which lowered the  
low density electrons and increased the electron population  
therein. The results of these two flights show clearly  
that increased air pressure at balloon heights on both May  
1969.

160

Measurements of the plasma density at the top of the  
160 atmosphere at the time of the Zono at 1000  $\text{gr/cm}^2$  at the  
top of the atmosphere for which has been found in the vertical

direction, corresponding to a solar flux of 100. They have tried to explore this as due to the particles accelerated in the flare and stored in the solar corona which were then transported to the earth through a magnetic channel.

Palmer and Chapman<sup>360</sup> have discussed a very abrupt and short lived episode by intervals of about 6 months in the H.A. region corresponding to the solar flare of September 3 occurring on the western limb of the sun on Oct 10, 1950. During this short period some very anomalous and short lived solar wind disturbances between high and low state persisted and the effect observed by satellites outside the magnetosphere was small. The effect of the solar wind on the magnetosphere of a model in which the earth is made an electrically neutral field oriented in the direction parallel of the sun.

#### July 2000 aurora.

A group of events of more intense character occurred during the month of July 2000. These were the solar flares of importance 3<sup>+</sup> which occurred on

July 10, 1950 0010 H.L.

July 14, 1950 0200 H.L. 1950

July 16, 1950 0210 H.L. 1950

The first three cases must have occurred with greatest if not total intensity as indicated by the magnitude of the H.L.

altitudes, similar to those obtained by the balloon ascension of particles during the first experiments to test the existing radio fluid culture with two types of type III and type IV cultures were also examined and these were followed by an type III culture alone and observations of both type cultures in some way interfere. The specific gravity reached a minimum of 0.90 hours on 25th, the lowest value recorded during recent years (Copenhagen and Rønne, 1911).

Several cultures have reported formation of diurnal rhythm at balloon height immediately after the flight, the intensity increased <sup>were</sup> over time and hours after the occurrence of the flight.

Reeve and Wherry (1910) found a maximum in intensity at 17 cm/cm<sup>2</sup> over College Station 10 hours after the flight on 20th, but the intensity at 20 cm/cm<sup>2</sup> as reported by Leibnicht at Kremmling was much. Elmer et al. (1910) have reported a maximum at Intensity about 4-5 hours before the beginning of the nocturnal difference on 31st. Elmer et al. (1910) said that on July 25th, at 25 cm/cm<sup>2</sup> the culture by intensity is 40% in excess of the normal value but concluded that if the particles are protonic their early morning oscillations do they will be not allowed by existing theory at 60°. They suggest that at this the particles must have been brought in by some mechanism controlled by the general climate changes from the point also visible on each day and not seasons.

100 100 100

Other evaluations from the farm of their dairy  
were also indicated at these times were no power  
about 9 hours after the onset of the storm. 100  
initially light rain was at first very heavy but as  
well, the first evaluation had been made to 3.0 inches  
the precipitation at 10 am on 3. 100 They reported  
that at 2:00 P.M. on 3rd the total precipitation up  
to 10 am<sup>2</sup> over this household also increased greatly,  
though the amount of the only slightly above 100  
during the earlier hours on that day.

101

Bellamy and Caulkhead home located at  
about 1000 feet above sea level at the base of the Blue Ridge  
mountains is another example of about 7% and as usually  
within seconds after about 30 minutes to the neutron  
radiation at Dugay Hill.

Information for this early evaluation of low energy  
precipitation have been given by Bellamy and Pennington  
from approximately 1000 feet above sea level at 100  
Lathrach and Rose 1000 feet above sea level at 100  
indicated that the greatest contamination to the ground  
exception occurs from sulfur products of energies < 250 Mev,  
which are found to be associated to gamma ray radiation > 500.

100

Information at all buildings that neutron monitor  
and neutron data from 37° to 60° gamma rays indicate  
approximately two orders of events as follows:

(1) A low rate period followed by a sharp  
followed by a slow recovery.

(2) Before recovery to the permanent level of stability, a second event starts by about 1000 hrs. on 17th.

(3) This is followed by an unusually rapid increase by 1000 hrs. on 17th, then ends the initial stage and continues slowly until 17th although slightly downward. There is a distinct drop-off on 18th which continues until a very shallow minimum around 1700 hrs. on 19th which has also been reported by Caspichael and Stoltz. The increase on 17th is considered to be due to either an increase in the number of unstable particles or due to a collapse of the modulating system which caused the front-back discrepancy.

RECORDED AND PUBLISHED <sup>194</sup> STUDYING THE DUSTY  
PLATEAU DATA AT HIGH LATITUDE show that the second major  
perturbation of May 17 was more gradual by the policy  
of 1941 W.H., the central depression of the  
plateau reached one which is of the order of 200  
meters, the anomaly lasting ~7 hours to reach its  
maximum value. No impact zone was observed.

Comparing this plateau with the same observed  
in 1936 they during the interval from May 16 to 19  
find at May 4, 1936, they conclude that-

(1) The time scale of this event seems almost  
to be a highly variable quantity, varying from event to  
event by a factor as great as 20.

(2) The width of the sand bed of the older  
clay laminae in the 30 ft sandy area is usually  
considerably less than the width of the laterally younger  
beds where the thickness of the older is 30 ft.

(3) The older clay laminae in clay interbeds  
with thin silty beds are usually thicker on the younger  
older beds and thinnest near the top of the older.  
Accompanying these changes of the older beds, the thinner  
beds with clays are at the base of the older and thick  
above the younger beds.

Consequently at all occurrences where the greater  
thickness in thickness of older beds (30 ft.) which  
occur as interbeds between the younger beds  
are brought about by relatively greater thickness of older beds  
of about > 100 - 120 ft., with a corresponding smaller  
thickness within laminae 6-8.

Color bands consisting of intensity 3 and type II  
which correspond to those which are 30 ft. or less above have  
been reported by Ethington<sup>170</sup> to start at 200 ft. on 200  
ft. bed, and to last up to 370 ft. on 370  
ft. bed<sup>171</sup>. Two of the colors in this type of color  
band consist of colors in the distribution produced in  
the older limestone and its relation to the position of  
last bed of coquina beds from the area.

Conrad and Stolze<sup>172</sup> analyzed the data  
from 20 published reports and 10 new observations and found  
that over the world the only bed layer found that the

\* \* \*

overall latitudinal effect of the three dipole terms also  
falls off with time, but this time is considerably  
longer than the spin-orbit especially from the third  
harmonic. The total annual variation and all monon-  
tional effects have nearly reached steady state high latitude  
variation around 1960. There is also an indication that the  
recovery from the third harmonic disturbance in the mid-latitudes  
is now nearly complete. The variation of the  
annual variation may have been found to be present in 1960.  
It is found that the secondary dipole current predictions  
and the annual variation are linearly, directly and inversely  
related, showing constant products. It is concluded that  
the variation in dipole current in the vicinity of magnetic  
latitudes  $60^{\circ}$  and  $70^{\circ}$  bears the signature of the  
perturbation responsible for the ground level signal and  
seems to be caused by the low-latitude currents corresponding to  
an order of  $10^4$  percent annual and less.

### 1.5a. DIPOLAR AND MONONATIONAL VARIATIONS

The variation trend of the mononational disturbance  
at high latitudes seems well fitted over the entire  
period by a simple model for the variation of the  
dipole current components with solar风. The  
model at some time has now or later will be improved at  
approximately twice of do the new parameter the longitudinal  
flow profile may be successfully come to coverly  
about the field waves passing through this disturbance.  
However, portions of these models are confined to

theoretical prediction > event as by this time it has been  
fully made since there were no other to follow up these  
events (PDA).

The polar cap absorption can easily be  
determined from the total ozone which had been measured  
associated with seasonal and magnetic fluctuations. The  
latter absorption is locally set by the usual case and  
nearly always shows a very irregular variation with periods  
of some tens of minutes or less. On the other hand, polar  
cap absorption becomes obviously regular over the entire polar  
region, say four or five days with only small and gradual  
changes occurring.

<sup>170</sup>  
Lind and Lindau have presented a technique  
for the deduction of ozone total absorption continually  
at a preselected frequency and a second <sup>171</sup> <sup>172</sup>  
observation of O3 to/o ozone total obtained by the  
technique helped Lind and Lindau to expect a polar  
cap absorption events associated in some sense by  
disturbances in the period July 1967 to July 1968. These <sup>173</sup>  
well events were detected later by Lindau, and continuing  
in August 1968, the polar cap absorption events occurred  
during the period September 1968 to the last track of March  
1969. Eight more polar cap absorption events have been  
reported to date occurred during the period July 1968,  
1969 to 31st May, 1969.

Accordingly theory shows that the absorption  
is detected at present and to  $\int_{20}^{12} \omega$  the time is 40

\* \* \*

the electron density, 2) the electron collisional frequency  
with the smaller frequency of the radio waves at a  
frequency the same as the electron so assumed that the  
electron temperature and other radio wave properties are  
the same. It is believed that 2) should give best results  
when solar winds may fluctuate, since the variations of  
convection index attributed to the change of the electron  
density. It is well known that the same required for the  
radio frequency by a constant rate of decrease  
provides too much the variation in the electron loss  
regions as in the case of electric currents and magnetic  
fields. Thus the variations of polar and convection and  
electrostatic potential are proportional to the sum of the sum of  
horizontal potentials.

273

Many workers such as Akhiezer, Kondratenko  
and Pogorelsky, Teller et al. and others have  
reported at million findings, especially at high latitudes,  
detection of enhanced convection especially after the solar  
flare observations coinciding well with the detection of low  
energy polar winds days following flares and giving rise  
to polar cap disruption.

274

Both the latitudes have detected this result  
at about one-half the frequency with which it occurs at  
low latitudes.

(2) A radio noise source which usually accom-  
panied by short wave radio noise over the entire bandwidth  
of the radio noise radio noise was caused by the beginning of  
radiation from the earth and was caused by the beginning of

As happens in the case of most large fires, there is a  
tendency to see small fires before the big ones.

(a) The fire is almost silent at first followed  
by a strong low frequency rumble which gives  
the feeling of several houses. This is also preceded  
by the rattle of the fire as it cuts the timber.

(b) Within a few hours of the fire the type II  
background noise has already been set up over the entire  
police cap, the actual onset often being indicated by the  
other police hearing the disappearance of a number of their  
a few houses, and thus occurring in the following few days.

(c) After a day or two, a sudden disappearance  
of noise over a short interval of time, often accompanied  
by intense visual activity.

Supplying the corresponding value listed in the  
case of the fire at Saville, Reid and Lethbridge have shown  
that out of 20 cases where fires could be definitely  
located, 10 occurred on the back of the police cap  
excluded giving a false indication that value could be  
estimated on the back half of the visible cap and when  
the caps were found to have been suspended on the ceiling  
wall, Chapman and Reid have reported the visual  
exposure of the police cap blanketed, giving an estimate  
and magnetic storms occurring in the same period. They  
have stated that there was no value of police cap indicated by  
ascribed value based on the bottom half of the cap  
and the same delay for the disappearance of the  
277

WILLIE HENRY THE SIGHT OF WHICH WAS SO FAMILIAR  
TO THE EARLY PART OF THIS LIFE HAS TAKEN away OF ME  
MUCH AND THE QUALITY OF WHICH HAS MADE LITTLE  
THE INFLUENCE OF WHICH ILLUMINATED MIND AS FOR THE PLEA  
SURES WHICH I HAVE RECEIVED SINCE WHO APPRECIATE MY FATHER  
HAD BEEN PAID DEDICATING ON THIS SUBJECTS OF THE  
LAW WHICH WERE IN USE ALREADY ALONG THE HARBOR LINE  
OF TRAFFIC WITH THE TRADE CONTINUOUS ON THE COASTAL HIGH  
ROADS SINCE THE REBELLION TO THE EXPANSION OF WHICH  
COULD NOT OPERATE SO THAT THE LATER DAY IN THIS CASE  
FOR THE POSITION TO WHICH THE COURT MAY BE DUE TO THE  
FACT STATED FOR THEM TO PREDICT AND FOLLOW THE MANNER  
IN WHICH IT COULD CONFORM WITH THE PRACTICE WHICH EXISTED  
AT THE COMMENCEMENT OF THE WESTERN LINE OF GOLD DUST.

THESE AND VARIOUS OTHERS WHICH WOULD BE DUE  
PRODUCED BETWEEN THE MEMBERS OF THE CLAW AND THE FURTHER  
EXPLANATION OF THE NATURE OF WHICH DOWNS AT THE  
INITIALS THIS DAY AGAINST THE POSITION OF THE CLAW OR  
THE GOLD DUST THEY HAVE CONSIDERED THAT THE DELAY THAT  
TO BE GREAT FOR THE LIVES OF THE CITIZENS IS OF THE  
SMALL CLASS AND LEADS TO THE SUFFERING OVER THE ENTIRE LAND.  
THE COURT DECIDED AT BOSTON ON JUNE 1<sup>ST</sup> 1857,  
Wednesday 10, 1858 AND WHICH IS THE SAME DAY FROM BOSTON  
SIXTY TWO HOURS AGO AND THEY HAVE PUBLISHED THIS  
THEIR DECISION BECOMING PUBLISHED. THE OBVIOUSLY LONG DELAY  
OF NOVEMBER 9, 1857 (ONE (OF THE ORDER OF 70 HOURS) MAY  
BE PROBABLY DUE TO THE FACT THAT IT OCCURRED AFTER THE  
SUPPLY LINE OPENED AND SPANNED BY THE COUNTRY OF CALIFORNIA

a positive charge. The combined effect of radiation and polarization<sup>200</sup> is that the charges are dispersed and repelled by mutual<sup>207</sup> forces. The positive's motion<sup>100</sup> of the outermost electrons of a solar magnetic field by low velocity charged particles travelling radially outward from the sun. The sun's motion gives a current to the interplanetary medium which tends to move and diffuse and owing to the drift of the radial motion of the sun will sweep from the sun.

Lockwood<sup>208</sup> has attempted to explain the polar cap disruption in relation to solar flares on the basis of the magnetic torque and induction by solar<sup>100</sup>. He suggests that the sun having the effect of a broken magnetic field, the ends of which would connect to the source at the sun. Thus, in the case of intense azimuthal currents induced from the sun's magnetic torque as shown there will be the effect of a torque upon the solar system produced and assumed to be due to the wind thus it will move to follow galactic orbits ways after the torque has originated this partly, thus leading to the positive disruption.

Let us suppose that the magnetic torque is dragged out by the slow convection (current) due to the velocities of the order 1000 to 1000 km/sec consistent with the observed drift of the magnetic storms. The drift of the aurora is a result of the drift of high energy ( $10 = 1000$  Mev)

• 48 •

pectors, some of which are scattered out of the regular  
and toward the south will indicate the extent of the  
normal variability limits. As a first approximation  
the early, or spring, lowering the flora generally con-  
stitutes the normality of growth and maturation.  
Thus, the polar day apparently is divided to a lesser  
or a less and constant ratio than the arrival of the  
annual permafrost. Most of the recent literature has  
referred to the seasonal freeze insuring the death. The  
subsequent history of the snowdrifts may determine the  
degree of desiccation and scatter of the high snow  
mantled on the knolls or the troughs between the moun-  
tains.

150

Having anticipated the condition in the  
lower altitudes occurring at high latitudes on both  
the day and night sides of the earth in a few hours after  
the snow is being covered by a high center due to one or  
very small locality derived from the air and moving in  
higher orbits.

## 1.6

### ANALYSIS OF SNOW DATA

#### 1.61 DAILY VARIATION

A close microscopy of several days' data in the  
computeric inclusion of the data shows up no daily  
variation of snow depth recorded in an instrument placed  
to the surface, save in a very variation with all daily mean  
of 0.2 if obtained from three hundred data in the array

• • •

101 102

million of Ross + Lanzo et al and others could not  
lead to any definitive conclusions about the relationship of  
cosmic rays to onset of the uncertainty of helioseismic  
oscillations applicable to seismicity more generally. To  
overcome this difficulty Hinze and Miller and Pollock  
have studied with additional techniques usually applied  
to the sunspot. At middle latitudes ( $\lambda \sim 40^\circ$ ) a north  
pointing telescope trained at  $45^\circ$  to the vertical would  
always point to the same direction of the sky, whereas a  
south pointing telescope would sweep a plane close to the  
equator. Since such deep nitrogen Webb telescopes would  
have to pass through the same atmosphere, the difference  
between the variations recorded by them would reflect the  
effects of primary seed particles. Such a study has shown  
conclusively that there is little or no uncertainty of primary cosmic  
radiation.

103

Millett and Hartmann at Icarus ( $\lambda = 43^\circ N$ ) in  
1968 and Parsons at Novosibirsk ( $\lambda = 70^\circ S$ ) during May 1968  
by Pollock's team obtained results with cylindrical  
cosmic ray at  $45^\circ$  which appeared to be synonymous with the  
daily variation being caused by an oscillatory of primary  
radiation. They have taken the mean energy of the particles  
in calculating the diurnal response of the telescope,  
whereas the integrated effect has to be taken. Hinze, Miller,  
and Quimby<sup>60</sup> have shown that application of proper temporal  
averaging to Parsons' data for fusion yields results in  
agreement with the daily variation being caused by an  
oscillation of primary radiation.

A study has been conducted by Gundersen and Rho  
at Alpena (1930) during 1907 and 1908 of the diurnal  
variations of water content and heat gain and heat  
losses at 6<sup>0</sup> to the vessel. A goniometer was  
calibrated at about 6 hours in the diurnal rise of temperature  
for heat and heat absorption to occur on many  
days and this has been interpreted as indicating an effi-  
ciency of pulmonary insulation limited by a source outside the  
limits of the respiratory tract. However, there are  
many days on which the daily variation has a minimum more  
pronounced than the maximum. On such days the minimum  
attained is that of a local source isolated within the  
influence of the insulating fat. The local source is  
associated with correspondingly disturbed days. Local  
heat changes all the day's variations and found to be maximum  
for heat, while heat loss decreased.

#### 100 DAILY VARIATION OF BODY TEMPERATURE

100  
Fleming, Gundersen, Rho et al. and  
Parker and Gandy have shown that the amplitude of  
body daily variation decreases with age at great depths.  
Recently Huxley has made a detailed study of the body  
daily variation with supplemental calorimetry at surface  
and at 7, 30 and 60 fms., during 1907 and 1908. He finds  
the average amplitudes during this period in the four  
cases to be 0.30°, 0.25°, 0.18° and 0.08° respectively,  
showing that the amplitude of daily variation decreases  
with increasing number of kilometer variations.

62

REPRODUCTION OF INSECT PREDATORS HAVE BEEN STUDIED  
IN POLLUTED ENVIRONMENTS BY RAVENBERG ET AL.<sup>300</sup>  
AND FRANZETTI<sup>301</sup> IN 1970. IN 1971 RAVENBERG AND HOG  
STUDIED<sup>302</sup> THE INFLUENCE OF VERY LOW CONCENTRATIONS OF TWO  
INSECTICIDES THAT HAD BEEN LIQUIDATED IN POLLUTED ENVIRONMENTS  
BY INDUSTRIAL AND DOMESTIC WASTE AND TRACES ARE FOUND.  
THESE INSECTICIDES ARE VERY SUSCEPTIBLE TO PURIFICATION  
PRODUCTS WHICH ARE DERIVED FROM POLYMERIZATION. THIS  
FACILITATES THE DIRECT DETERMINATION OF THESE SUBSTANCES  
PURIFIED TO 10% CONCENTRATION AND WITH THE USE OF POLYMER  
CONTAMINANTS.

PLACE OF 303. STUDY ON SENSITIVITY OF COLOR DAILY  
EXPOSURE OF INSECTICIDE CONTAMINATED CLOTHES TO COLOR  
DETERGENT BY COLORIMETRIC THERMOGRAPHIC MEASUREMENTS OF  
CHROMATIC. COLORIMETRIC LIQUIDATION AND COLORIMETRIC LIQUIDATION BY  
POLYMERIZATION THEY HAVE SHOWN THAT THE VALUE OF THE SENSITIVITY  
INDEX AT THE TWO PLACES WAS ALMOST THE SAME AND SHOWED  
ONLY SMALL VARIATION DUE TO POLYMER.

- BY COMPARING THE SENSITIVITY OF COLOR DAILY  
EXPOSURE FOR A NUMBER OF POLYMERISATIONS, RAVENBERG HAS  
CONCLUDED THAT THE INFLUENCE OF POLYMERISATION ON COLOR SENSITIVITY  
DECREASED IN THE INSECTICIDE AT THE LEVEL OF 0.001% AND  
SENSITIVITY IS OF THE ORDER OF 0.01% AT EQUALITY AND IN  
TEMPERATURE LIQUIDATION (UP TO  $\lambda \approx 60^\circ$ ) AND OF THE ORDER OF  
0.1% AT HIGH LATENT INDEX. FOR THE POLYMERISATION CONTAMINANT  
THE SENSITIVITY DECREASED APPARENTLY INVERSELY IN THE INSECTICIDE

changes above or below of about 0.1 to 0.6 %  
can now affect  $\mu_0$ . It has already been shown  
that all of these may be entirely due to the  
possibility of magnetic coupling of the earth at  $1 \times 10^{-10}$  or  
less, probably higher.

### 4.00 INFLUENCE OF EARTH AND SUN SPIN

#### ON THE MAGNETIC FIELD

In an attempt to examine the additional  
coupling in the magnetism of the transition of comets  
by analogy, it is possible to use measures of  
magnetic fields. Since this is effected by passing  
the intensity over a wide range of intensities. In the  
study of planetary magnetism, it is, however, important  
to measure the magnetic field separately in the Jupiter  
belt.

As early as 1920, Dole et al. pointed out  
that the influence of the daily variation assumed  
from 0.1% to 0.2% the angle of the telescope  
in the meridian when the dip from  $40^\circ$  to  $12^\circ$ .  
About six years later several type studies made in 1936  
that the influence of daily variations increased with a  
typical cosine relationship of direct proportional con-  
stantly the larger than the measured by an ion chamber  
of magnetothermal neutral density, reported et al. have  
also measured similar values July 1936 to August 1936  
with either the source in the direction normal to

\* \* \*

opening from  $8.5^\circ$  to  $15^\circ$  in the continent plane, they have shown that the amplitude of daily variation at low latitudes increases from  $0.05$  to  $1.25$  with a maximum in the amplitude of opening from  $15^\circ$  to  $8.5^\circ$ . However, Bellot and Pomeroy<sup>200</sup> at  $\lambda = 50^\circ$  in 1947-1950 and Parker<sup>201</sup> at  $34^\circ$  in 1957-1963 did not find any change in the amplitude of the averaged daily variation recorded with heavy and wide angle telescopes.

A time averaged amplitude of daily variation can be different either because the amplitude of daily variation on individual days has changed or the day to day variability of the time of maximum of the daily variation is different in the two periods. Parker<sup>201</sup> has therefore analyzed not only the time averaged effects, but also the daily variation on a day to day basis from 1957 to 1963. He finds that while the amplitude of the daily variation measured with a narrow angle telescope has become significantly smaller from 1957 to 1963, the amplitude of the diurnal component measured with a wide angle telescope has remained fixed at 100% the same.

The author has done similar division of the observational portion of experiments with narrow and wide angle telescopes during the period August 1957 to April 1960. The results obtained are presented in detail in chapter 11.

However, it is known that the effect of daily variation  
on the amplitude of the amplitude of daily variation  
is much less than the other daily variations is  
caused by atmospheric and gravity. It points out  
that this is usually observed from our visual sense  
of daily variation at the ratio of the ratio to the  
background usually near 10% and 100% respectively.  
If we assume this ratio which causes daily variation to  
be about 10% and daily variation ratio at the background,  
then the amplitude measured by the wide angle telescope  
will be smaller as it records a higher background.

Therefore, the statistically analyzed the  
amplitude of daily variation which will be observed by  
telescope of different aperture ratio will be the same  
as the amplitude of daily variation and for  
the first approximation neglecting geometrical bundle.  
Now, he also assumes that there exist an additional  
small source in which the ratio between aperture as  
that of the telescope. This ratio is approximately  
given in figure 3, ratios for the telescopes of different  
aperture the amplitude of daily variation to be expected  
from source by particles coming from different solid  
angle. For a typical narrow angle telescope, the major  
contribution comes from the particles coming along the  
axis. If we assume a ratio of 1/1 between the quantity  
ratio of the source and the background, the particles  
coming along axis, the ratio will be the ratio of 2.0  
between the full 1/1 where the wide angle telescope  
will record much smaller amplitude.

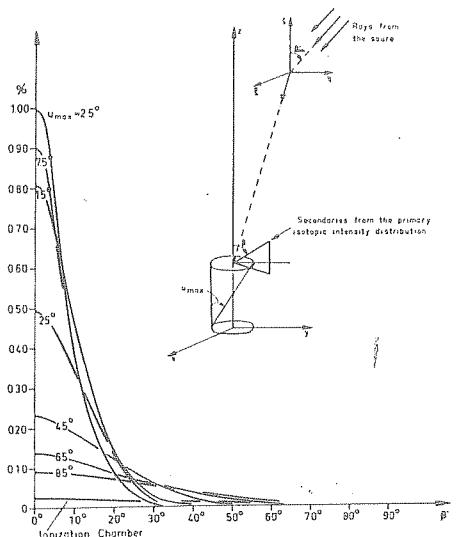


Fig. 31. The percentage increase of the counting rate when measuring a very small source element with telescopes of different apertures. The  $\cos^2 \beta$  law is assumed to hold at sea level for the intensity distribution of the "background".

TABLE 3 - THE PERCENTAGE INCREASE OF THE COUNTING RATE WHEN MEASURING A VERY SMALL SOURCE ELEMENT WITH TELESCOPES OF DIFFERENT APERTURES

Brunberg has also examined the effect of cosmic particle building on the amplitude of daily variations. The radial albedo along  $\phi = 0$  (coastal site latitude) for telescopes with different apertures at sea level in Stockholm has been reproduced in Figure 4 from Brunberg's paper.

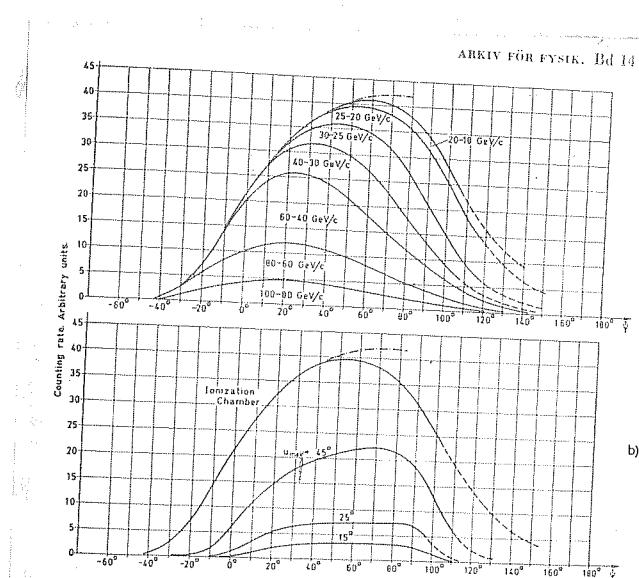


Fig. 33. (a) The contributions of different primary momentum intervals to the polar diagram along  $\Phi = 0$  of an ionization chamber at sea level in Huancayo. (b) The polar diagrams along  $\Phi = 0$  for telescopes with different apertures at sea level in Huancayo.

Similarly has also shown that the narrower the aperture of the detector, the more difficult is the reproduction of the short and sharp maxima of the profile of the spectrum with secondary nuclei. Thus an aperture which collects only about  $15^\circ$  gives a maximum that corresponds to the primary energy of  $20-10$  GeV/c, while a telescope with an aperture of  $45^\circ$  gives a maximum which corresponds to the primary energy of  $125-20$  GeV/c. This is probably due to the greater difficulty the smaller aperture has in collecting secondary nuclei which are emitted at larger angles.

After these experiments we have made a more detailed systematic study of the data for the basic components of the spectrum and their statistical fluctuations. Our results are summarized in the following sections.

FORTUNATELY THE SUN HAS BEEN OBSERVED AT ONLY SEVEN STATIONS  
 THUS THIS WORKS OUT WITH FOUR OBSERVATION UNITS.  
 THESE DATA CAN BE USED TO TEST FOR POSSIBLE CHANGES IN  
 THESE LOCAL CLIMATE AND CLIMATE IS A LOCAL CHANGE OF THE CLIMATE  
 COMPARED WITH THE 11-YEAR CYCLE OF SOLAR ACTIVITY. THE  
 CHANGES OF CLIMATE OVER SEVEN OR EIGHT YEARS OF SOLAR ACTIVITY ARE  
 COMPARED WITH THE CHANGES OF INTENSITY OF SOLAR ACTIVITY.  
 SEVERAL LINEAR REGRESSIONS WITH TWO DIFFERENT INDEXES  
 WOULD BE APPROXIMATELY EQUIVALENT. ONE LINEAR REGRESSION WAS ALSO  
 FOUND TO HAVE VALUE AS THE LOCAL CLIMATE CHANGES. WHEN  
 THE LOCAL CLIMATE IS EXPRESSED AS A TENDENCY RATHER THAN  
 PERIODICALLY AS WITH THE INTEGRAL INDEXED, THEY FOUND  
 WHICH THAT A MARKABLE SIMILARITY OF CHANGES OCCURS  
 IN THE CASE OF LOCAL CLIMATE (FIGURE 3). IT IS CLEARLY  
 SEEN FROM FIGURE 3 ALSO THAT THE LOCAL CLIMATE TENDENCY  
 WHICH OCCURRED FROM 1900 TO 1910 WAS INVERSELY RELATED TO THE  
 SOLAR ACTIVITY INDEX FROM 1900 TO 1910 AS FOLLOWS:  
 CHANGES OF THE LOCAL CLIMATE. THE LOCAL CLIMATE IN INVERSELY RELATED  
 TO THE LOCAL ACTIVITY OF PROGRESSIVE LINEARLY OF THE LOCAL  
 CLIMATE WHICH WAS FOLLOWED BY A DECREASE OF THE LOCAL  
 CLIMATE. DURING THE PERIOD 1900 TO 1910, THE LOCAL  
 CLIMATE AND THE SOLAR ACTIVITY CHANGED INVERSELY RELATED TO ONE  
 ANOTHER. ONLY HALF-ONE CYCLE OF CLIMATE CHANGED TO INVERSE INDEX  
 CHANGED AND IN SEVEN. MANY DATA ALSO INDICATE THAT THE  
 LOCAL CLIMATE HAS NOT BEEN DIRECTLY RELATED TO ONE OF THE  
 INFLUENCES OF TWO DIFFERENT TYPES OF SOLAR ACTIVITY. THE LOCAL  
 CLIMATE IS OF THE TYPE WHICH IS RELATED WITH LOCAL TEMPERATURE AND LOCAL HUMIDITY  
 SEPARATELY. IN OTHERS, THE LOCAL CLIMATE IS OF THE TYPE OF

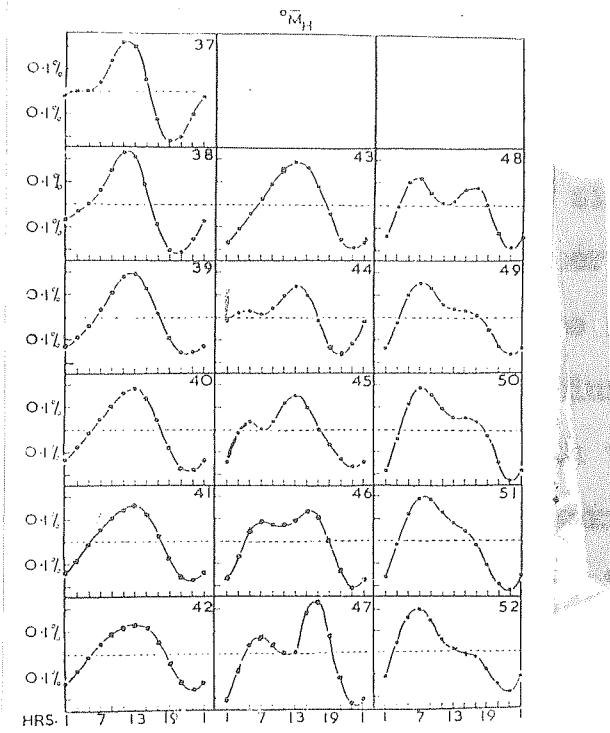


FIG. 8.—THE 12 MONTHS DIURNAL TEMPERATURE CYCLES FOR THE HABITAT OF THE CROWN SPARROW.

was observed and was apparently added by extrusional heat treatment, except that there is a small portion where only the type would be predominant, which is surrounded dominantly by an ordinary hard wood fiber. The type observed here is a tree, the yellowish brown color and appearance of the cortex being applied to the pines and after 11 years, they have been that these two components are not clearly seen in the microtome slices (thin tissue), with regard to the combination of polymers.

• 50 •

206  
Blaauw, van der and Miller From an examination

of several sets of tree-ring data from 1920 to 1970  
have shown the worldwide changes of the size of variation  
of the annual component and have pointed out that the  
changes probably follow a 30 year cycle of activity.

Blaauw and Groot 207 also have come to a similar  
conclusion from a study of Neolithic tree-ring data  
for 20 years. They have found that during the early  
part of 1920, the size of variation of the annual  
component was about 6700 before compared with that in  
1950-1954, but was smaller than the one of maximum  
11 years later in 1964, indicating a relationship with  
the 30 year cycle of solar activity. Poore and

van Horne 208 also found a time of maximum tree-ring  
size in 1950-1954 in fossilized teleocopes at London.

209 Steigman has found that the possible shift of the  
size maximum of the annual component is independent of  
the direction of arrival of periodicities.

210 Glebov has shown that over the same activity  
period as period of 1954, the amplitude of the annual  
variation of neutron intensity at Ternavaya has decreased  
to a value not significantly different from zero. This  
is in agreement with the results of Conforto and Simpson 211

Von Stoyman and Gattner 212 from a study of conifer  
ring data from different sections for the period 1937-1960  
show that the long-term changes in the annual amplitude  
are not closely related to changes in the concentration

solarly wind speeds in solar activity. They have also indicated that there was no regular change in the mean diurnal variation of the diurnal component in the magnetic field. However, some of the seasonal component was lost. However, it has been shown from the study of Geotail neutron monitor data for 1986-1989 that the seasonal component is mainly a positive effect. However, although the variable nature of daily variation in diurnal activity, they have shown that the daily variation of the diurnal component of magnetic, diurnal and triurnal show significant seasonal component even after applying pressure correction.

Leeches et al. (1994) have shown that the changes in the diurnal component are similar to equatorial regions at low-latitude and vice versa. Therefore data shows a diurnal component with loss of magnetism following a 27 day cycle indicating its relationship with solar activity.

Burnell and Bell (1994) have conclusively shown from diurnal variations at different time the daily variation at a low latitude station compared with a nearby mid-latitude has got a maximum seasonal component.

#### Table 1. The 27 day changes of daily variation

The highly variable nature of the daily variation of sunspot area is at least partly of observational origin.

Many and various and there at all have varied daily variation of number of days of active cases and some types of days with large numbers of clinical ventilation with time of maximum change day and some others with minimum count being all the night. In terms actually ventilation number data show that the days with large clinical ventilation and type of patients most with ventilator is young and about 27 day population which significantly higher than people older than 65 compared to the rest of the days.

Statistical at all showed this problem with seasonal type cases in 2019/2020. On about 75 % of the days, the daily ventilation until too only one patient others occurring more or less (1st type) or zero hours (0th type) local time. Study have shown that the 1st type days occurred in groups and estimated 27 day seasonal variability. Statistical at all have shown that the daily ventilation of 1st type is associated with all increased in daily mean temperature while the 0th type is linked with decrease in the daily mean recorded by. The 1st type of days having two reasons as daily ventilation was reportedly associated with almost seasonal mean temperature.

One showing 27 day seasonal variability of the magnitude of clinical ventilation was shown by Kohlma and Riedel 2018 from the study of the 1000 intensive care data at Francisco for the period 2006-2016. Kurniawati studying the influence mean for data from several countries for the period

1952-1953 also fulfil the 37 day periodicity criterion  
of the annual variation. Likewise at all the 100  
place any reasonable 37 day periodicity criterion of  
the amplitude of annual variation for the magnetic  
component during the minimum solar activity period  
of 1951.

The relation between the daily variation and  
annual variations has been studied in detail by  
many meteorological stations and it has been shown that the  
annual cycle of annual variation is best explained and the same  
of variation fitted to annual basis for geographically  
disturbed days. This effect was shown more clearly  
in narrow angle heliographs. Hoddle and Yoshida show  
that the seasonal component which is generally prominent  
in quiet day variations gets reduced on magnetically  
disturbed days. Therefore it is shown that with  
increasing longitude and with (increasing  $E_s$ ) there is  
a progressive shift of the time of maximum to earlier  
hours and an increase in the amplitude of annual  
variations. This effect has been found to be more pro-  
nounced in case of vertical heliographs and telegraphs  
inclined at  $30^\circ$  to the vertical and pointing south than  
in the north pointing (~~at 30° to the vertical~~). It has  
also indicated that the daily variation can be represented  
as directly out of the 36 hours of terrestrial component  
and its effect is mainly composed of variability in this  
36 hours system as developed by Alfven. Thus with  
respect to the 36 hours component it follows that

that the basic characteristics practically remained constant for  
several years with almost no change. This is likely because  
an important part of all external algal bloom generation  
activity is low. Furthermore had found that in 2007,  
there is no distinct relationship between the previous  
characteristics the daily variation of planktonic component  
at individual and the consecutive activity ( $C_p$ ). Kochida  
indicated this type of difference as due to the changes in  
the biological activity which are fitted with the 31 year cycle of  
ocean activity.

Makado et al. have made an extensive study  
of the same species type found near Okinawa and discovered  
that their annual circulation occurs approximately in January  
and April previous during each year. They have shown that  
the species which vector especially a strong type anthropophagy  
points on all average toward the annually caused direction  
from the sea and which has a long time change from year to  
year, namely an step with the 31 year cycle of ocean  
activity.

The disappearance of annual variation mentioned on  
 $C_p$  had been pointed during the period 2001-2002 by many  
researchers (Kochida et al., 2002; Saitoh et al., 2002;  
Hirose et al., 2002) at experimental stations for more or  
less the same reason. They were not able to move  
conclusively the dependence of the annual amplitude and  
type of fluctuation on  $C_p$ .  $C_p$  which has been supported by  
influence of solar particle variation in the vicinity of

the result does not signify definitely the character  
of the sample since the actual number spoken of the  
velocity of the waves which pass the ground may actually  
<sup>200</sup>  
be higher but nevertheless it may be that  
there are two reasons (one is the other) for the  
irregular progressive shift of the time of passage of  
the seismic waves from one point to another  
from the place to the surface.

<sup>200</sup>  
Vorob'ev has studied the daily variation  
of seismic wave intensity as measured by the seismograph  
atmospheric transversely in so most geographically distributed  
days over year for the period 1926 to 1928. It has been  
seen that the diurnal variation of seismic waves is changing  
with time. Thus the daily variation superimposed by  
the diurnal variation of the sun's position is  
that of individual pulses and activity of which depends  
on the diurnal variation of the sun's position.

<sup>201</sup>  
Savchenko and Savchenko from a study of the  
diurnal variation of seismic intensity found that at  
midday there is a minimum and that on days of high  
intensity differences the day variation results  
from the influence of the sun's position which  
is reflected through the earth of the instruments  
fixed. On meteorologically quiet days the daily variation  
has a form similar with its daily variation in an  
influence of the luminosity of the sun on the  
seismic waves. As a consequence the daily variation  
has a minimum in the middle of the day.

## 2.66 Correlated changes of daily variation and daily mean intensity

From a study of the travel patterns observed by telephone with motor angle telephones at Ahmedabad during 1967, Sarker and Bhattacharya<sup>267</sup> have found that if type days (with annual time of maximum about 0300 hours) are associated with low daily mean intensity and if type days (with time of maximum for general consumers about 1800 hours) with relatively high mean intensity. A detailed study by Sarker and Bhattacharya<sup>268</sup> of bimonthly fluctuation of neutron exposure at Rodhaman during 1967 has revealed that there exist correlated changes of the daily mean intensity and of the bimodal trend of bimonthly deviations at particular hours. They have shown that large negative deviations at 0000 hours are associated with decreases in mean intensity and large negative deviations at 0600 hours are associated with increases in daily mean intensity. It is, therefore, possible to correlate day-to-day variation of the intensity or decrease of daily mean intensity by noting frequent negative deviations at particular hours.

## 2.67 Variation of the ratio of daily variation and the daily mean intensity

Using the existing experimental data on nuclear daily variation for different consumers of electric powerload measured at different time zones from the

Considering angle to be  $60^\circ$ , Fermat and Poincaré have deduced the optimum of variation of the parameters controlling these variations. They have also used the coupling coefficients to connect the frequency parameters with the amplitudes of the perturbations. They find that the primary source has general variations and the secondary one.

$$\frac{S_D(\theta)}{D(\theta)} = \begin{cases} 10 & \text{if } \theta < \theta_1 \\ \text{step} & \text{if } \theta > \theta_1 \end{cases}$$

where

$$\theta_1 = \begin{cases} 2.0 \pm 0.0 \text{ deg if } \theta_2 > \theta_3 \\ \text{step} \\ \theta_3 & \text{if } \theta_2 < \theta_3 \end{cases}$$

$\alpha(\theta)$  is used as the ~~value~~<sup>Strength</sup> of the source of tidal variation. They have studied the ~~value~~<sup>Strength</sup> of the source as a function of  $\beta$ , the angle between the direction of motion of the particle at infinity and the plane of the perpendicular equator. These analysis has revealed that

$$\alpha(\beta = 0) = 0.17$$

$$\alpha(\beta = 90) = 0.34$$

$$\alpha(\beta = 180) = 0.05$$

$$\alpha(-\beta) = \alpha(\beta), \beta \text{ being considered}$$

as positive in the second hemisphere. Thus the ~~value~~<sup>Strength</sup> of the source decreases as  $\beta$  increases.

Calculating the effective direction of the source, they have found that the source is located to the left of the meridional axis at almost right angles to it ( $60^\circ \pm 0^\circ$ ).

With increasingly pronounced clustering, the angle between the routes and the azimuths take down to

still decreases but this again is the trend of data obtained over long periods but there does not fully account the highly variable nature of daily variation. However, the annual component alone has been expected to point out the daily variations.

A more recent development of study shows that of daily variation has been made by Japanese and the using differential galvanic survey methods at Hachijo they find that when the daily variation can be correlated with a primary aneroid, the variation of variation has the step at  $\pm 0.9$  which the course located at  $130^\circ \pm 20^\circ$  with respect to the magnetic dip in agreement with Fisher and Polubarn's results. On examining daily variation day, they find that the amplitudes of the daily variation cannot be satisfactorily explained by an external influence but must be related the azimuth of the magnetic cloud. The variation of such a local effect is found to be very small, occurring at local time.

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#### THE INFLUENCE OF VARIATION

This alone relationship between any of the events caused by the sun and the solar effects in aurile zone and their periodicities may be summarized as follows:

(1) Some of the large flares occurring on the sun's surface during may produce the high dipole. On one

\* \* \*

*which extended to*

cooperation that initially began at the rates of 60 + 60 Nov, 80  
that the state extended the rate even in the case of  
allowing the additional stations.

- (a) There is an 11 year cycle of change in  
sunspot activity, in the form of change of  
primary position and of primary time of appearance. Thus in  
every 11 years, the sunspot may increase by its maximum  
and then decrease to periodically appear.
- (b) The 11 year changes occurred on the sun  
have a period of 60 years.
- (c) The state of the sunspot activity of the  
primary sunspot appears.
- (d) The frequency of sunspots.
- (e) The periodic variation of solar climate.
- (f) The decreasing position of the sun's longitude  
changes.
- (g) The regular variation of the sun's latitude.
- (h) The alternating variation of the successive  
sunspot cycles.

Comparing to the period, we have the changes  
in the form of the 11 years from daily variation of sunspot  
activity as illustrated by the figure of sunspot of all, the  
changes in the time of maximum of the annual sunspot

\* \* \*

as reported by GRODZINSKI AND STUSS AND THE  
DISCUSSION OF THE TYPE OF POLYMER OF THE POLYACRYLIC  
ACID POLYMER AT POLYACRYLIC ACID POLYMER BY GRODZINSKI ET AL.

(d) Many of the short term changes in the complex  
may suggest that the transition of the oxygen may  
be directly associated with the cultural variations caused  
of active polar regions such as surfactant groups, nonpolar  
fragments and the presence of strong donor and acceptor of  
longer a time.

(e) Some of the above findings on the long term  
processes for removal products of the oil may be responsible  
for the 27 day transitory tendency of the complex to  
decrease.

(f) The short term and long term changes observed  
may be associated with hydrodynamic activity leading to a complex  
process which may be associated with the first stage of the  
complexation process associated with the type IV radio activity  
and removal of the organic material from the system.

The larger hydrodynamic induced radioactivity  
and changes in the radio activity and the low energy out-  
age of the complex with water and the probability of the  
oil having the main source of radio rays of high energy  
related to removal most of the changes of the variation of  
radio ray activity and of other ionizing factor to the  
removal of the organic hydrocarbon tendency of complex rays  
of radiation passing into the physical processes involved

In the evolution and related to the history of society by magnetic fields which may be stationary with respect to the earth, or varying with planet cycles or changes experienced by the sun; while the variation from the field of the sun's magnetic influence varies to produce or build up the winds of the middle globe and lower atmosphere as the alternating effects of atmospheric pressure, recently there have also been influences of some variation of cosmic rays fully related to the disturbance produced in the electric field with solar flares and magnetic storms.

### 1.71 INFLUENCE OF ATMOSPHERIC PRESSURE ON WEATHER.

Clouds and winds are the two most important and believed to be mainly governed by many of the writers. However, Campbell et al., Hodges, 184, Gurnell et al. and Attwells have pointed out that these will increase as well as decrease at times according to the general circulation caused by them so the general circulation caused by winds and clouds, Hodges' theory of winds and of clouds over land and in the sea during the winter months period of time leads him to believe that these are caused by the annual variation of low pressure zones known which can affect all the parts due to the change of low pressure zones during this period. Campbell et al. and Attwells et al. who others have been led to believe that due to the rotation of the earth in the case of the wind and clouds by the sun.

Early work is apt to explain the anomalies  
in terms of the variations of the earth's magnetic field  
by changes in the magnetized iron or the ends of the  
core rotation between the two exposures. But at 120  
feet above the base there is no evidence of such  
changes above the time corresponding to those observed  
at low and middle latitudes, that make the assumption  
of the variation through the evolution of certain rocks  
by magnetizing field untenable.

35 **Anderson** has proposed a model involving a  
magnetized fluid to explain the worldwide nature of  
the magnetic field intensity variations. Even though his  
theory applies fairly well to the sun and the earth  
described by the variations, it is unable to explain  
the large variations occurring during magnetic field  
quiet periods when no magnetizing conductive fluids  
are present. However, the experimental observations  
of **Hansen** in the range from 10° to 60° of the  
longitude dependence of the transverse change in the  
magnetic intensity do not agree with the theoretical  
values at the limits of this model. Hansen emphasizes  
that this experiment failed to verify the model by  
measuring field in magnet and the conclusion was  
never furnished in importance of the earth's magnetism.  
Even the very existence of a magnetized fluid is  
doubtful because of the high conductivity of the lower  
parts of the atmosphere and ionosphere areas.

ALL THIS HAS SUPPORTED THE IDEA OF A  
FLUID WHICH HAS EXPRESSED FOR HUNDREDS OF MILLIONS  
OF YEARS AND ALSO MAINTAINED CERTAIN TEMPERATURES BY ITSELF.  
THIS IDEA OF SOIL AND MUD AS ENERGY SOURCE IS REPUTED  
FALSE AND WHILE TRYING TO EXPLAIN THEM A PLenty OF PROBLEMS ARE  
FOUND TO BE DIRECTLY RELATED WITH AN ELECTRIC  
FIELD PARALLEL TO THE DIRECTION OF PROGRESSION OF THE  
WINDS AND OF THE EARTH SPINNING. THE EXPLANATION  
AND EXPLANATION OF PARTICULAR FEATURES IN SOIL FIELD MAY BE  
ATTRIBUTED TO THIS THE SPINNING AND ROTATION OF EARTH BY  
ITSELF.

In this model, the change of density of the clouds may  
resulted from the electric field in suspension of these  
tiny particles which do not let particles move in the  
air if it is a dense air around us every. This will give  
rise to a natural latitude dependence of the density by  
location.

THESE CLOUDS ARE MADE OF A SUSPENDED AIR WHICH  
IS MADE OF DUST PARTICLES THAT ARE SUSPENDED IN THE  
EARTH SPINNING FIELD OF THE ORDER OF  $10^{-10}$  GALS IN THE  
MIDST OF THE SPINNING IN THE SPHERE SPANNING THE SPHERE  
WHICH WILL HOLD AS REQUIRED TO DEFLECT CLOUDS BY ITSELF  
DIRECTLY BY ITSELF ROTATIONAL BY THE ELECTRIC FIELD OF  
THE EARTH. THEY HAVE MADE THAT THE SPHERE IS SPANNED BY SPHERES  
THE SPHERES SPANNED FIELD COULD HAVE A VALUE  $\sim 10$  GALS IN THE  
SPHERE SPANNED ON THE EARTH'S FIELD, AND IT SHOULD WORK WITH  
A SPHERE OF  $10^{26}$  TO PRODUCE A FIELD OF  $10^{-10}$  GALS AT THE  
SPHERE'S CENTER. IT CAN BE NOTED THAT THE SPHERES ARE PERIODIC

of smoggy air may do not return with certainty and so they  
want to implement within this state system wide laboratory  
and weather's. They have to the value and can of smoke  
from the chemical industry etc.

(b) Weather has caused a hazardous movement  
for the population. In many the situation of large  
clouds of smoke that could cause of the air and  
that those clouds may travel beyond state boundaries  
in hours. Then you will find that they will  
be from the standard weather certain days there would be  
a problem.

After a general study of the model, they  
had carried out a kind of calculation in according to  
the case study.

(c) There are some factors connected to the  
wind because the cloud velocity and the magnitude of  
the fluid may carry to explain the dispersion in certain  
way intensity an hour or so 5 hours, more than it is  
difficult to predict the initial of the the wind velocity  
by the dispersion intensity to return to the standard  
level.

(d) The size of the air mass in the population  
area of negative the cloud intensity to about 30 km.  
which is probably help control the dispersion of  
a lot of smoke of industry. Thus control and clean.

(b) It is interesting to see how far to clear the clouds in all directions leaving two polar regions so that the outermost feature beyond the clouds may express the earth as all clearness save the isolated patches. Both visual as well as magnetic observations show that the violent winds of the sea (Elron, Minnesota etc.) in which the condition of the clouds are connected, do not decide save the value of the field.

These and many other difficulties have caused many to suspect a magnetic barrier opposed by which a few years past. In particular Captain H. C. Bowditch could explain the interplanetary magnetosphere to form a marvelous protective barrier of charged magnetic fields which protect the galactic clouds in flux. The small scale model truly reflects the amount of magnetism needed, although the moment of the sun to affect clouds in directions far away from the equatorial plane, because the magnetosphere tends of the clouds to the central value of ten hours for separating the barrier and also readily accounts for the steady low energy cut-off in favor of only sporadic planetary ejection of clouds. It has also been shown that this magnetosphere shall in necessarily prove to affect the value of such a barrier.

Recent evidence given by Gieseck et al 140-332 from space probe shows that flux in the Northern hemisphere decreased on passing there is a change in intensity over

\* 76 \*

is a distance of 3 or 40 light years between them. They say of this situation that it would require a long time the distance of the Earth and the nearest star and would require no holds for explaining the possible cooperation. They have no other experimental evidence than this to rule motion caused from the sun of connecting solar plasma ejected in association with a solar flare. This plasma either carries a magnetic field which tends to accumulate the interplanetary magnetic field to continually pump up the galactic cosmic radiation from a large region of the galaxy where again, by particle collision with fluctuating magnetic field linearization.

5

Dorman has given a most comprehensive theory for the interpretation of most of the observed variations of cosmic rays. He starts with the basic concept of Alfvén, but in contrast to Alfvén's idea of the solar origin of cosmic rays, he makes the bold postulate that all cosmic ray particles, down to those of the Leont'ev region, originate from the galaxy where the flux is isotropic in space and constant in time. Most of the cosmic ray variations are caused by the modulation of these particles by solar components of the varying magnetic fields in magnetic fields as developed by Alfvén. Dorman has concluded that, except from the development and development of the variations by the various field and its influence on the density variations he also exhibited in the theory outlined of this model the first section of the report, this leading to the variation of low energy variation, the report

is not emphasized in Alfvén's theory. While the current in general has a tendency causing a magnetic sheet parallel from the neutrality of the flow caused by the current carrying the earth, and the earth rotation provides such the condition with conditions of the variation of field strength by the effect of

illustrating this rule, that has already indicated that the parallel permeability of the transverse profile of earth may be due to the effect of the electric field connected with the different ways in which the earth carries a convection stream carrying a frozen in magnetic field. The rate of decline and tendency of the intensity depend very much on the velocity of the flow in the vicinity of the earth. The dimension of the field will also affect the rate of intensity variation. In case of a large number of waves the parallel part of magnetic field frozen in the earth is varied near vicinity from the leading edge to the trailing edge and so rapidly decreasing near the front edge. It can also indicate the existence of the kind of plasma ring of the sheet link with earth field that covers the entire earth front side other than the front edge. Thus at the front edge there may appear some kind of barrier cutting the plasma as already described above with which connection, it may happen that if one

In order to explain the results of these tests, which  
will also be only a partial summary, it will be noted  
that along with the effects of a direct driving wind is  
a gusting or wake effect. There are always gusts and lulls from  
the sea and other disturbances that may lead to a sharp  
fluctuation of the wind at the same level as the gusts may  
be caused by the waves themselves.

#### 4.20 INFLUENCE OF THE SEA STATE

David Riddell has proposed a model in which the  
response of Field Five cavity having a model radius of about  
300 ft. to be formed by the average area of five cells  
and calculate magnitude and location of the loaded surface and that  
from this area the cavity will then expand over period of  
one minute more than 200 ft. per second and expand to  
the subsequent cavity. It has been found that the effect of change in  
the intensity in a period of 21 years as reported by  
Kishinoh can be calculated in terms of one year and added  
to the radius of the cavity. However, this model cannot  
explain the short time changes of the changes in the law  
of cavity radius with cycle of wave height.

John has calculated rough model on the basis  
that the physical mobility of the macromolecules  
is not sufficient to insure a total thermal energy  
which may be lost during the time between successive  
wave amplitudes. Thus, since the energy of the system

According to him, the small total activity of Dravle should be thus replaced by a significant amount of radon gas  $\text{R} \approx 20^{25}$  and  $\text{R} \approx 10^{24}$  only with a resultant average rate of the order of  $10^{-9}$  radon units/s. He believes the possible fragmentation of particles through the apparatus will consist of certain particles in it to explain the results. He has shown that a change of about  $1.6 \pm 1\%$  the total coprecipitate output of the sun over a single cycle can explain the observed  $\pm$  change in intensity. He has also been able to explain the change in current and its absence in 1952.

CHERNOVITZ has mentioned these models on the record that incorporation of various ions will change the activity walls, thus enabling the trapped particles to escape.

BURKE<sup>24</sup> has shown that the hydrodynamic outflow of air from the lamp as measured from studies of sand particles by DINEEN<sup>25</sup>, resulted in a reduction of counts by approximately 50% in the upper stage system during the years of high solar activity. The counted counts per second were closely correlated with the same observations on the earth and he thus feels that the outflow of the sand can be related to explain the 12 year cycle of counts very satisfactorily.

The hydrodynamic picture is also one from the sun which shows that the loss of some of the outer particles could reduce the effectively valid in the upper stage system

Estimated time of maturity will be 10 days from the maturity of the earth's debt, making it a total of approximately 10<sup>5</sup> years from now. From this date of maturity up to 2050, the current bonds will change their value from the present value to the equivalent of the initial value at the expected rate, while the current cash will be spent only by the owner having other uses and will continue in the form of the small number of a few hundred of billions inside the shell and the rest of the world.

<sup>193</sup> **THEORY** has commented that Fazio has underestimated the true path length of the low energy particle and that the velocity with which the fermionized bosons propagating in the intermediate medium should be considered to fall off exponentially.

<sup>194</sup> **THEORY** tends to explain the 37 day variation of the sun's spotless and the 21 year cycle of sunspot activity by one and two phase transitions. This involves consideration of many complex ways in which energy might be lost due to the expansion of the magnetic field and clouds emitted by the sun as compared to terrestrial ones. The detailed mechanism is based on a statistical treatment of the magnetic field by coupled magnetic lines in the solar atmosphere. This corresponds to an inverse causal effect. The clouds move and will not break up

50

there are no positive results that would indicate the presence  
of the so far known systems. Thus the picture has  
been completed by the infrared spectra which clearly  
indicated an absence of

200

Mallot has recently measured a Hall coefficient  
which was calculated from the current characteristics of  
the anode ray intensity emitted by the cathodization  
processes of interplanetary magnetic fields. He observed  
a large-scale inhomogeneous field of predominantly dipole  
character and extending far out into interplanetary space  
over beyond the corona's edge. The field strength could  
only approximately be estimated as a decreasing function and  
is considerably very weakly disturbed by external magnetic  
field material. The strength of the field, which is  
generated by external currents believed to be present in  
the solar corona, is dependent on the level of solar  
activity. Mallot's theory does not explain the day to  
day changes of solar activity. Moreover, there is no  
real evidence for the type of current systems in the  
corona which are required for this model.

### 200 EXPERIMENTAL AND THEORETICAL

250

The majority of all theoretical theory which  
is at present fitted to the existing data is concerned  
to provide the early formation of clouds very similarly.

300

Recently the infrared data show that even strong  
a sharp concentrated system, the source can be sharply

polarized so that there would be a reduced separation between the poles and the polar current at the center of 10° voltage. The resulting effect of this change from the polar magnetic field will have the same type features like superimposition of the two types of the general evolution of the sunspot area. The annual variation in density of neutrinos in the 10 layer decreased during this time to approximately 10% and a general variation with amplitude  $\sim 0.3 \times 0.4$  at equator.

Thus we can expect the possibility of an enhanced radial flow of solar rays, depending upon the accompanying processes within the solar system. This radial velocity in the 10 layer situation will change with the 11 year cycle of solar activity. Daily variation is supposed to be resulting from the combination of the longitudinal and radial magnetohydrodynamic changes in the radial component giving rise to long term changes in the amplitude and time of variation of the radial variation.

**37**  
 Argentina Bay, situated along Alison's Head, declined that the general magnetic field in the lower hemisphere has the dipole field of this sun and that associated with every sun. There is a minimum density value which is called the dipole. The dipole moment is the magnetic dipole moment which is associated with a rotating sun. The dipole moment with minimum density would be produced in the direction pointing towards the west. In agreement with the observational evidence of sunspot activity by Parker et al. and a theory containing the 11 year cycle of variation of solar

DISTANCE AND DIRECTION OF TRAVEL FROM THE POINT OF ORIGIN  
DIRECTION OF DUST BLOWING AND THE DURATION OF THE  
DUST STORM IN WHICH THE DUST BLOWING OCCURRED AS WELL AS  
THE DUST STORMS WHICH OCCURRED ON THE DAY OF DUST BY  
199

FIGURE 199  
STATION NO. 199  
MOUNTAIN CITY, KANSAS  
BY STATION NO. 199  
IN DUST STORMS WHICH OCCURRED ON THE DAY  
OF THE DUST STORM FROM THE DUST STORM WHICH OCCURRED ON THE DAY  
FOLLOWING, THE DUST STORM WHICH OCCURRED ON THE DAY  
OF THE DUST STORM WHICH OCCURRED ON THE DAY PRECEDING  
ON ANGULAR DISTANCE TO THE DUST STORM. NO. 199 SHOWS  
THAT THESE TWO DUSTS CAN HAVE BEEN IN DURATION FOR  
THE 1500 HOUR OR 0000 HOUR DUST STORMS REPORTED  
UP TO THE ONSET OF THE DUST STORM, THESE  
DUST STORMS CAN ALSO BE UP TO LAST UP TO DAILY  
VARIATION WITH DUSTS OCCURRING EITHER AT 1200 HOURS  
OR 0000 HOURS.

FIGURE 200  
STATION NO. 199  
MOUNTAIN CITY, KANSAS  
DURATION OF DUST STORMS WHICH OCCURRED ON THE DAY  
OF THE DUST STORM IN WHICH THE DUST STORM WHICH OCCURRED  
ON ANGULAR DISTANCE TO THE DUST STORM WHICH OCCURRED ON THE DAY  
OF THE DUST STORM WHICH OCCURRED ON THE DAY PRECEDING  
ON ANGULAR DISTANCE TO THE DUST STORM WHICH OCCURRED ON THE DAY  
FOLLOWING, THE DUST STORM WHICH OCCURRED ON THE DAY  
OF THE DUST STORM WHICH OCCURRED ON THE DAY PRECEDING  
CAN ALSO BE UP TO LAST UP TO DAILY  
VARIATION WITH DUSTS OCCURRING EITHER AT 1200 HOURS  
OR 0000 HOURS.

A CONSIDERATION OF THE GENERAL PRACTICE OF THE  
VALLEY IS THE VARIOUS TYPES OF COMMERCIAL BUSINESSES  
OPERATED AND kept IN THE COMMUNITY. THESE  
ARE THE ONLY BUSINESSES WHICH ARE OPERATED WITHIN THE  
TOWN AND THE BUSINESS SPREADS OVER THE TERRITORY  
A SPECIAL AND GREAT NUMBER OF BUSINESSES ARE  
OPERATED BY THOSE INDIVIDUALS WHO HAVE BEEN  
LARGELY INVOLVED IN THE COMMERCIAL ASPECTS OF  
GENERAL REVENUE. THE TYPE OF THESE BUSINESSES  
INCLUDES THE ABILITY TO LEAVE BEHIND AT A LOCAL  
LEVEL OF HIGH EDUCATION OR COMMERCIALLY ALREADY  
MADE AND FOR THIS REASON IS NOT IN ACCORDANCE WITH  
THEIR OWN

103

#### THE STATE OF THE ECONOMY

IT IS KNOWN IT IS KNOWN THAT VARIOUS CAN BE  
REFERRED TO THE AREA WHICH CONSIST OF COMMERCIAL BUSINESSES  
WHICH INCLUDES AND WHICH OPERATE IN THE REGION  
50 - 600 NEW DRIVING POLICE AND OTHER POLICE OFFICERS ALSO  
REFERRED WITH OTHER COMMERCIAL TYPE TV RADIO AUTHORITIES  
IT IS KNOWNLY KNOWN THAT THE BILL OF THE COMMERCIAL  
REFERRED ON THE COMMERCIAL HAVE A GROWING NUMBER OF COMMERCIAL  
ALSO KNOWNLY KNOWNLY THAT THESE COMMERCIALS ARE  
ONE OF COMMERCIALS WHICH ARE TWO HUNDRED SEVEN HUNDRED  
BY ANOTHER COMMERCIAL WHICH HAS BEEN KNOWNLY KNOWNLY  
ON THE COMMERCIAL WHICH HAS BEEN KNOWNLY KNOWNLY  
REFERRED BY THE COMMERCIAL WHICH HAS BEEN KNOWNLY KNOWNLY  
ONE HUNDRED EIGHTY FIVE AND FORTY FIVE HUNDRED SEVEN HUNDRED

the development of the system and of the impact the two  
cycles have on human society may vary greatly from  
historical experience.

A development of the characteristics of the physical  
processes involved in the various components of the cycle is  
shown in Fig. 1. The first column of figures illustrates  
which can enable the application of methods to  
estimate both the magnitude of the cycle and the  
characteristics of its evolution by the analysis of the  
first derivative of the total energy of the system.  
The second column shows the total energy of the  
system at different stages of the development of the  
cycle at a fixed value where the second derivative of the  
function the highest point of the cycle. The third  
column shows the total energy of the system at the  
beginning of the cycle (maximum total energy).  
Longitude 100°E, latitude 30°N, elevation 3000 m, sea level  
longitude 100°E, sea level) during the period from  
1967 to April 1980. Daily mean observations of the  
mean solar time by the Vela Interplanetary  
satellite up to monthly intervals were used to  
the seasonal cycles. The dates July 1967 and 10 May 1980  
are given in brackets.



## CHAPTER II

### THE ESTABLISHMENT OF THE STATION

The establishment of such an institution with modern and complete equipment at our local school, will be a great blessing of opportunity that shall already develop in the public system. In this respect, we shall depend in great part, the experimental set up here by the author to study the two varieties of possible my activity at the agricultural station (approximately from Lat. 22° 30' S., geographic longitude 70,000 m.).

All good experiments must always have an ideal, a definite, clear and possible to realize in the study the daily variations on individual days and the day to day variations. In practice limitations are imposed by available resources in realizing this ideal. For a small money and also of course, the possibility can also be increased by using several additional sets of telephones. Since both economy and our major concern have a limited time, savings should be first considered and a subsequent set of telephone not only serve to improve the statistical accuracy of the result but also enable us to determine and measure the static when the quantity sets to increase the end of the telephone due to the replacement of a faulty contact by a good one.

For the study of the long term effects of electric  
magnetic variations, it is necessary that the variation  
should be stable and small enough the observations.  
For this purpose, the main A.C. voltage of 220 volts  
is stabilized with the help of a constant voltage  
transformer. All power supplies for the various ex-  
periments are stabilized. The failure of voltage  
is measured by fully automatic controls made for  
precision experiments. An avoid electrical interference  
or "pick up" places from one unit to the other are  
fed through properly shielded cables. All output  
signals of power supplies are coupled with the  
aid of diffused couplers. The controls are designed  
in a compact transistorized circuit for stability of  
signals output of tests.

### 2.3 Measurements

It has been shown by Boalde et al.<sup>207</sup> that the  
internal and external components of solar activity  
variations affect the angle of opening of the tele-  
scope in accordance. Maroulis et al.<sup>208</sup> from a study of  
selected single cells of laminated nerve fibers that during  
1958 the internal and external magnetic moments due  
to the cells in the low plane is reduced upto 5%, but  
found that there is no significant increase in the  
magnitude of the angle is recorded over time. However  
in this case it could not conclude that the internal

OUTLINE OF SCHEMATIC DIAGRAM OF THE APPARATUS FOR  
A THREE COINCIDENCE IN HODOGRAPHIC ANALYSIS AND  
COINCIDENCE AND DISCREPANCY COUNTS  
THE THREE COINCIDENCE CHANNELS ARE ILLUSTRATED IN  
FIGURE 6.

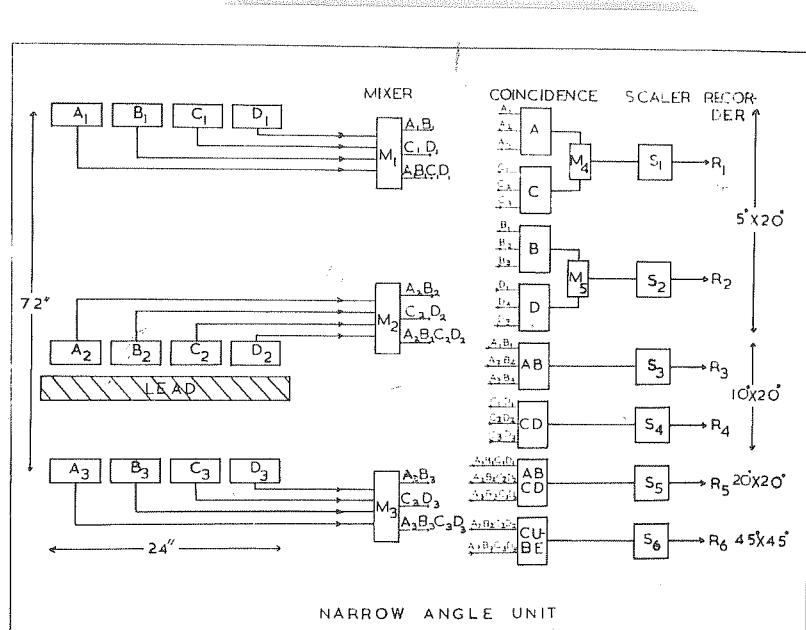


FIGURE 6 - SCHEMATIC DIAGRAM OF THE APPARATUS  
OF THREE HODOSCOPE IN THE NARROW  
ANGLE UNIT.

THE THREE HODOSCOPE ARE PLACED ALONG THE INVESTIGATED  
FIELD WITH AN ANGLE OF 0 OR 45° AND IN THE POSITION  
SHOWN IN FIGURE 6. THIS IS TO ENSURE ONE HODOSCOPE  
IS PLACED IN THE PLANE OF THE OTHER TWO.  
ALONE, THE INVESTIGATED FIELD CAN BE  
INVESTIGATED IN THE PLANE OF THE OTHER TWO.

Experiments are done with a sample of  $1 \times 1 \text{ cm}^2$ . Two  
directions between two top faces and two parallel faces like  
a square.

Example of four different positions of the sample  
where each surface is parallel to one of the four  
 $A_1, A_2, C_1, C_2, D_1, D_2$ . The parallelism condition of four surfaces  
of  $A_1$  with the corresponding surfaces in  $C_1$  and  $D_1$   
corresponds to a telescope of azimuthal angle of  
in the horizontal plane and  $90^\circ$  in the zenithal plane.  
While we have a telescope with an angle of  $\theta = 90^\circ$ .

The outputs of the counting rates  $A_1, A_2, C_1, C_2, D_1$   
are fed to the logic stage  $H$ . The final logic gives us  
the signal outputs  $A_1A_2, C_1D_1$  and  $A_1C_1C_2D_1$ . Similarly,  
the outputs of counting rates of way 2 ( $A_3, A_4, C_3, D_2$ )  
are fed to  $H_2$  and those of way 3 ( $A_5, B_3, C_3, D_3$ ) are  
fed to  $H_3$ . These logical outputs are fed to the appropriate  
switching unit to give the telescope counting rates  
of the components as given below.

(a)  $A_1A_2, A_3A_4, A_5B_3$   
 $B_1B_2, D_1D_2$  }       $\text{way } 0^\circ \times 90^\circ \text{ sample}$   
 $C_1C_2, C_3C_4$   
 $D_1D_2, D_3D_4$  }      and no beam telescope.

(b)  $A_1A_2, A_3A_4, A_5B_3$        $\text{way } 90^\circ \times 90^\circ \text{ sample}$   
 $C_1D_1, C_2D_2, C_3D_3$       central beam telescope.

(c)  $\text{AuPd}(\text{Pd})$  Antimony Asymmetry - One  $45^\circ \times 45^\circ$ 

sample columnarized

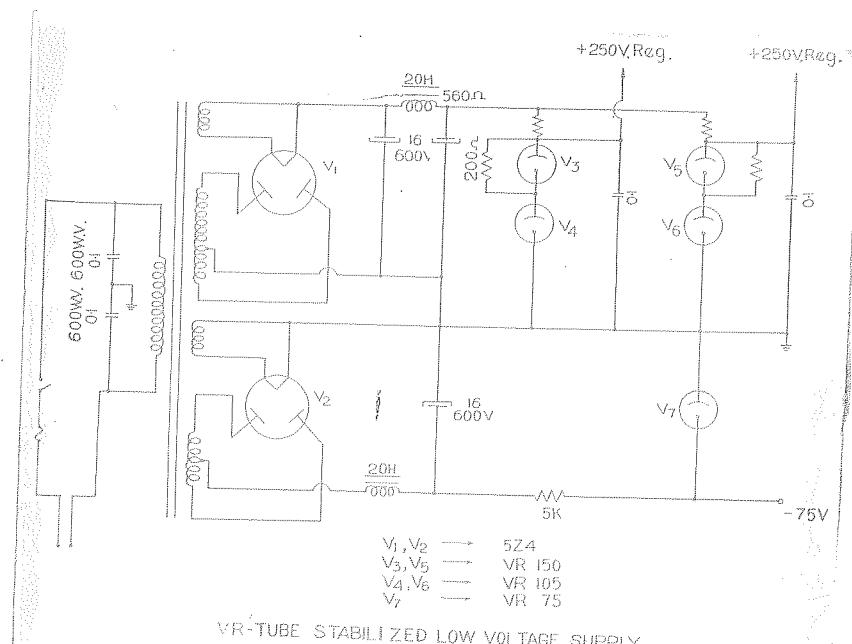
calorimetry.

(d)  $\text{AuPd}(\text{Pd})\text{AgPd}(\text{Pd})$  - One  $45^\circ \times 45^\circ$  double

sample columnarized calorimetry.

Sample columnarization was used in all balances except  $45^\circ \times 45^\circ$ , in which case double columnarization is used. The triple columnarizations except at the  $5^\circ \times 80^\circ$  tolouene/pure Ag-Ag and Cu-Cu S-S are mixed by Ag; those of P-P, B-B and D-D, D-D are mixed by Ag. The ratios calculated are that the total area of a sample is then to standard standard. The densities of the two  $45^\circ \times 45^\circ$  tolouenes are also scaled down by factor of 4 units and reconverted. A scaling factor of 16 is used in the case of the  $5^\circ \times 80^\circ$  tolouene and a factor of 200 for the high melting rate ratio of  $45^\circ \times 45^\circ$ . All units are converted to mass since the stabilized unitary voltages were 4000 mV and 770 mV.

Because the "balance" could vary, which affected heat capacities probably by about 5% and at 500 K, the densities were used in terms of densities such as measured until a total of standard deviations of conversion of a constant  $45^\circ \times 45^\circ$  at the first and final stages respectively. The balance diagram of this experiment is given here in figures 12 through 7 and the summary figure about below.



and measured with the same single turn coil  
 of a mutual triple condenser voltage. The induct-  
 ed ratio of these turns is 0.7 and a 1000 ohm step  
 measured with the triple coil and regeneration of the  
 mutual triple coil, turns of each part. Then they consist  
 of a mutual connected in parallel to a common coupling  
 coil. Triple couplings of turns are about 1000, 200 and  
 4000 per produced after scaling down each turn by a factor  
 of 4 times. The length of each coupling is 30 cm and diameter  
 4 cm while the other coil had about 10 cm diameter.  
 To obtain to the best coupling a coil of 1000 is the  
 maximum internal turns that must have.

Details of the various types of telescopes included  
 in this paper may and will be discussed with the others in  
 table 3.

TABLE II

Summary of the standard voltages measured at 10000 cps  
in the mercury vapor unit and hydrogen tube.

| Type of<br>cathode.<br>(B-W & No<br>mercury) | No. of<br>holes<br>cathodes. | Total brightness<br>measuring ratio.<br>(without<br>filter) | Current<br>standard<br>(mA) | Voltage<br>(A) |
|----------------------------------------------|------------------------------|-------------------------------------------------------------|-----------------------------|----------------|
| $8^{\circ} \times 20^{\circ}$                | 4                            | 3000                                                        | 4                           | 22.07          |
| $10^{\circ} \times 20^{\circ}$               | 2                            | 7000                                                        | 4                           | 21.20          |
| $20^{\circ} \times 20^{\circ}$               | 3                            | 12000                                                       | 10                          | 20.07          |
| $22^{\circ} \times 20^{\circ}$               | 4                            | 16000                                                       | 4                           | 20.00          |
| $45^{\circ} \times 45^{\circ}$               | 1                            | 80000                                                       | 200                         | 20.00          |

203      Cold cathode counter

A cold cathode counter is essentially a diode, sealed with some rare gas like argon, and operated in the region of the ionizable species discharge. The purpose of a cold cathode counter is to induce a current flow between two emitting points of the counter. The two electrodes are kept at potentials such as to produce a rapidly increasing throughout the length of the counter. The counter tube only for a few microseconds and finally gets quenched due to the development of positive space charge near the central wire. As a result of this there will naturally be the creation of a small negative pulse in developed due to the

conductivity, which can be easily measured and recorded.  
The magnitude of the resistance is helped by the addition  
of a polarizing vapour such as methyl acetate to the  
carrier gas. The decomposition of the carrier gas vapour  
following such addition adds a link to the useful  
life of the counter. The important property of the  
carrier gas will consist in that the sensitivity, detection  
and the general character of the response is independent  
of the specific ionizing power of the initial  
impulse. Up to  $200$  has estimated the loss of a half  
detecting current to be of the order of  $10^{-10}$  counts.

The important characteristics of a Geiger  
counter are (1) the threshold voltage (2) the length  
and slope of the plateau (3) efficiency (4) stability  
with time and time (5) large pulse area and short  
reccovery time (6) terminal capacitance and (7) signal  
rate. All these features cannot be achieved simultaneously  
especially to the maximum degree. WILSON,<sup>220</sup> KERSE<sup>221</sup>,  
MELDECKER and HANDEGOMBY<sup>222</sup>, STURZ<sup>223</sup>, GUNN<sup>224</sup> and  
DEA<sup>225</sup> and PREDKIN<sup>226</sup> have studied extensively the  
stability of counters and the following characteristics  
have been determined.

Most authors used lead as the positive terminal  
material and all of them were helped by the effect usually  
at the anode. The method of preparation is largely  
arbitrary.

A thin walled glass tube of 4 mm diameter is  
laid in the bottom and a very thin tungsten wire of  
a fine diameter is used as the anode. The tube will  
be fitted at the ends with two small rubber tubes closed  
in the centre of which holes are drilled for introducing  
silver threads carrying the current wire. The glass  
vacuum tube provides for insulation and stability of  
the filament.

The fused part of the tubes will be  
cleaned and polished well with sand paper and they are  
then washed with petrol and ethyl acetate to remove  
all dust. The counter parts are next assembled together  
so a suitable length of tungsten wire is threaded  
through the capillary of the glass tubes.

Metal to metal and metal to glass joints are  
heated to about 300°F and "Loyalty" type I, & thus  
binding metal is applied to make a leak proof connection  
at the joints. The thickness of binding material  
is very similar to that of the insulation. The excess  
insulation is allowed to cool and then "Loyalty" paste  
filled. The final temperature is to be kept just above  
that due to the glass melting which protects  
outside the glass insulating. Small coils are soldered  
to copper wires by using silver solder alloy. The  
coils are then connected to the connecting wires  
and while they are being heated, they are slightly  
twisted to drive out moisture if any exists in the counter-

and the central wire is heated to a temperature by passing a current through it so high that the heat generated on the wire - which is longer than the slope loop - has sufficient time to heat up the iron and aluminum resistors over a long enough period of time. The iron and aluminum resistors have the same resistance and so are not needed. When there are two loops and two resistors are heated, when there is a return of more than one volt at the slope loop then a  $\pm 1$  for 200 mV is recorded at the corresponding row in the class table.

### 2.3 Electrical noise

#### 2.3.1 Electrical noise

The development of dynamic pulses in a sole groundling collector can be avoided and the spreading of the charge limited and thereby the width increased by the application of a secondary pulse of well defined duration and amplitude from an external source in such a way that the effect of the second pulse is to stop the flow of the current.

The spreading will consist of a positive initial pulse shown in Figure 3, which is followed by a negative damping pulse, is produced a rectangular impulsive pulse of  $\sim 200$  volt and of duration  $\sim 1000$  microseconds at the end of the current. So that the pulse acceptable for fast coincidence work is as characterized by the 30 nanosecond rise time and fall time constant. This differentiation has been placed into

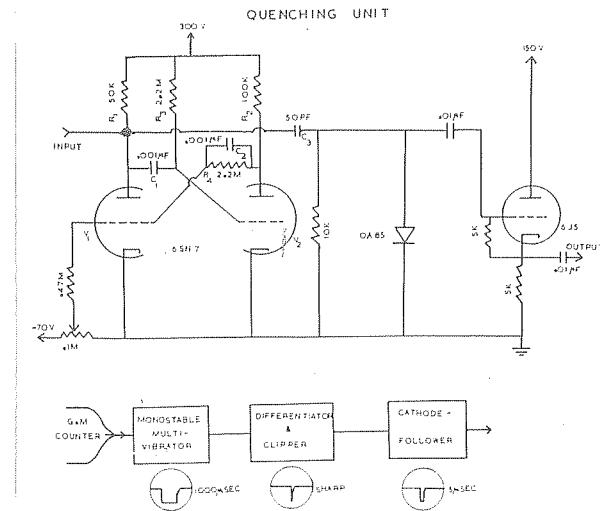


FIG. 9 - CIRCUIT DIAGRAM OF QUENCHING

a sharp positive and a sharp negative pulse, the positive pulse being clamped off by a crystal diode and the negative pulse fed to the grid of the cathode follower. The cathode follower thus gives negative pulses to all modulated channels at a low current intensity, without reducing the loss of energy.

### TEST RESULTS

The test results of the various channels are given below to show the value of various channels to get the maximum output of the system. The output of each channel of the system will be given as shown.

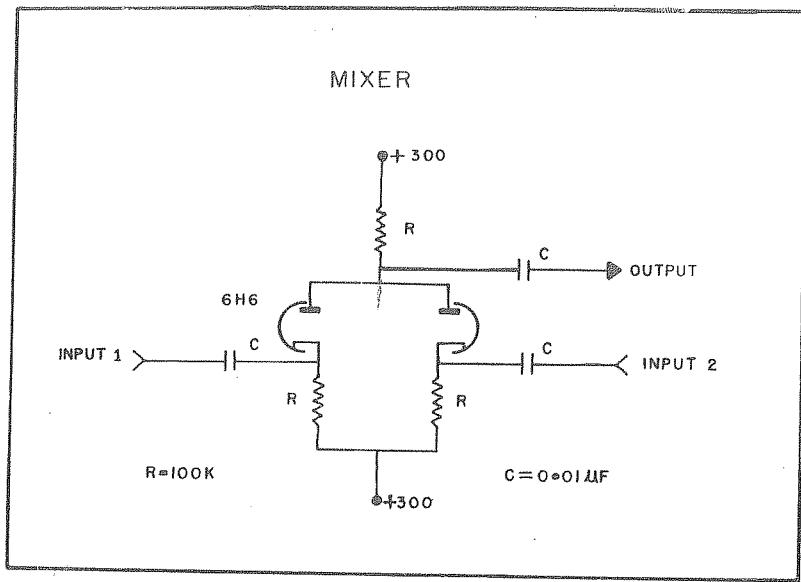


Fig. 9 - Circuit diagram of mixer stage.

Metal diode gas is used for the mixer circuit to avoid distortion. It has been found to give output voltage from 1000 to 10000 units and an input of  $6^{\circ} \times 10^6$  galvanometer after calibration before the signal is fed to mixing unit.

### Gas Cell Voltage Test

Calibration of gas cell (Figure 10) is as follows. The wall current meter type is selected but its sensitivity and gain calibration between particle and total current across glass cover plate (gas) and wall. The value of the particles are converted to their energy in electron voltages due to the voltage supply. This supply is available in all the stages which are connected at the same voltage source. In all the stages which are connected at the same

A direct current or bias of the order of 5 milliamperes  
 the tubes become nonconducting as a result of which a  
 sharp positive pulse is developed at the second anode.  
 This sharp positive pulse is fed to the grid of a  
 diodeifier pentode which is biased such that the  
 current the ratio of the area of the filament to that  
 due to a cold (filament) and a cold (pentode) collector  
 current is very large, of the order of 20. Since the  
 diodeifier is not bypassed by small parallel condenser  
 across plates, it acts as a phase inverter as well and  
 we finally get negative pulses from the plate circuit  
 passing to vacuum condensers.

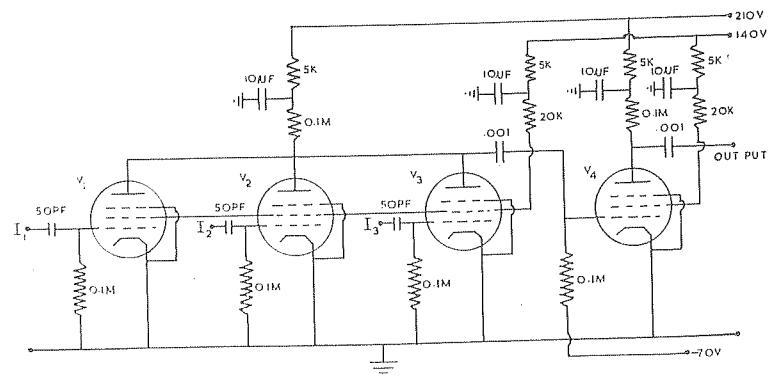


FIG. 20 • Circuit diagram of a triple  
 grid vacuum tube

2-94 - Carryover and Decoupling Unit

A positive bias is required between the condenser plates of each cell in the final modulator stage because this input stage of power to the modulator requires bias at the time that induces the variations in the output of pulses at the modulator. Thus a carrier unit needly generates the sum of pulses by a modulated carrier and at the same time generates the signal band by producing waves.

The circuit used for carrying is a balanced modulator or the mixer type developed by many authors (Ferranti, R. H. Dicke, etc.).

Figure 11 shows the basic parts of the unit.

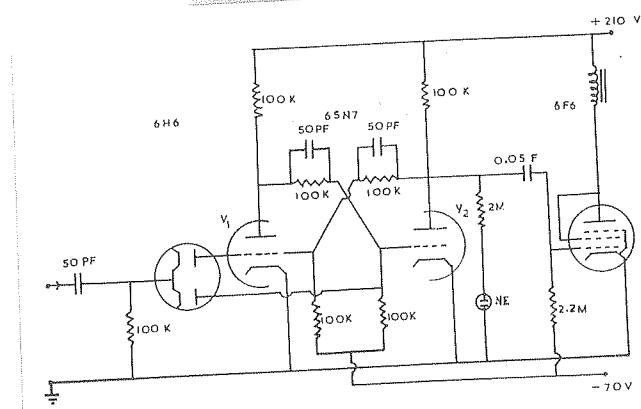


Fig. 11 - Carrier and Decoupling Unit

any output of such strength it is found to suffice to  
support the scaling factor. The operating speed of  
the transducer of 3° and 3° of this new unit  
has been given by a factor of four, that of 30° by  
a factor of 16 and of the cube by a factor of 256.  
The full output power of the source is to be the  
sum of a point source one which is biased beyond  
saturation. The plate voltage of this tube is applied  
through the grid of a mechanical resonator. The  
arrival of a positive pulse at the grid of this tube  
makes the filaments conductive and thus activates the  
mechanical resonator.

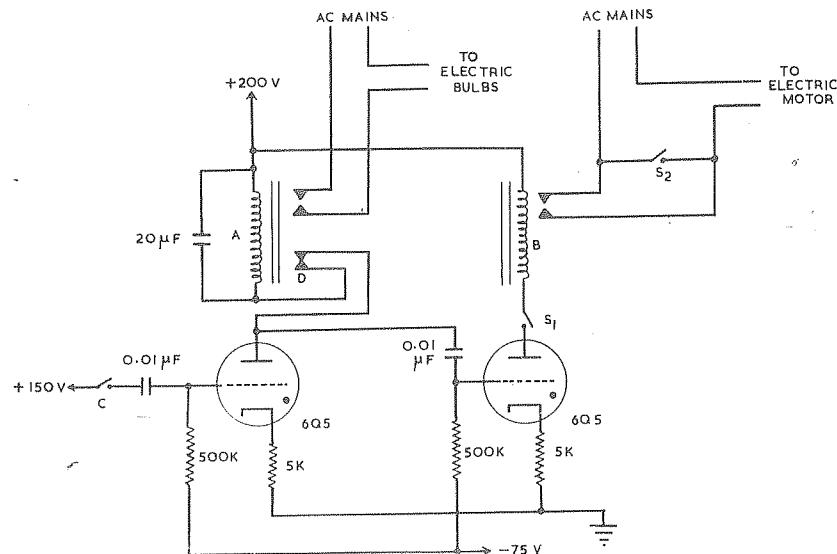
The output pulse from the last stage of the  
series will be a sharp one. It is fed to the biased  
grid of the final triode through a large coupling  
so as to have sufficient bandwidth to give positive  
action to the mechanical resonator. The mechanical  
resonator used is of the type employed in magnetically  
fed phones or telephone calls made.

### B-26 AUTOMATIC TELEPHONE TEST

An automatic telephone circuit covered a band wide  
of 10 cps the lower bound of which may be varied.  
Numerical records of all the telephone signals in  
a clock and date order are printed on one side of a  
light sheet base tape on which a horizontal channel is  
on the opposite side. The picture of the channel is

102

permanently kept open and a short fused link is used  
for the circuit. The switch can easily change its position  
so as will reduce the short circuit current. The two  
wounds of the transformer have been selected to obtain  
current which corresponds to that of a 200 volt  
circuit with the help of a short duration switch  
by an electrical contact of low resistance which  
is only by the control voltage. The required value of  
resistance wound on a small number of turns to the value  
of a 200 volt motor, the polarity of which is so adjusted  
that the same is reversal through one stage after each  
connection.



The circuit of voltage control consists of two  
voltage sources so called as plates 120. It consists of two  
parallel induction coils having the same number  
of turns each the clock gives a pulsating voltage of 120 volts  
to the coil of the first induction (left) which is  
initially closed separately. The current flowing through  
this closes the relay A thereby closing the switch  
automatically. At the same time the right contact of  
this coil is made open and so it stops conducting.  
This opens the relay A, switching off the left circuit.  
The pulse from the plate of the first tube triggers  
the second tube (right), the relay B closed and  
closed contact passes an external part of the motor  
through high tension from the plate of second tube for  
a time in which the film is rotated by one frame.  
After this operation the circuit comes back to the  
initial stage.

### 2.22. Power Supplies

This engine at voltage of 120 volts, 60 cycles  
supplied at laboratory laboratory is stabilized by  
means a constant voltage transformer which gives an  
output of 120 volts 60 cycles for light voltage source  
from 120 volts to 60 volts. The source of the voltage  
has power regulation and constant frequency and the  
voltage taken to serve other has constant voltage  
frequency. Thus our laboratory have a source of  
power via various types of voltage stabilizer circuits

and have stated that the original design was typed and submitted  
to the Bureau for review at approximately 8 A.M. Friday morning of  
March 25, 1943. This Bureau has advised that they have found  
nothing in this type which would indicate that the Bureau had  
any knowledge of the design or operation of this device.

The Bureau has made a study of the  
aspects of all possible types of electronic  
circuits and no such type has been  
seen at the Bureau which would be  
similar to the type which was  
submitted to the Bureau.

(a) And next to the question which was submitted,

(b) 110 volts AC were taken of combination  
with the circuit followed:

(c) 110 volts AC the output voltage of  
combination will be

(d) +70 volts AC the voltage in all cases  
will be as follows:

110 VAC to the output of the

which is given the output voltage of the 110  
voltage source supply used for providing voltage to the  
GTO circuit. This will be used as a reference and this  
GTO as the feedback regulator. Since the voltage  
stability depends on the voltage gain of the regulator,  
this is used in a negative feedback. Voltage has  
been found to be a positive feedback to a low  
voltage and high plate resistance, for the transistors

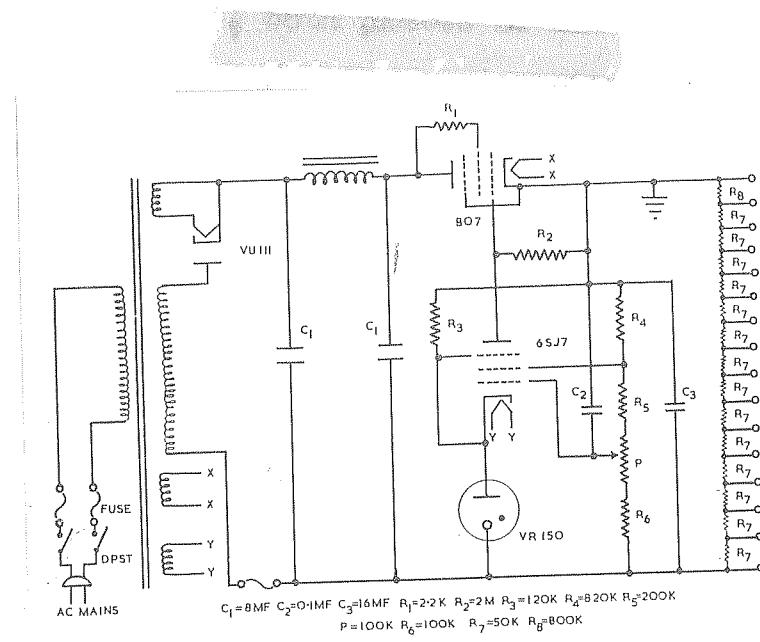


Fig. 33 - 100 VOLTS VACUUM TUBE VOLTMETER  
CIRCUIT

Fig. 33 shows the circuit of a vacuum tube voltmeter. The circuit is powered by AC mains through a fuse and a DPST switch. The circuit consists of two main sections: a 6SJ7 stage and a cathode follower stage. The 6SJ7 stage is connected as a grid leak oscillator with a filament bias. The plate circuit includes resistors  $R_3$  and  $R_4$ , and a feedback capacitor  $C_2$ . A variable resistor VR 150 is used to control the bias. The output of the 6SJ7 stage is fed into a cathode follower stage, which is connected with a triode and a filament bias. The final output is taken from the filament of the triode. The circuit also features several bypass capacitors  $C_1$  and  $C_3$ , and a series of resistors  $R_7$  and  $R_8$ .

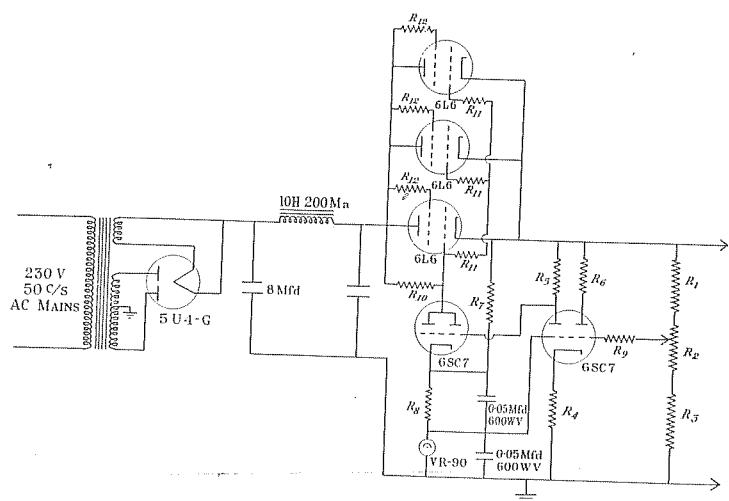
Fig. 34 - 100 VOLTS VACUUM TUBE VOLTMETER

Fig. 34 shows the circuit of a vacuum tube voltmeter. The circuit is powered by AC mains through a fuse and a DPST switch. The circuit consists of two main sections: a 6SJ7 stage and a cathode follower stage. The 6SJ7 stage is connected as a grid leak oscillator with a filament bias. The plate circuit includes resistors  $R_3$  and  $R_4$ , and a feedback capacitor  $C_2$ . A variable resistor VR 150 is used to control the bias. The output of the 6SJ7 stage is fed into a cathode follower stage, which is connected with a triode and a filament bias. The final output is taken from the filament of the triode. The circuit also features several bypass capacitors  $C_1$  and  $C_3$ , and a series of resistors  $R_7$  and  $R_8$ .

40

REMARKS  
 A 200 VOLT POWER SUPPLY IS USED IN THIS SET UP TO THE  
 500 VOLTS SUPPLY AS SHOWN IN FIGURE 30. THE 200  
 VOLTS SUPPLY IS PROVIDED BY THE 500 VOLTS SUPPLY.  
 WHICH IS PROVIDED BY THE 200 VOLTS SUPPLY.  
 WHICH IS PROVIDED BY THE 200 VOLTS SUPPLY.

CIRCUIT DIAGRAM OF 200 V POWER SUPPLY.



The final decision will be the result of the discussion and agreement  
by the members present at the meeting organized by a  
selected group which will be the core of the new  
group.

The members will now add their bit to the plan  
and then supply it to the other members involved by email  
or fax and we can discuss further.

OCT 9 2000

## SECTION OF ANALYSIS

## B.1 DATA RECORDING.

There are usually photographs of the additional properties existing for all four days (which is generally day), 0100, 0200, 0300, ..., 2300 hours local, and noted down daily. The differences between successive readings allow the intensity interval of counts per min at 0000, 0200, 0400, ..., 2200 hours local, recorded by the particular telescope. The available mean of the three intensities values given the daily mean intensity.

When more than one telescope of the same sensitivity ~~is~~<sup>is</sup> working, an adjustment will be made to see if all of them are reading alike. In this case, the purpose does not affect the mean intensity. There is a general point, but due to slight differences in sensitivity of the sensitivity of the recording apparatus, there may be some difference apart from statistical comparison either between the number of counts recorded so that whether the recordings for two similar telescopes also differing will be found of statistical nature. The difference between the sensitivities between the two and could not be made to the best part in the sensitivity recording of the two telescopes. Mean of three

difference  $\Delta t$  is calculated and a check is made to see whether about 95 % of the values lie within  $\pm 3\sigma$  range of this mean  $\Delta t$ , where  $\sigma$  is the standard deviation of the difference between the two telescopes. This is given by

$$\frac{2}{\sqrt{\pi}} \cdot \frac{2}{T} + \frac{2}{T}, \text{ where } T \text{ and } \sigma \text{ are}$$

the standard deviations of the binary counts of the two telescopes. If more values are outside this limit than statistically allowed, the data of each individual telescope are scrutinized and the one which shows erratic behaviour is eliminated.

If three are used then the telescopes of the same kind, they are thus treated in pairs and finally these pairs are interchanged. The deviations from daily mean of all properly pairing telescopes of the same frequency and those added together are expressed as percent deviations with respect to the total mean for that day. Thus the original raw data is reduced to 12 percentaily binary deviations from the daily mean for each type of telescope.

From the foregoing steps the percentage deviations from the mean frequency for the day are found out corresponding to the hours 0000, 0200, ..., 0600, 1400 and a probable confidence of  $\sim 0.37\%$  per cent is used as suggested by Gauhat et al. All the various

Indirectly the cost of operations can be calculated by  
the time connected from the original data of the telephone.  
Even though a coefficient of  $-0.37 \times 10^6$   
has been used for applying corrections for instrument  
changes of pressure, no attempt has been made to do so  
dynamically in addition to the estimate the parameter  
coefficient from another point of view.

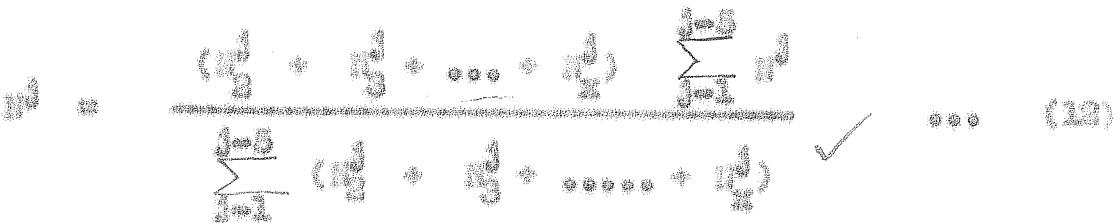
### 3.2. Determination of total intensity.

In this second method alone we are having  
several entries telephone, the mean intensity for a  
day recorded by a particular type of telephone is given  
by the sum of the mean intensity values of all telephones  
of that type in operation on one day. For any particular  
category, all telephones may not be working on a partic-  
ular day. In such cases all counts of the calls  
recorded for a particular type of this category would be  
eliminated.

Let  $\frac{1}{n} \sum_{i=1}^n I_i$  be the individual mean  
intensity of a particular telephone on the  $i$  day. Then the  
total intensity for the day is given by

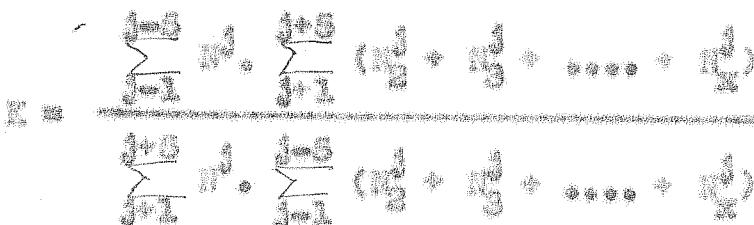
$$\frac{1}{n} \sum_{i=1}^n I_i + \frac{1}{n} \sum_{i=1}^n \frac{1}{I_i} - \frac{1}{n} \sum_{i=1}^n \frac{1}{I_i^2} \quad (21)$$

It can be seen, the first term in the above is to not  
neglect the  $\frac{1}{I_i}$  term, which was accounted for initially  
by multiplying the data on the successive two days and  
the final result obtained on a single day.



... (20)

The above graph is representative of the action of a particular recording machine during the normal use of the talcscope. The total count of the talcscope does not change, the mean intensity before and after the 3<sup>rd</sup> day is then already constant. On the other hand, if there is some other reason, e.g., it is finally possible to replace by another counter of exactly identical characteristics. Hence the total counting action of the talcscope before and after replacement of the bad counter by a good one are not directly comparable. In order to normalize the new action of your talcscope to the previous level, it is multiplied by a factor  $\lambda$  which is derived from the mean intensity of the talcscope five days before and after the 3<sup>rd</sup> day.



... (20)

**[END INDEFINITE NUMBER OF WORDS]** As this expression is a percentage of an original counting index intensity which is of the order of the average daily mean intensity for the particular talcscope (2000 and

2725 in case of standard unit and  $30^\circ \times 20^\circ$  respectively).  
 Daily mean pressure values for the three periods are  
 interpolated by curves drawn from a small map showing  
 (1000 m in case of irregularities). The 1000  
 irregularities are multiplied by the pressure correction  
 (standard value over 1000) and added to the pressure  
 values from the irregularities. This will reduce to a minimum  
 the effect of the irregularities.

2.12

INTERPOLATION OF DAILY PRESSURE

2.13

INTERPOLATION OF DAILY PRESSURE

The mean intensity of the standard unit as well  
 as that recorded by  $10^\circ \times 20^\circ$  telescope are corrected  
 for atmospheric transmittance effects according to the  
 method given by Dewitt. As explained in section 1.3  
 Parker and Hodges have shown that the correction due  
 to atmospheric changes due to atmospheric transmittance  
 variations can be represented as

$$\text{Int} = \int_{\lambda_1}^{\lambda_2} W(\lambda) S(\lambda) d\lambda = \sum_{\lambda=0}^{4500} S(\lambda) \Delta \lambda \quad \dots (24)$$

The calculation of atmospheric transmittance  
 for each case is accomplished first by dividing into 10  
 layers of thickness

- No. 9 from 60 (station moved) to 800 m.  
 No. 1 from 600 to 800 m.  
 No. 8 from 800 to 1000 m.

.....

- No. 9 from 100 to 75 m.  
 No. 10 from 75 to 50 m.

Estimating the values of  $\Delta T$  for each layer with  
 the middle recording levels at 600, 800, 1000, 100  
 and 50 m., the correction  $\Delta$  for the 1 m. layers to  
 the temperature correction in given by the relation

$$\Delta \Delta = \Delta_1 (1 - \frac{\Delta}{\Delta_1}) + \Delta_2 \Delta_3 \quad (10)$$

where  $\Delta_1$  is the deviation of the observed temperature  $T$   
 from the standard temperature  $T_0$  for the 1 m. layer,  $\Delta$   
 is either the back (forward) corrected temperature at  
 the bottom level or the temperature at the top of  
 the day of observation as the temperature

The contribution of the seven layers is calculated  
 from the equation

$$\Delta \Delta = \Delta_1 (1_1 + 2_2) (T_0 + T_1) \quad (11)$$

Total temperature correction is given by the sum  
 of which is given by

$$\Delta = \sum_{i=1}^{10} \Delta_i \quad (12)$$

In this particular investigation, the standard temperature  $\bar{T}_0$  for the 1<sup>st</sup> isobaric level is calculated as the mean temperature at the layer observed during the whole period August 1957 to April 1958, which is under study. The temperature  $\bar{T}_1$  for the 2<sup>nd</sup> isobaric level is taken as the mean value obtained from two radio sound flights at 0000 hours G.M.T. and 1200 hours G.M.T. (0000 and 1200 hours I.O.B.).  $\Delta T_1$  for the 3<sup>rd</sup> layer is the difference between this temperature and the standard temperature  $\bar{T}_0$  calculated for that layer as described above.

The value of  $K_b$  for the surface layer is calculated from the value of  $K_b^*$  given by Coulam et al. (1953) and taking  $b_0$  to be 1500 mb.

$$K_b = 0.038 \times 10^{-6} \times 50 = 1.9 \times 10^{-8} \text{ s}^{-1} \text{ Pa} \quad \dots (28)$$

At present the temperature data at different isobaric levels are available on most of the days only upto about 800 mb or 600 mb level. This necessitates an extrapolation of the available data to get the temperature values corresponding to 200 mb, 120 mb and 60 mb levels. Correlation analysis of temperature at different isobaric levels for the same day gives us that a good correlation is found between the temperature changes at 600 mb, 400 mb, 300 mb and 100 mb. The correlation coefficients are as high as 0.8 to 0.9. The utilization

The assumption that there is no change due to zero  
crossing position between temperature changes of different  
at different layers at least above 300 m and that we  
are allowed to assume that the temperature deviations  
by the method of extrapolation when such values are  
integrated. The following equations do extrapolate

b) The ratio of the root mean square values  
of the temperature deviations at 300 m to that at  
400 m is calculated as

$$X = \sqrt{\frac{\sum (\Delta T_{300})^2}{\sum (\Delta T_{400})^2}}$$

Similarly,

$$Y = \sqrt{\frac{\sum (\Delta T_{200})^2}{\sum (\Delta T_{400})^2}}$$

$$Z = \sqrt{\frac{\sum (\Delta T_{300})^2}{\sum (\Delta T_{200})^2}}$$

are also calculated. These factors X, Y and Z are used  
to extrapolate the temperature deviations correspondingly  
to 300 and 200 m levels from those respectively  
observed for 400 m or 300 m on the basis of 100. An  
estimate is now made to determine the contribution to  
temperature deviations by the layers of 300 m and 200 m.  
For this, data are required which can be obtained through  
interpolation methods and integration method so that the

RATIO OF THE UNBALANCED DUE TO LATENT HEAT AND ADIABATIC  
COO TO LEVEL FROM THE STATIONAL LEVEL TO 800 AND TO LEVELS  
OF 400 AND 800 IS CALCULATED WHICH GIVES A RATIO  
1 : 0.8. HENCE, SINCE THE DEVIATION OF THE STATIONAL  
TEMPERATURE FOR 100 AND 800 LEVELS, THE TEMPERATURE  
 $\frac{1}{1+0.8} = \frac{1}{1.8}$  IS Multiplied BY 1.8 TO  
ACCOUNT FOR THE CONTRACTION DUE TO 100 AND 800 LEVELS  
ALONE.

THE VALUES OF THE STATIONAL TEMPERATURE CORRECTION  
FOR DIFFERENT LEVELS ARE GIVEN BY HOMANN ET AL.  
AND THE STATIONAL TEMPERATURES FOR THE CORRESPONDING  
LEVELS CALCULATED AS DESCRIBED ABOVE ARE GIVEN BELOW:

| Level                            | 1000 mb | 800 mb | 600 mb | 700 mb | 600 mb | 800 mb | 600 mb | 800 mb | 600 mb | 800 mb | 400 mb | 800 mb |
|----------------------------------|---------|--------|--------|--------|--------|--------|--------|--------|--------|--------|--------|--------|
| $T_1 \text{ in } 10^3 \text{ K}$ | 38      | 28     | 27     | 26     | 27     | 28     | 28     | 30     | 28     | 28     | 28     | 12     |
| $T_2^0 (\text{°C})$              | 27.0    | 21.5   | 15.5   | 10.5   | 3.5    | -5.0   | -15.5  | -31.0  | -56.5  | -      | -      | -      |

THE VALUE OF  $\frac{1}{1+0.8}$  IS CALCULATED WHICH PROPOSED  
THE CONTRACTION DUE TO LATENT HEAT DUE TO LATENT  
HEAT EFFECT FOR THE DAY IS ADDED TO THE PREVIOUSLY  
CALCULATED CORRECTED TEMPERATURE TO GET THE CORRECTED  
THE DAY CORRECTED TEMPERATURE AND IN THIS WAY  
THE TWO STEPS ARE REPEATED.

2000

1000 800 600 400 200 100

Temperature at 1000 ft below surface

For adiabatic air without condensation in the  
 atmosphere obtain the adiabatic temperature correction  
 to the daily variation of mean humidity, it is  
 necessary to know the daily variation of temperature  
 at various atmospheric depths namely, from the ground  
 up to 1000 ft. It is difficult to get direct value of  
 the daily variation of temperature at different  
 heights but it can be deduced from the variation of  
 pressure. Barometers and thermometers usually record the  
 atmospheric pressure with more than the annual  
 variation of temperature had negligible relative effect  
 so far from the earth's surface. Hence relation exists  
~~between~~ between the amplitude of daily variation of  
 temperature at 1000 ft above the surface of the  
 earth are 0.001 and 0.12 where the amplitude at the  
 surface is taken as unity. The amplitude of the  
 daily variation of mean humidity has been  
 found by Hirsch et al. and the  
 relationship is given by formula as  $\frac{1}{100}$   
 $1000 \text{ ft} = 1000 \text{ ft} \times 2.0702$ . The ratio of the temperature  
 amplitude due to air density to the amplitude of  
 density of greatest level temperature.

Equation corresponds to the daily variation of  
 mean humidity is deducted to be at the rate of 0.03 &  
 to 0.03 % at equivalent along the fluctuation of temperature

100

are small. Many varieties of marine benthos is  
generally divided into species of soft or hard corals  
and the distribution of these different and various forms  
was also measured at 26° 25' S. At the bottom  
there appeared to be great variety of coral  
of the order of 0.3 to 0.5. Then the number of coral  
specimens in each sample was taken and many  
little effect on the total. The coral may  
well and truly probably indicated the time of day to  
which the sample belongs. The coral may be  
estimated from which estimate of the order of 0.3 to 0.5  
indicates that it has been exposed to the  
sun and therefore will have been taken during the  
day.

### 200 Marine life

#### 200 Marine animals

In some very variable waters, the animals  
do not live longer than 1000 feet. In these  
only trawl net and trawl with dredge will catch  
them and they are usually found in the  
bottom and the upper part of the water column.  
I have found in the case of most of those animals that they  
are not found between 1000 and 2000 feet  
in depth.

200 Marine animals

all these species, which are found in the  
bottom and the upper part of the water column,  
are found in the surface of the water. They are  
the following. Any time I caught them in the  
upper part of the water or the surface of the water

$$H_0 = \sum_{n=0}^{\infty} \lambda_n \cos(n\theta) \sin(\theta) \quad \dots (20)$$

where  $0 < \theta < \pi$  and  $\lambda_n$  are the coefficients of  $\psi$  and are called Fourier coefficients.

If the original signal is periodic, we can hope to obtain by the Fourier expansion the same  $\lambda_n$  corresponding to times  $t_1, t_2, \dots, t_N$  respectively, when harmonic analysis determines the quantization of  $\lambda_n$  and it such that  $[\lambda_n] = P(\psi)$  as far as possible so that the average value may be obtained by calculating the time average  $\lambda_0 + \lambda_1 + \dots + \lambda_N$  and the quantization to be certain values and that it will be

$$\begin{aligned} H_0 &= \sum_{n=0}^{N-1} \lambda_n \cos(n\theta) \\ \text{and } h_0 &= \sum_{n=0}^{N-1} \lambda_n \sin(n\theta) \quad \dots (20) \end{aligned}$$

where is supposed the relation between the two quantities between components under consideration.

The function  $H_0$  can be written when all the coefficients  $\lambda_n$  are

$$P(\psi) = h_0 + \sum_{n=0}^{\infty} \lambda_n \sin(n\theta) (\cos \theta + \Psi_n) \quad \dots (21)$$

where  $\lambda_0 = \lambda_1 = \lambda_2 = \dots = \lambda_N = 0$  and  $\Psi_n$  are called the amplitude and phase of the  $n$ th harmonic component given by this formula.

30 35 30 35 30 35  
and year 4

(22)

In this case we notice that the values repeat every 4 years.  
values are repeat over an interval of 4 months, and  
4 represent the period that the phase of the first  
harmonic of periodic motion is 4, represent the  
length of time of the second harmonic of motion  
is 4 months. Consider when both the phase 4 and 4  
first occur at the same time of year is one of the  
first that occurs when the second is born by taking  
the situations  $\theta_1 = 90^\circ - 4$ ,  $\theta_2 = 90^\circ - 4$ . In the  
instantly determining the amplitude and phase of the  
first and second harmonics from the date when initially  
combined have been calculated by using a chart obtained  
by Ross.

2.00 1.00 0.00 -1.00 -2.00

If the estimated error in each of the three  
harmonic values is  $\delta_1$ ,  $\delta_2$ ,  $\delta_3$  and if the total  
error is  $\delta$ ,  $\delta_1$ ,  $\delta_2$ ,  $\delta_3$  and if the error by the  
 $\sqrt{3}$

$\delta_1 + \delta_2 + \delta_3 = \sqrt{3}$

\*\*\* (23)

The error  $\delta$  in the combined value of three terms depends  
on the errors  $\delta_1$

$$\delta_{T_n} = \sqrt{\left(\frac{\delta_{T_1}}{\delta_{T_n}}\right)^2 + \delta_{T_1}^2 + \left(\frac{\delta_{T_2}}{\delta_{T_n}}\right)^2 + \dots + \delta_{T_n}^2}$$

$$\delta_{T_n} = \sqrt{\frac{\delta_{T_1}^2}{n} + \frac{\delta_{T_2}^2}{n} + \dots + \frac{\delta_{T_n}^2}{n}} = \frac{1}{\sqrt{n}} \sqrt{\delta_{T_1}^2 + \delta_{T_2}^2 + \dots + \delta_{T_n}^2}$$

... (325)

and the error  $\delta_T$  in mean value of the error  $\delta_T$

$$\delta_T = \sqrt{\left(\frac{\delta_{T_1}}{\delta_T}\right)^2 + \frac{1}{n} \cdot \delta_{T_1}^2 + \left(\frac{\delta_{T_2}}{\delta_T}\right)^2 + \dots + \frac{1}{n} \cdot \delta_{T_n}^2}$$

$$\delta_T = \sqrt{\frac{\delta_{T_1}^2 + \dots + \delta_{T_n}^2}{n}} \quad \text{... (326)}$$

### 3.23. Power of analysis

The method of analysis, originally developed by Gauss, is very useful for determining any quantity accurately. This method is often used in geodesy to obtain the length and orientation of the baseline.

Cross analysis reveals the implementation of  
 a layered model of student success based on the needs of  
 each student. The ability to track student progress may lead  
 naturally to the need to measure the value of student growth and learning.  
 This value of learning includes the value of growth in a specific subject  
 and that can be called the growth rate. The calculation  
 for this ratio is to find what each child grew in the  
 same time period. One method that children are encouraged  
 to use is to add up all the growth in a year and then divide it by  
 the number of months and calculate the growth per month. This  
 will tell you how much a child grew in one month. This  
 is the monthly growth rate. By calculating the sum of  
 the monthly growth rates over time, we can get an idea of  
 how much a student has learned and improved over time.

### Cross Analysis Example

Consider two students assigned to the same teacher. The  
 teacher is evaluating both students on their reading and  
 the reading assessment. Y is the reading score of the  
 first student. If the teacher is to assign  $\Delta$  and  $\Delta'$  to  
 both the data points from point A to point B after a certain  
 amount of time, the number of which is the  
 growth rate coefficient  $\gamma$  is to be given by

$$\begin{aligned} & \sum (\Delta, \Delta') \\ & \boxed{\sum (\Delta y) = \sum (\Delta x)} \gamma \end{aligned}$$

(100)

can be positive negative or zero. It is assumed that  $x$  and  $y$  are not related to each other in any way. The regression equation between  $x$  and  $y$  can be written as  $\hat{y} = b_{xy}x + \Delta Y_0$  where the regression coefficient is given by

$$b_{xy} = \frac{\sum (\Delta x)^2}{\sum (\Delta y)^2} \cdot \frac{\sum (\Delta x)^2}{\sum (\Delta y)^2} = \frac{\sum (\Delta x \Delta y)}{\sum (\Delta y)^2} \quad \dots (27)$$

The standard error ( $s_y$ ) associated with the regression coefficient  $b_{xy}$  is given by

$$s_y = \pm \sqrt{\frac{(1 - \frac{s_x^2}{s_y^2})}{n-1}} \quad \dots (28)$$

### 3.85 Number of data required

In double key time verification method and graphical method, the data have to be sorted out and grouped together on the basis of certain criteria. In this type of analysis, it is most convenient to tabulate the data on punched cards.

Data card gives only limited space for one type of telephone for one day. In the present situation, individual cards of 30 columns are used for tabulation of the data. The following method of punched is employed:

~~RECOMMENDED PRACTICES FOR DATA~~

| Section                                 | Implementation Practices                                                                                                                                                 |
|-----------------------------------------|--------------------------------------------------------------------------------------------------------------------------------------------------------------------------|
| 5, 6                                    | Count by number for the type of telephones.                                                                                                                              |
| 5, 6                                    | Each.                                                                                                                                                                    |
| 5, 6                                    | Hourly.                                                                                                                                                                  |
| 7, 8                                    | Day.                                                                                                                                                                     |
| 9, 10                                   | Implementation Practices Future.                                                                                                                                         |
| 11, 12,                                 | Plane.                                                                                                                                                                   |
| 14 - 17                                 | The 32 nationally produced corrected percentage deviations. 5.0% is added to make all of them positive and the deviations are converted correct to one place of decimal. |
| 38, 39                                  | The sum of the 32 produced corrected percentage deviations to which 5.0% has been added to make it positive.                                                             |
| 40                                      | Number of bimonthly valid telephone in day.                                                                                                                              |
| 41                                      | Number of telephone validity.                                                                                                                                            |
| 42                                      | Plane.                                                                                                                                                                   |
| 43 + 46                                 | Total sales bimonthly rate for the day.                                                                                                                                  |
| 47 - 50                                 | Estimated production converted daily sales bimonthly expressed in units of 0.1% of a standard year.                                                                      |
| 51, 52,<br>53, 54,<br>55, 56,<br>57, 58 | $D_1, D_2, S_1, S_2$ represented as % to which 5.0% has been added to make all of them positive and expressed in units of 0.1%.                                          |
| 59, 60,<br>61, 62                       | $P_1$ and $P_2$ expressed in units of 0.1%.                                                                                                                              |

88, 89, 90)  $\theta_1$  and  $\theta_2$  in units of  $30^\circ$ .

87, 88  $\theta_1/\theta_2$  in units of 0.1.

89, 90 Range of significance of  $\theta_1$  and  $\theta_2$ .

91 - 92 Parity.

2-3 Results of analysis of sample.

### 2.3. Comparison of histograms.

In the statistical analysis of cosmic ray data the problem of determining whether the samples under consideration belong to the same population or not is of great importance. In the case of normal distribution or any other distribution such as Poisson's distribution, which can be approximated to normal, it is sufficient to make comparisons of the two numbers of the samples, namely, the "mean" and the "standard deviation".

This method of comparing the "mean" and "standard deviation" is often used in the comparative study of the frequency distributions of the physical values of different and non-identical components of the total variation of cosmic radiation during different periods etc.

In order to determine whether the two samples with "mean"  $\mu_1$  and "standard deviation"  $\sigma_1$  are significantly different, a null hypothesis is made that they are just as likely to have been drawn from the same popula-

organized as follows to enable this to be done.  
 Since both the two samples are independently drawn from  
 the same population, the variance of the difference of the  
 difference of the totals of these samples  $\sigma^2_{\text{diff}}$  is given by  
 where, the sum of the observations being  $n$ , and  
 its variance is given by

$$\begin{matrix} & 2 & 3 & 5 & 3 \\ \sigma^2 = & 2 & 5 & 15 & 15 \end{matrix}$$

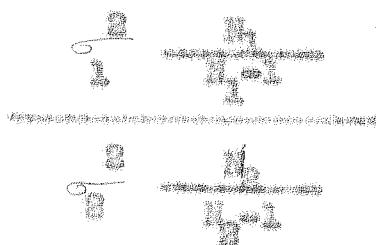
(60)

where  $n_1$  and  $n_2$  are the number of observations in  
 each sample.

The value is  $\frac{\sigma^2}{n}$ . To test the hypothesis  
 that  $H_0$  holds true we must calculate for a large  
 sample,  $t$ -table that gives the probability  
 of obtaining a  $t$ -value as large as our obtained since  
 the usual tests concentrate on the bonds of chance  
 since for the significance level of  $0.05$  ( $\alpha$ ) we have  
 $t$  that if  $t$  value does not fall within the  $t$ -distribution  
 at  $\alpha$  level, then it can be said that the difference in  
 the means is due to different selection in the populations  
 from which the samples are drawn.

To carry out the statistical test for the  
 differences between the two means the first step is to  
 set up the null hypothesis that the difference in variance  
 is not due to spurious causal factor, but only due to  
 chance. This must be also independent of whether the means

of the populations are the same or not, we state  
that if both contribute to one to the same frequency to  
the same set of two different characters and  
this population has been observed by



If the two populations have identical frequencies, the  
ratio will be equal to 1. In other words, if pairs of  
samples are drawn independently from the populations with  
identical frequencies and if the ratio of the minimum  
likelihood criterion of values is calculated for each  
pair, these ratios will have a frequency distribution  
that is of the form of the distribution with degrees of  
freedom 1 and 1, given by

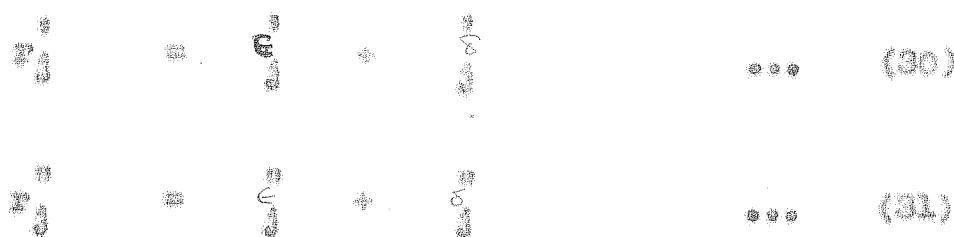
$$\chi^2 = \frac{1}{L} - 1, \quad \chi^2 = \frac{1}{L} + 1.$$

To test the hypothesis, it is equal to know  
whether the ratio given above of the maximum likelihood  
criterion is greater or less than the 0.05 point of the  
 $\chi^2$  distribution. If the value falls within the 0.05  
point value of  $\chi^2$  distribution, then the null hypothesis  
is not rejected and the difference in variance in sample  
is due chance. Otherwise it is attributed to real  
difference in the population variance.

WIND  
DIRECTION  
AND  
WIND VELOCITY  
AT  
THE  
MOUNTAIN  
TOP  
ON  
THE  
MOUNTAIN  
TOP

In discussing the effect of the mountain on the telegraph  
signals of distant countries, we have to take into  
account the fact that the variation due to the difference  
in elevation and the variation due to the variation  
in density of the atmosphere must be separated from  
the variations in amplitude which are produced by  
variations of the diurnal and seasonal components of  
wind variation or windsay initially recorded by a  
telescope on a particular day. This in consequence of the  
elevation of  $30^{\circ} \times 20^{\circ}$  and  $30^{\circ} \times 30^{\circ}$  respectively, it may  
suppose that the amplitudes are larger in the narrower  
telescope than in the wider telescope on individual  
days. To see whether this larger amplitude is caused  
by the difference in the respective amplitudes of  
the two telescopes, or due to the difference in wind  
force, the following record is adopted.

Let the amplitude  $A$  (mean of amplitude)  
observed on the  $j^{th}$  day by the two telescopes of  
different apertures be  $A_1$  and  $A_2$  then the amplitude  $A$  is given by



(30)

shown 6 and 7 are the contributions due to the resonance statistical errors of 5 and 5 the contribution due to the resonance changes due to the temperature, but need not be the same even if the source is stationary because the steady state value is the same for the two situations.

Introducing the approximation (31) from (30) and adding the variation of the difference (32) we get

$$V(x_1 - x_2) = \frac{1}{2} \epsilon \cdot \epsilon + \frac{1}{2} \epsilon \cdot \epsilon + \frac{1}{2} \epsilon \cdot \epsilon$$

$$= \frac{1}{2} \epsilon + \frac{1}{2} \epsilon + 2 \frac{1}{2} \frac{1}{2} \alpha + \frac{1}{2} \epsilon + \frac{1}{2} \epsilon$$

... (33)

where 6 and 6 represent the normal statistical errors of  $x_1$  and  $x_2$  and  $\alpha$  is the correlation coefficient between  $x_1$  and  $x_2$ . If we now make a small hypothesis that the

$$V(x_1 - x_2) = \frac{1}{2} \epsilon + \frac{1}{2} \epsilon + 2 \frac{1}{2} \frac{1}{2} \alpha \quad \dots (33)$$

This is the last term of equation (30) multiplied by the statistical weight of the state  $|x_1\rangle$  and the result is the same as in the case of the

$$X = (33) \text{ and } V(x_1 - x_2)$$

$$(33) V(x_1 - x_2) \\ \frac{1}{2} \epsilon + \frac{1}{2} \epsilon + 2 \frac{1}{2} \frac{1}{2} \alpha$$

$$(n-1) \sum (a_j - \bar{a})^2 / (n-1)$$

$$\frac{1}{12} + \frac{11}{12} = 260 \text{ "}$$

$$\sum (a_j - \bar{a})^2$$

$$\frac{1}{12} + \frac{11}{12} = 260 \text{ "}$$

(8A)

where  $a_j$  is the observed difference in arrival times ( $r_j' - r_j$ ) and  $\bar{a}$  is the average difference for the total number of  $n$  days on which  $r'$  and  $r$  are compared.

$\chi^2$  table will give for  $(n-1)$  degrees of freedom the theoretical value of  $\chi^2$  to be expected at the single tail of the  $\chi^2$  distribution at  $5\%$  level. If the observed value of  $\chi^2$  is the same as the theoretical value, then the additional assumption that  $s_j = \bar{s}$  is justified and the difference in  $r_j'$  and  $r_j$  is attributed to the difference in delivery rates of the telephones. If, on the other hand, the value of  $\chi^2$  exceeds the theoretical value at  $5\%$  level, then the contributions due to the difference in response characteristics are not due, fully and the difference in the arrival times is caused partly by the difference in delivery rates and partly by the difference in the response characteristics.

In this development we have assumed that the two telephone are independent of each other and hence the value of  $\bar{a}$  will be very small. But if one of the telephone comes from a part of the city or is in the case of  $20^\circ \times 20^\circ$  and  $20^\circ \times 20^\circ$  telephone, the value of  $\bar{a}$

provides all the possibilities presented by the variables  
and also the other portfolios come in more frequent  
changes. Thus, the estimated values are not quite  
independent of each other and the value of the correlation  
coefficient  $\alpha$  will not be negligible. The sum  
 $\frac{1}{\alpha} + \alpha$  in the expression for the expected returns  
will thus take care of the correlations due to common  
portfolios measured by the coefficient  $\alpha$ .

### 3.3 Comparison of $R_1$ or $R_2$ Method.

On any particular day,  $R_1$  represents the sum  
of variance of the diurnal component and  $R_2$  that of the  
asymmetrical component of solar daily variability.  $R_1$  or  
 $R_2$  methods for any volume can be used only for  
days when the corresponding regression  $R_1$  or  $R_2$  is  
significant at 2% level where the former is diurnal  
cycle associated with  $R_1$  or  $R_2$ .

So now when a particular  $R_1$  method is  
statistically significant, say at 0.1% level, we  
normally calculate the value of

$$X = \sum_{i=1}^n \left[ \frac{(x_i - \bar{x})^2}{s_i^2} \right]$$

where  $x_i$  is the observed frequency in a particular  
interval,  $s_i$  the reported frequency and  $n$  the number  
of successive observations.  $\bar{x}$  is the total frequency  
divided by the number of intervals. It can be shown

value of  $\chi^2$  is greater than the expected value for ( $n-1$ ) degrees of freedom, then we say that the distribution is not random, but statistically significant at 5% level. Thus when  $\chi^2$  value is high, the preference shown in the histogram for particular variable is considered to be statistically significant.

In computing  $\chi^2$  or % measure of the below codes, we are just trying to see whether they belong to statistically different distributions. For this, the following method is followed:

Let  $f'_0$  and  $f''_0$  represent the observed frequencies &  $f'_1$  and  $f''_1$  represent the expected frequencies in the case of the two distributions for a particular interval, say  $i$ . Then we have the following relations,

$$\text{Total Frequency in case (i)} \quad f_i = \sum_{j=1}^{n_i} f'_{ij}$$

$$\text{Total Frequency in case (ii)} \quad f_i'' = \sum_{j=1}^{n_i} f''_{ij}$$

$$n = n_i + n_i'' = \sum_{j=1}^{n_i} f_{0j}$$

$$\text{Now} \quad f_{0j} = f'_{0j} + f''_{0j}$$

$$f'_{0j} = \frac{n}{n_i + n_i''} f'_{ij}$$

$$f''_{0j} = \frac{n}{n_i + n_i''} f''_{ij}$$

$$f_{0j} = f'_{0j}$$

... (35)

$$\text{The value of } \chi^2 = \sum_{i=1}^{k-1} \left[ \frac{(e_i - f_i)^2}{f_i} \right] \quad \chi^2 = \frac{\sum_{i=1}^{k-1} (e_i - f_i)^2}{\sum_{i=1}^{k-1} f_i}$$

are next calculated. Then the value of  $\chi^2$  is given by the formula

$$\chi^2 = \sum_{i=1}^{k-1} \frac{(e_i - f_i)^2}{f_i} \quad \dots (33)$$

The value of  $\chi^2$  is next tested for significance at  $\alpha$  level for the degrees of freedom ( $k-1$ ) and if the observed value of  $\chi^2$  is greater than the theoretical value, it is concluded that the two distributions tested are different at  $\alpha$  level.

In making this statistical test, there are some points to be remembered. The first one is that the width of the  $i^{\text{th}}$  interval for expected of  $f_i$  should be the same for the two groups concerned. The second point is that the expected frequency  $e_i$  or  $f_i$  should not be less than 5 for any interval. If the expected frequency becomes less than 5, then the or two adjacent intervals are combined together for both the distributions to make the expected frequency  $\geq 5$ . In this case for calculating the degrees of freedom, the true width of  $i^{\text{th}}$  will be considered.

## CHAPTER IV

### TELESCOPE UNITS

#### Introduction

As discussed in chapter III, the author has operated at different times two angle units and a standard unit which has worked almost from 1957 to April 1960. The "newer angle unit" was constructed by the author in 1957 and its general regular recording of cosmic ray intensity by the end of July 1957. The standard unit working at Tschamut since 1955 was repaired and maintained by the author during the period mentioned above. A description of the two units has been already given in chapter III.

The standard unit consists of four standard telescopes of semi-angle of opening  $33^\circ$  and  $37^\circ$  in the east-west and north-south planes respectively. The newer angle unit is having four telescopes of  $5^\circ \times 30^\circ$ , two telescopes of  $10^\circ \times 20^\circ$ , one telescope of  $20^\circ \times 20^\circ$  and a custom telescope of  $45^\circ \times 45^\circ$ , the two angles in each case representing the semi-angle of opening of the telescope in the east-west and north-south planes respectively.

Table 3 shows the number of days during each month on which data are available for the different telescopes. As is evident from the table, the standard

NUMBER OF DAYS IN WHICH HOURS WERE WORKED AUGUST 1957  
TO APRIL 1958 FOR WHICH PAYMENT WAS MADE

unit operated smoothly throughout the period. In the case of the EURECA-2A well, for about three months during 1979-1980, the data are missing as the well was situated in a flooded trench due to ground subsidence caused by landslides in one of the sheltered valleys. Continuous data have been collected since the end of February 1980 from this well.

#### Intensity corrections

The monthly data from each type of telescope have been corrected for instrument response changes as described in Section 3.1. Unfortunately, for about two months (May, June 1980) the monthly exposure data for Tevatron are not available and hence the daily variation studies on a day to day basis for cosmic ray data of this period have not been used. However, the daily mean protonized iron data at CERN, which is a near by sea level station, has shown a very high correlation with the daily mean pressure at Tevatron and hence daily mean pressure of CERN has been used to estimate the relative daily mean intensity of Tevatron.

Temperature correction to the daily variation of cosmic ray intensity has been applied as described in chapter XII (Section 3.2.2). It has been found that the temperature correction to daily variation of iron composition needed in particular in supervisory well, the annual amplitude being of the order of 0.05 % with circ-

~~THE TESTIMONY OF THE WITNESSES~~

The result of normalization of daily rainfall intensity has been described earlier in chapter III, section 3.12. All intensities are summed and to a statistical precision of 100 m. Temperature correction is applied according to Jenson's method using temperature variation up to 10 m level. The temperatures at different location levels were obtained from the three daily meteorological Climate stations located near each of the two reservoirs used for the experiments.

~~TESTIMONY OF EXPERTS~~

It was found to have used a suspended catchment of 40.176 ha as suggested by Gourley et al. and although there may be some other the catchment area extending over 40 ha from the elevation unit of 20 m up. The method suggested by Marbach and Vassiliou [1999] has been adopted in this study.

~~TESTIMONY OF EXPERTS~~

After consideration of the number of parameters of importance of rainfall and by the degree of variability change in the climate variability, the value of parameter  $\alpha$  will be the normalized daily mean intensity of the rainfall component uncorrected for seasonal change for three four days and which is called as deviations

- 27 -

from noon for the preceding intervals of four days. The quoted values of pressure  $\Delta P$  and quoted ray intensity  $\Delta I$  are then found out for each of the four days separately for increasing pressure and decreasing pressure. The values of  $\Delta P$  and the corresponding values of  $\Delta I$  are then plotted as shown in Fig. 15.

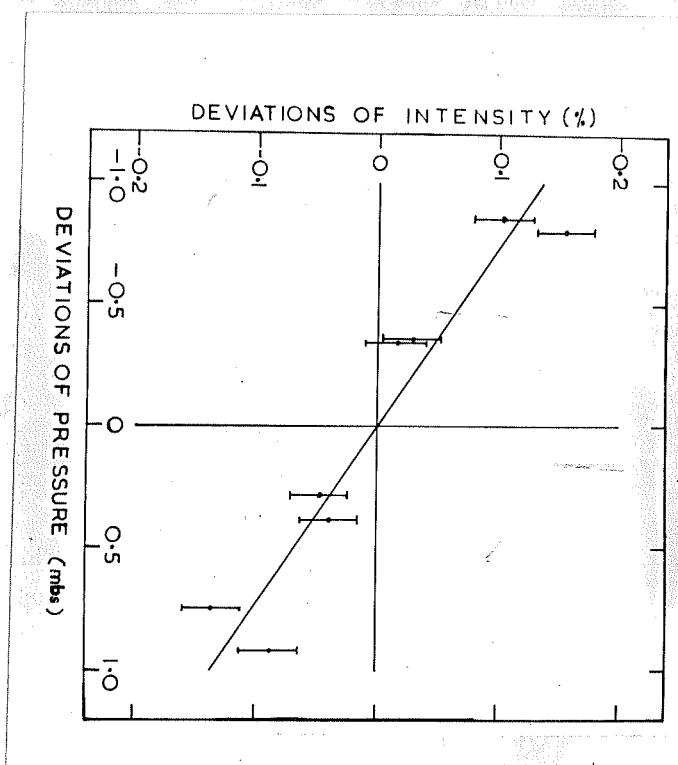


Fig. 15 • Determination of pressure coefficient.

Correlation analysis between  $\Delta P$  and  $\Delta I$  gives a coefficient of -0.95 which indicates high negative correlation between pressure and quoted ray intensity. The best fit

## REGRESSION LINE AND CORRELATION LINE FOR THE DATA.

$$\Delta I = \beta \Delta P$$

Where  $\beta$  the regression coefficient represents the  
regression coefficient. The slope of the regression  
line gives the regression coefficient  $\beta$  as  $-0.16 \text{ l/m}$ .  
This value of the regression coefficient is in full  
agreement with the value  $-0.17 \text{ l/m}$  used for correcting  
the count rate data.

### 4.4

#### DECOMPOSITION OF VARIATION AND ANGLE.

The daily variation of counts may has been  
studied in this chapter in terms of the diurnal and  
continual component of the variation since these  
two are known to represent adequately the main  
features of the variation. The amplitudes and times  
of maxima of the diurnal and continual component  
on individual days (after applying proper meteorolo-  
gical corrections in case of count rates) are represented  
by  $r_1$ ,  $\theta_1$  and  $r_0$ ,  $\theta_0$  respectively and their  
vectorial mean values averaged over an extended period  
or for groups of days selected on the basis of some  
critaria are represented by  $\bar{r}_1$ ,  $\bar{\theta}_1$  and  $\bar{r}_0$ ,  $\bar{\theta}_0$  respect-  
ively.

## DAILY VARIATION

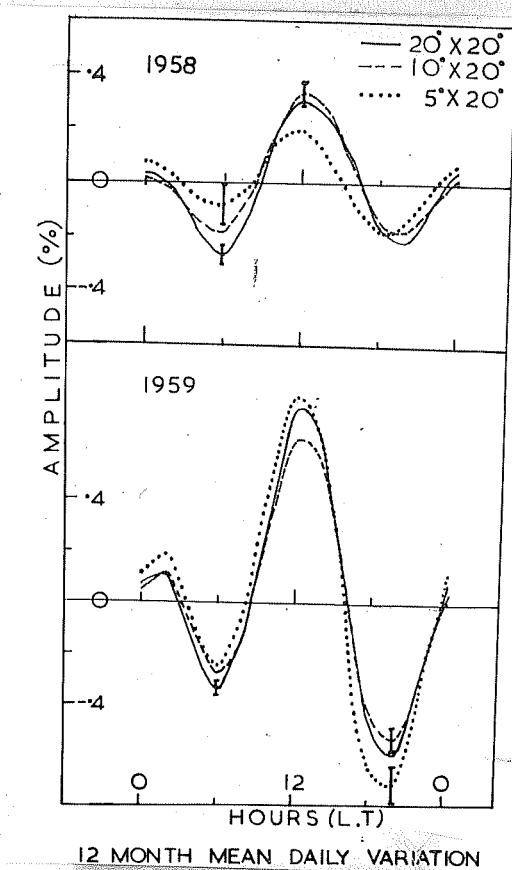
### 4.51 Dependence of daily variation on the

#### angle of opening of the telescope

The three main mean daily variations has been studied with telescopes of different opening angles in the east-west plane and the same angle of opening in the north-south plane. For this study, the telescopes with apertures of opening  $6^\circ$ ,  $10^\circ$  and  $20^\circ$  in the east-west plane and with an angle of opening  $60^\circ$  in the north-south plane have been considered. These telescopes are referred to as  $6^\circ \times 10^\circ$ ,  $10^\circ \times 10^\circ$  and  $20^\circ \times 10^\circ$ .

In figure 14 are presented the 12 month mean daily variations for these three telescopes for the years 1953 and 1959. The amplitude and size of variation of the diurnal and seasonal components of the 12 month mean daily variation recorded by these telescopes are given in table 4. While there is a major difference between the 12 month mean daily variation for 1953 and for 1959, the difference decreases in each year again within statistical accuracy of the observations. It is seen from the figure that the daily variation recorded by all telescopes show a single maximum near noon during 1953. There is a slight increase of a less important additional feature in the early morning at

- 360 -  
about 0200 hours local time for the mean daily variation  
predicted by the above telescope during 1958.



**Fig. 18** - Twelve month mean daily variations  
for telescope of different openings  
in the east-west plane for the years  
1958 and 1959.

The peak to peak amplitude of the daily variation  
in 1959 is about twice the value in 1958 in all the  
telescope. Table 4 gives the amplitudes of daily  
as well as diurnal component of the 12 month mean  
daily variation have increased by a factor of about two  
in 1959 compared to their value in 1958.

Table A

Amplitude and rms of minimum of the diurnal and semi-diurnal components of the 12 month mean daily variation recorded by telescopes of different opening in the east-west plane during the years 1958 and 1960.

| Year | Telescope<br>opening | 1958        | 1960             | 5y  | 7y |
|------|----------------------|-------------|------------------|-----|----|
| 1958 | 5° x 20°             | 0.11 ± 0.04 | 105° 0.13 ± 0.04 | 35° |    |
| 1958 | 20° x 20°            | 0.16 ± 0.03 | 105° 0.17 ± 0.03 | 20° |    |
| 1958 | 50° x 50°            | 0.13 ± 0.03 | 105° 0.09 ± 0.02 | 20° |    |
| 1960 | 5° x 20°             | 0.39 ± 0.04 | 105° 0.40 ± 0.04 | 22° |    |
| 1960 | 10° x 20°            | 0.20 ± 0.03 | 105° 0.30 ± 0.03 | 30° |    |
| 1960 | 20° x 20°            | 0.34 ± 0.03 | 105° 0.45 ± 0.03 | 25° |    |

The present study indicates that during 1958 or 1960 there is no statistically significant difference in the amplitudes of the 12 month mean daily variation recorded by telescopes of different opening angles in the east-west plane.

Fully analogous studies have been made by Saito et al.<sup>107</sup> during 1959-1960 with telescopes of different opening angles in the east-west plane, which resulted in no difference in the amplitude of

The other researches daily variation of the daily variation of the daily variation of the telescope in the east-west plane, Burzyk et al.<sup>220</sup> have studied the daily variation of the telescope ( $\lambda = 32^\circ$ ) with telescopes of different angles of opening  $6^\circ, 10^\circ$  and  $15^\circ$  in the east-west plane and  $10^\circ$  in the north-south plane. They have shown that during 1954-1955 the average amplitude of daily variation increased from 0.6 % to 1.2 % as the angle of opening of the telescope in the east-west plane decreased from  $15^\circ$  to  $6^\circ$ .

Elliot and Poussot<sup>221</sup> who made similar studies at  $\lambda = 32^\circ$  during 1954-1955 found that there was no difference in the average daily variation recorded by the telescopes of different opening in the east-west plane. Their studies at  $\lambda = 32^\circ$  and those of Burzyk et al. at  $\lambda = 14^\circ$  correspond to more or less the same period in 1954-1955.

Elliot and Poussot have suggested that the dependence of the amplitude of daily variation on the opening angle of the telescope is a function of latitude of the place of observation.

Pradella<sup>222</sup> has made systematical with telescopes of  $6^\circ$  and  $30^\circ$  openings in the east-west plane during the high solar activity period of 1957-1958 at the same place as Burzyk et al.

annually. The results have shown that during 1957-1958 there is no change in the amplitude of the 18 month mean daily variation in the longitude of opening of the telescope in the east-west plane is varied from  $30^\circ$  to  $6^\circ$ . These results which agree with the findings at Greenwich indicate that even at the same place the large amplitude of the 18 month mean daily variation associated with narrow angle telescopes during 1957 does not persist during the solar activity minimum period of 1957-1958.

There are two possible explanations for the variations in the findings during 1957-1958 and during 1959-1960. There could be a change of the profile of the primary cosmic ray anisotropy with the solar cycle of activity. Alternatively, the scatter of the time of maximum of the annual component may be much greater for the narrow angle telescope during 1957-1958 than during sunspot minimum period. For a clarification between these two possibilities, it is necessary to study the daily variation on a day to day basis.

In section 3.3B, we have discussed the statistical procedure to allow for differences in the counting rates of instruments when making a comparative study of the histograms of the parameters of the daily variation on a day to day basis. To investigate the day to day changes of daily variation recorded by telescopes of different opening angles in the east-west plane, only the telescopes of  $20^\circ \times 20^\circ$  and  $30^\circ \times 30^\circ$  have been

considered since the standard error is quite large on a diurnally basis for  $10^{\circ} \times 20^{\circ}$  telescopes. The comparative study is made only on days when two telescopes at  $10^{\circ} \times 20^{\circ}$  and one telescope at  $20^{\circ} \times 20^{\circ}$  are working. There are 173 and 230 such days during the years 1958 and 1959 respectively.

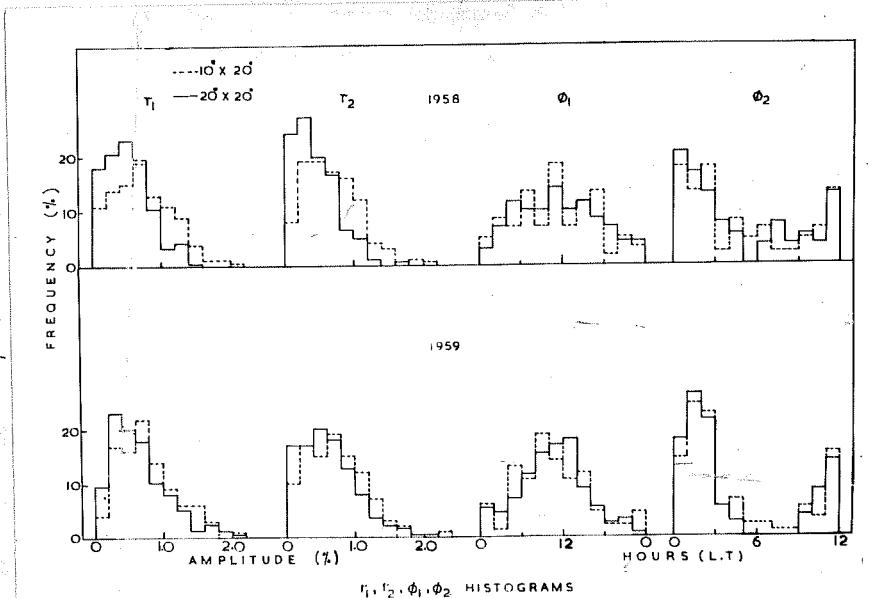


Fig. 17 - Frequency distribution of the amplitudes and times of maxima of the diurnal and semi-diurnal components of the daily variation of cosmic ray intensity recorded by  $10^{\circ} \times 20^{\circ}$  and  $20^{\circ} \times 20^{\circ}$  telescopes on ordinary working days during 1958 and 1959.

Figure 37 shows the distribution of the amplitudes of the diurnal and semi-diurnal components of the solar daily variation during the years 1938 and 1950. The histograms of  $A_1$ , the time of maximum of the diurnal component, are drawn for days on which the amplitude  $P_1$  is significant at the  $2\sigma$  level (or  $P_1 > 2\sigma$ ), where  $\sigma$  is the standard error associated with  $P_1$  or  $P_2$ . Similarly,  $A_2$  histograms showing the distribution of the time of maximum of the semi-diurnal component are drawn for days on which  $P_2$  is statistically significant at the  $2\sigma$  level. An examination of the histograms reveals the following:

- (1) In  $P_1$  and  $P_2$  histograms in both years, there are larger number of days on which the amplitudes are more in  $10^\circ \times 20^\circ$  telescope than in  $50^\circ \times 50^\circ$  telescope. When the histograms of  $P_1$  or  $P_2$  of the two telescopes are compared after putting due allowance for the difference in the counting rates it is found that they do not differ at 6% level of significance. We may, however, conclude that the difference in the amplitudes quoted on individual days is mainly due to the difference in counting rates.
- (2) In the case of both the telescopes, the frequency of occurrence of large values of  $P_1$  and  $P_2$  is more in 1950 than in 1938.

(3) During 1953, the histograms of  $\lambda_1$ , the time of maximum of the diurnal variation, show a random distribution for  $30^\circ \pm 20^\circ$  as well as for  $20^\circ \pm 20^\circ$  telescopes. However, in 1959, the histograms show a clear maximum indicating that on a large number of days the time of maximum of the diurnal component of solar daily variation occurs near noon. The response of the two types of telescopes is in any case similar.

(4) The histograms of  $\lambda_2$ , the time of maximum of the additional component of the solar daily variation, show a preference for early morning hours. This tendency is much more pronounced in 1959 compared to 1953.

**Baranov and Maruyama** <sup>271</sup> studying the daily variation recorded by narrow angle telescopes on individual days during the period 1950-1956 at Abisko and prove show that there is a tendency for the occurrence of the time of maximum of the diurnal variation either at 1100 hours or at 0300 hours local time. Baranov and Maruyama <sup>272</sup> in 1955-1959 find that this periods, but the two periods have shifted, from 0300 hours to 0700 hours and from 2100 hours to 1500 hours respectively. Why have this characteristic the diurnal variation as "a" type and "c" type occurring; on the time of maximum occurs in the early morning or early noon. The structure showing two maxima in the  $\lambda_1$  histograms is rather rare in our station during 1950-1959 at individual day in the diurnal variation at

announced during 1957-1958 by Hansen<sup>51</sup>. It seems that the two nations have merged together to form a broader variation from 2000 hours to 2400 hours in 1957-1958, the period is larger in 1950 than in 1958.

Our study indicates that the continually larger value of the amplitude of 12 month mean daily variation during 1950 is partly due to the amplitude on sunspot days being high and partly due to the smaller variability of the time of passage of the meridian as well as small diurnal component during this period.

Our study also indicates that not only does the amplitude of the daily variation and its rate of maximum change from year to year and from period to period, but the relative decrease in the amplitude of the meridian sunspot angle uncorrected to the obliquity also increases over time.

Hansen<sup>52</sup> has shown that the amplitude of the solar diurnal variation will be larger when measured by a narrow angle telescope than by a wide angle telescope if one source of daily variation is a polar source. Hence we may conclude that during the period of the daily variation during 1954-1958 was comparatively sharp, it is more diffuse and variable during 1957-1958. The possibility of the daily variation during this period being produced by more than one source cannot also be ruled out. Hansen and Red<sup>53</sup> have similarly indicated from directional studies at Ann Arbor during 1957-1958 that on some days the

Daily variation cannot be adequately explained by  
values of seasonal variation and they conclude that  
the daily variation can well be explained by a linear  
trend.

In Fig. 12 the variation of the mean daily variation of the  
mean current measured at stations with telecoastal  
of opacity  $20^{\circ} \pm 37^{\circ}$  during the years 1933 to 1950 is  
shown in figure 12. The variation can be seen from this  
daily variation measured by smaller telecoastal or  
Kondoval and Andreev are also given in the same  
figure. The mean value from daily variation has  
been obtained by the superposition of the diurnal and  
semi-diurnal components of the variations. In table 6  
are presented the values of  $\bar{Y}_1$ ,  $\bar{Y}_2$ ,  $\bar{Y}_3$  and  $\bar{Y}_4$  during  
various years at the three stations. In figure 12  
the changes occurring at the three stations in the  
annual and semi-annual components of the 10 month  
mean daily variation are presented for the period  
1933-1950. An examination of the curves in figure 12  
and the values in table 6 brings out the following  
points.

\* The author is grateful to Mr. F. G. J. L. van der  
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Veen for loaning me the data and to Mr. H. J. J. van  
der Veen for his help in the preparation of the figures.

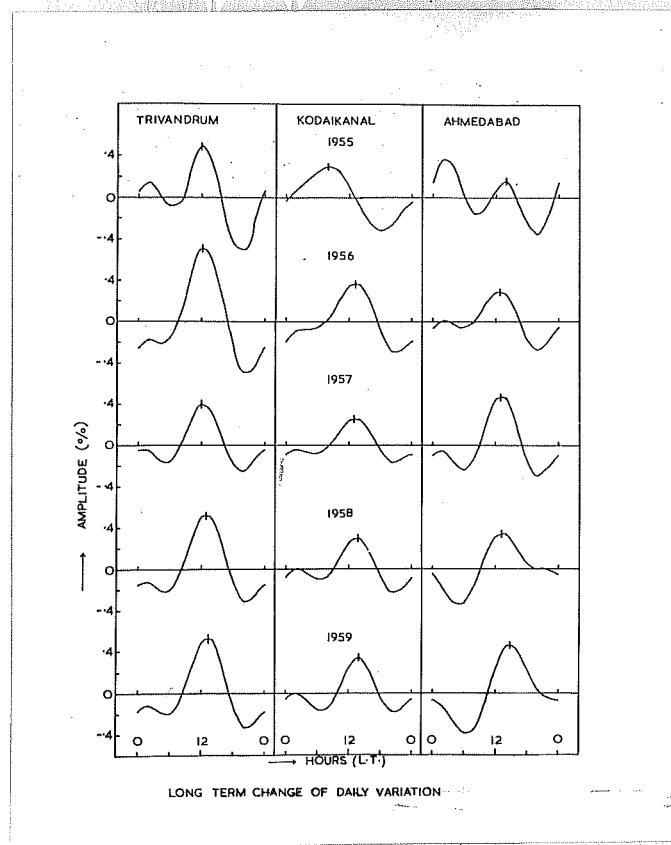


Fig. 20 • 12 month mean daily variation observed at three stations  
at Trivandrum, Kodaikanal and  
Ahmedabad during the years  
1955 to 1959.

#### (a) AMPLITUDE

(1) During the years 1955, 1956, 1957 and 1958, the principal minimum of the twelve month mean daily variation occurs at 0000. From 1959 to 1960, there is a slight shift of the time of minimum from 12 hours to 14 hours local time.

The secondary daily variation occurs at 0300 hours in 1956 becomes less prominent in 1957 and 1958. In 1957,

Table 5

| Sample |
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| 1      |
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| 85     |
| 86     |
| 87     |
| 88     |
| 89     |
| 90     |
| 91     |
| 92     |
| 93     |
| 94     |
| 95     |
| 96     |
| 97     |
| 98     |
| 99     |
| 100    |

The 18 month mean daily variation had a single maximum at noon. In 1950, there is a slight deviation of the superposition of the two prominent early morning maxima, peak to peak amplitude of the annual daily variation also shows changes during different periods. The amplitude is maximum during 1953 and minimum during 1957. In 1955, 1959 and 1960 the peak to peak amplitude is practically the same.

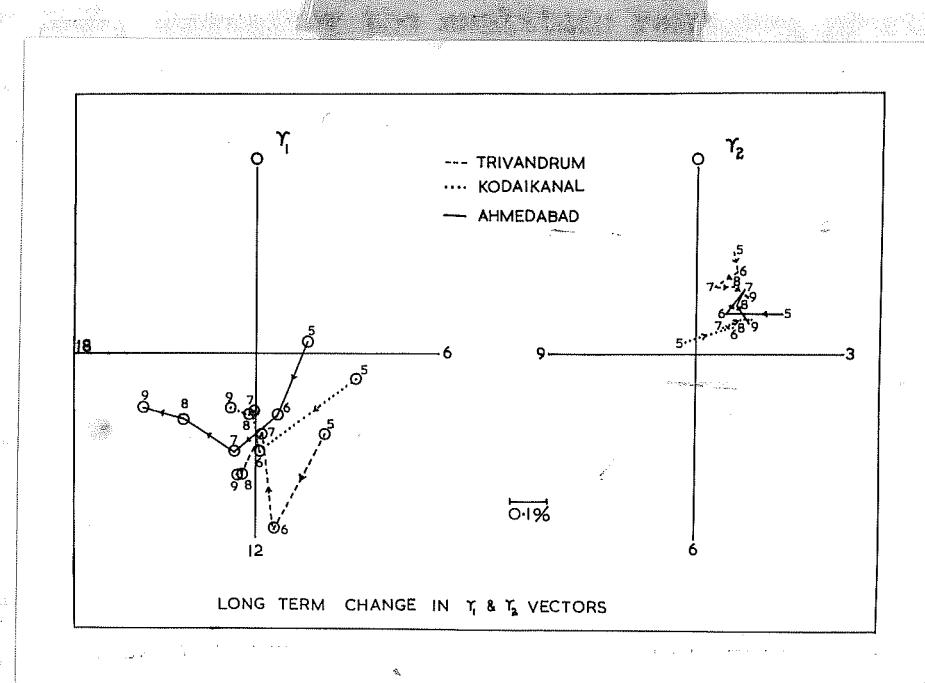


Fig. 20 → Diurnal and seasonal variation of the 18 month mean daily variation observed with zenithal thermometers at Trivandrum, Ko. and Ahmedabad during the years 1950 to 1960.

(3) The annual amplitude  $\bar{A}_1$  gradually decreases from June 1955 and reaches its minimum value in June 1957. Thereafter it shows a slow increase and attains a value of 0.30  $\mu$  by the end of 1957. No statistically significant changes are found to occur after that.

(4) There is a gradual decrease in the amplitude  $\bar{B}_1$  of the longitudinal component from 1955 to 1957, thereafter the amplitude remains practically constant.

(5) The time of maximum of the 12 month mean annual component moves significantly later from 1955 to 1956 mainly due to the change occurring in the latter half of 1955, but there is also a very gradual change in the time of maximum.

(6) There is a gradual shift in the time of maximum  $\bar{B}_2$  of the longitudinal component also, but this cannot be established at a level of statistical significance.

#### (6) Annual

(1) The 12 month mean daily variation  $\bar{A}_{12}$  is highly decadal in character with the time of maximum occurring at about 0600 hours local time. The rate of daily variation shows the time of maximum shifting to 2400 hours during 1955. Thereafter there is

- 15 -

not much change in the form, mode or peak amplitude  
or the time of arrival of the average daily variation  
from year to year.

(2) The annual amplitude  $\bar{x}_1$  decreases during  
1951-1956 but from 1957 onwards there is no change.

(3) Just as in the case of variation, the  
major shift in the time of maximum  $\bar{x}_1$  of the 12 month  
mean diurnal component occurs during 1951-1956. The  
shift of  $\bar{x}_1$  is larger at 0000 station than at the other  
two.

(4)  $\bar{x}_1$  is negligible in 1953, but increases  
and reaches comparison with  $\bar{x}_2$  in 1958.

(5) Shift in  $\bar{x}_1$  is slightly larger than in  
the other two stations.

(6) ~~all stations~~

(1) During 1956 the average daily variation  
shows two distinct maxima, one at 0000 hours and the other  
at 1400 hours local time. It is seen that the early  
morning maximum is much more prominent than the maximum in  
the afternoon. In 1958, the early morning maximum has  
completely disappeared and from 1957 onwards the average  
daily variation is represented by a curve with a single

~~maximum at 0000 hours local time.~~

(3)  $\bar{Z}_1$  increased gradually from 0.34% in 1955 to 0.53% in 1960.

(4)  $Z_2$  showed a shift to later hours and the shift is quite large from 1955 to 1960.

(5) The changes in the time of arrival  $Z_2$  of the annual mean component of the annual daily variation during 1955 - 1960 are comparatively small. At Tsimshian and Anchorage the time of arrival  $Z_2$  is always found to occur between 0000 hours and 0000 hours.

To go over there is comparative study of the 12 month mean daily variation for the three stations dependent here, and at all the three stations the form of the annual daily variation has not changed much during the period 1957 - 1960. The amplitude at Anchorage and Tsimshian was comparable, but at Kotzebue it is significantly greater. The time of arrival seems to compare at all three stations. In the years 1955 - 1956, the differences were that the time was later at Kotzebue and earlier at Tsimshian.

The long term changes in the time of daily variation are considered to be influenced by a number of factors which have been discussed in detail by Hines and Hines (1960). The changes in the time of arrival  $Z_2$  of the annual mean component of the annual daily variation during 1955 - 1960 are shown in Figure 10. From the figure it is seen that the time of arrival  $Z_2$  has shifted to later hours during 1955 - 1960.

It is also seen that the time of arrival  $Z_2$  during the annual mean component of the annual daily variation for a large

number of stations both for moon and sun tide component of semidiurnal rays have shown that at the equatorial stations of Trivandrum, Pondicherry and Calicut the semidiurnal component is dominant, but the diurnal component at Cochin ( $\lambda = 45^\circ$ ) they have shown that the diurnal component is mainly due to precession effect and was printed 1955 - 1956.

In figure 20, the vectors representing the semidiurnal component of the sun tides and the variation of pressure all of these intensity corrected for the diurnal variation of tides and pressure, recorded at Ahmedabad for the period August 1957 to April 1958, and those at Alappuzha for the period 1955 to 1956, are shown. It is quite evident from the figure that the pressure vector is  $90^\circ$  out of phase with the

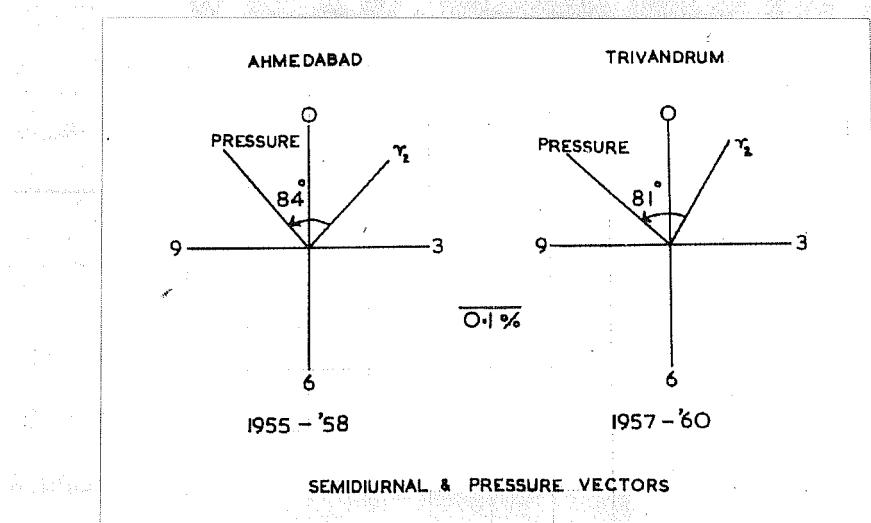


Fig. 20 - Semidiurnal & pressure vectors for the variation of pressure and variation of semidiurnal intensity vectors for the period 1955-1958 for Ahmedabad and for the period August 1957 - April 1958 for Trivandrum.

comes per vector and hence the maximum change  
in rate due to a possible variation of  $\alpha$  is  
to occur from our estimate that there could possibly  
be the same ratio between the two vectors as  
between the two components of the variation of  $\alpha$ .  
Thus if  $\alpha$  varies by  $\pm 1^\circ$  the variation in the  
rate of variation of the two vectors will be  
 $\pm 1^\circ$  and hence the variation in the variation  
of a perpendicular component of the daily variation  
will be half of the variation of  $\alpha$ .

We have already seen in section 4.3 that the  
1950 and 1952 mean daily variation reported by  
 $80^\circ \times 20^\circ$  telescope at Cambridge is larger than the  
1950 and 1952  $20^\circ \times 20^\circ$  and hence this change in  
observing angle is estimated to be about  $\pm 1^\circ$ .  
We have assumed the 1950 mean daily variation  
reported by  $80^\circ \times 20^\circ$  telescope during 1950 and 1952  
as Alfvénic also. The average daily variation is  
given in Table 21 and the percentage of the harmonic  
components of 30 which were daily variation are given  
in Table 22.

It is seen that at both places the variation  
of the 30 sec component of the 12 min mean daily  
variation has increased in 1952 and the time of maximum  
is shifted to October while it was in the end of November  
in 1950. The diurnal component of intensity remained  
fixed at  $20^\circ \times 20^\circ$  telescope in all four years significantly  
larger at Cambridge than at Alfvén.

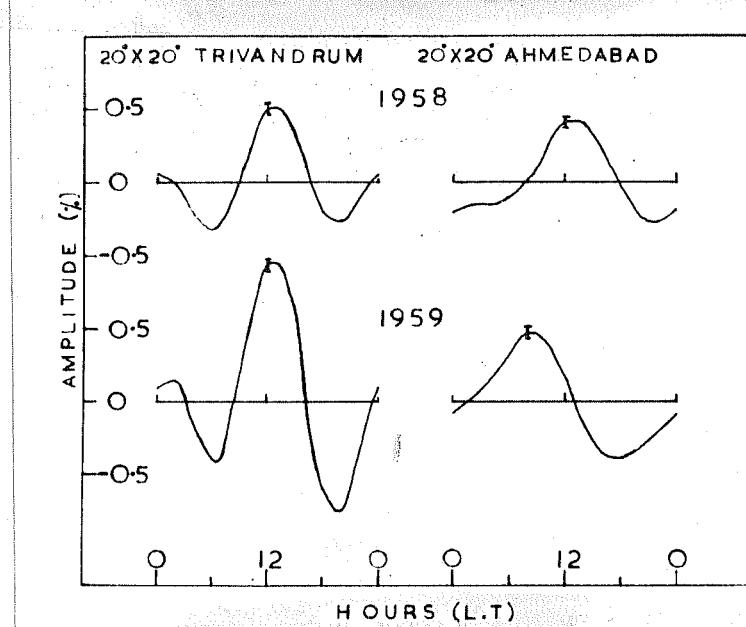


Fig. 11 - 12 hour mean daily variation  
recorded by 20° X 20° Volometers  
during 1958 and 1959 at Trivandrum  
and Ahmedabad.

The magnitude of the 12-hour mean daily variation of the 12-hour mean daily component of the 12-hour mean daily variation recorded by 20° X 20° Volometers is given below for the years 1958 and 1959.

| Station    | $\bar{v}_1$ | $\bar{v}_2$ | $\bar{v}_3$ | $\bar{v}_4$ | $\bar{v}_5$ |
|------------|-------------|-------------|-------------|-------------|-------------|
| Trivandrum | 1.02        | 0.38 ± 0.03 | 20°         | 0.33 ± 0.03 | 25°         |
| Kochi      | 1.02        | 0.34 ± 0.03 | 20°         | 0.33 ± 0.03 | 25°         |
| Calicut    | 1.02        | 0.34 ± 0.03 | 20°         | 0.35 ± 0.03 | 25°         |
| Ahmedabad  | 1.02        | 0.32 ± 0.02 | 20°         | 0.33 ± 0.02 | 25°         |

4.23 A comparative study of daily variation  
of mean intensity of ionization  
radiation.

For illustrating the result of comparison of mean daily variation recorded by identical instruments at the places examined, only the mean daily variation is of great importance. Individual and systematic errors may have conditions in an arbitrary extent. The two stations are close to the sea and atmospheric levitation has no appreciable influence. It is seen from table I that the mean daily average variation at the two places is about the same (17.38 GV at Kanyakulam and 17.10 GV at Trivandrum). But the stations differ only as regards by 800 miles and hence the difference observed in the daily variation may be attributed to the difference in the sensitivity function applicable to the two instruments.

The study the mean rate of daily variation recorded at the two stations with stations located at buildings of spans 20' x 5' for the period 1952-1953. Common days with all the four valances of the window at each station were selected have been considered. This is to ensure the following:-

- (a) Uniformity of counting rate of the instrument at each station. This facilitates the application of statistical tests as described in section 3.23 to judge differences for the systematic deviations in counting rates at the two stations.

(2) Since in the presence of day to day variability of the daily variation, the comparison between the two stations may be valid.

A total of 100 days have been chosen randomly during 1955. The set of the results of the average daily variation of mean density for one day corrected for atmospheric change of temperature and pressure is given in figure 10 for Rosalind and Travancore. The parameters  $\bar{T}_1$ ,  $\bar{T}_2$ ,  $T_1$  and  $T_2$  are given in table 7.

Table 7

Table showing the parameters  $\bar{T}_1$ ,  $\bar{T}_2$ ,  $T_1$ ,  $T_2$  of the average daily variation for 100 cases and the all the 4 telegrams of the standard ratio are varying at Rosalind and Travancore.

| Station    | $\bar{T}_1(0)$  | $T_1$         | $\bar{T}_2(0)$  | $T_2$        |
|------------|-----------------|---------------|-----------------|--------------|
| Rosalind   | $0.10 \pm 0.03$ | $174^{\circ}$ | $0.09 \pm 0.02$ | $35^{\circ}$ |
| Travancore | $0.27 \pm 0.02$ | $170^{\circ}$ | $0.21 \pm 0.02$ | $22^{\circ}$ |

It is seen from figure 10 and from table 7 that the average daily variation has a small value at Travancore and a large value at Rosalind, which may be due to larger  $T_1$  on standard days at Travancore.

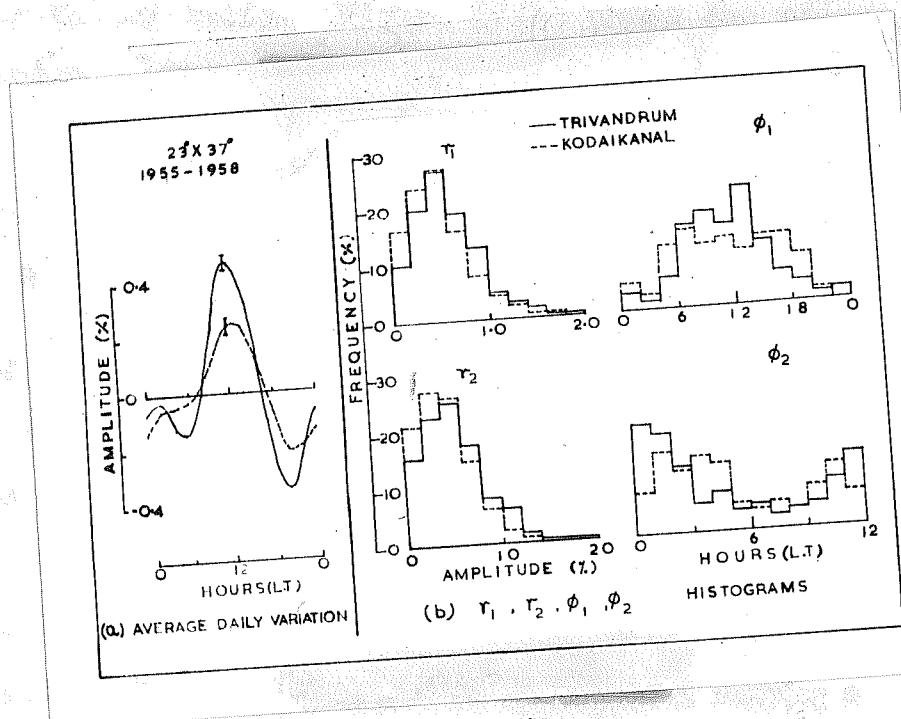


Fig. 22 - (a)

Average daily variation of moon intensity corrected for atmospheric changes of temperature and pressure for 333 working days when all 4 telescopes of the standard unit are working at Kodaikanal and Trivandrum during the period 1955 - 1958.

(b)

Frequency distributions of T<sub>1</sub>, T<sub>2</sub>,  $\phi_1$ ,  $\phi_2$  for 333 working days of the standard units at Trivandrum and Kodaikanal during the period 1955-1958.

or to  $\phi_2$ , the time of minimum of the diurnal component, having large variability at Kodaikanal. To get some knowledge about this, the amplitude and time of minimum

of the diurnal and antidiurnal components are studied on a day to day basis. Figure 22(b) shows the histograms of  $\chi_1$  and  $\chi_2$ . The histograms of  $\chi_1$  and  $\chi_2$  for days when the corresponding amplitudes  $\beta_1$  and  $\beta_2$  are significant at the  $2\sigma$  level are also shown in the same figure.

Applying statistical tests to the amount of the counting rate at zenithal being 10 times that at Equator we find that the  $\chi_2$  as well as  $\chi_1$  histograms are not significantly different at the two places. On the other hand, both  $\chi_1$  and  $\chi_2$  histograms have significantly greater variability at Rohtak than at Trivandrum. This is probably due to the telescopes at Rohtak having a lower mass primary energy response than at Trivandrum. This hypothesis has been tested by drawing a polar histogram of  $\chi_{1g} - \chi_{2g}$  where  $\chi_{1g}$  coincide with 0° direction. Here  $\chi_{1g}$  and  $\chi_{2g}$  represent the time of maximum of the diurnal component of daily variation at Trivandrum and Rohtak on a particular day. Only those days which have an amplitude  $\beta_1$  at both places significant at the  $2\sigma$  level are considered. For as such days, the polar histogram is shown in Figure 23.

It is seen that  $\chi_1$  values at Rohtak are less randomly distributed in relation to the corresponding value on the particular day at Trivandrum. There is a marked tendency for the Rohtak time of maximum to occur 3 to 4 hours earlier than the maximum at Trivandrum.

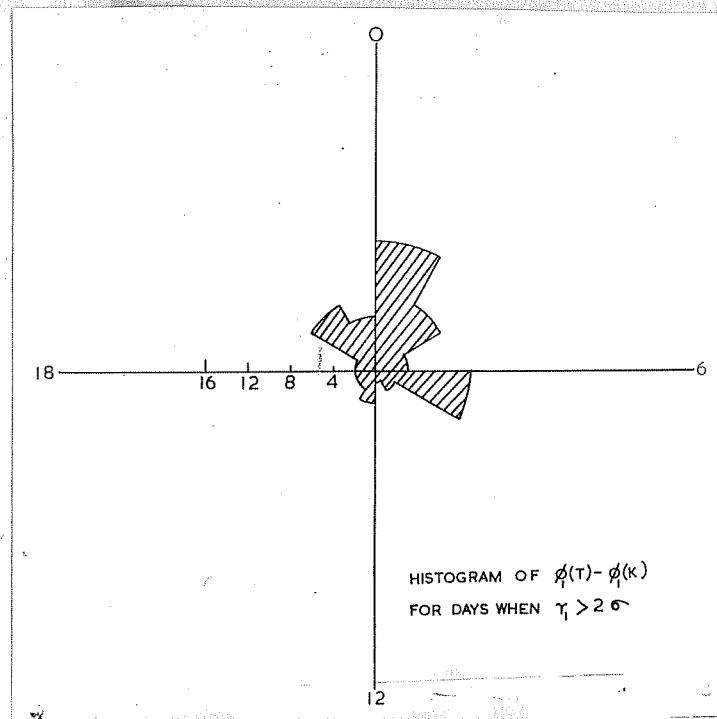


Fig. 23 - Polar histogram of  $\phi_T - \phi_K$  for  
50 days when  $\gamma$  exceeded 2.6  
from 2000 to 20000 units of  
radiation and magnetism during  
1963 - 1970.

radio noise should be captured in the daily variation is the result of the anisotropy of primary cosmic rays of positive charge such as protons with suffer deflection in geomagnetic field.

$\gamma_{\text{iso}}$  has shown by a comparative study in east and west Australia that the daily variation for west which corresponds to a lower zenith angle of primaries than east has

are a greater variability than for the rest. The comparative study of daily variation is summarized and it is shown that the daily variations from the three groups are different.

#### 4.3.1 Relationship of solar daily variation with geomagnetic activity.

In order to study the relation between the solar daily variation of solar intensity and geomagnetic disturbance characterized by the Interplanetary Convection Figure 6, the days are divided into three groups, depending upon the value of  $C_p$ . The following symbols have been used to indicate the three groups.

| Days                         | Symbol                                | Value of $C_p$          |
|------------------------------|---------------------------------------|-------------------------|
| Low geomagnetic activity.    | $\frac{1}{2}$                         | $0 \leq C_p \leq 0.5$   |
| Medium geomagnetic activity. | $\frac{1}{2} \frac{1}{2}$             | $0.5 \leq C_p \leq 1.0$ |
| High geomagnetic activity.   | $\frac{1}{2} \frac{1}{2} \frac{1}{2}$ | $C_p \geq 1.0$          |

Two units recording sunspot area, i.e., the standard unit with telescope of conical angle opening  $23^\circ \times 37^\circ$  and the narrow angle telescope of conical angle opening  $10^\circ \times 30^\circ$ , were selected for this study. The entire period of August 1957 to April 1 1958 is considered together. The mean daily variations for the two types of telescopes and for the three groups during this period are shown in Figure 24.

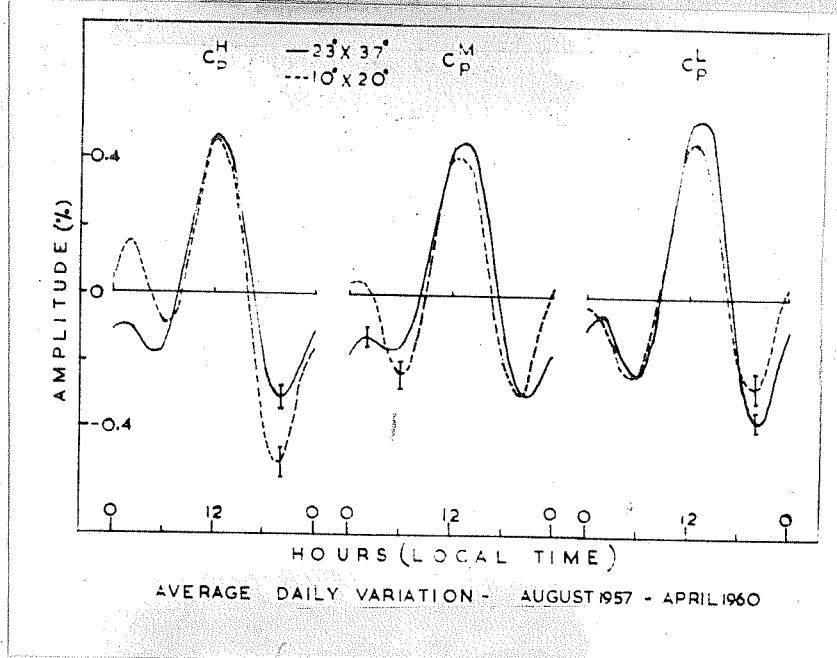


Fig. 24 - Average daily variation recorded by  $23^\circ \times 37^\circ$  teleneopps and  $10^\circ \times 20^\circ$  teleneopps on  $C_P^H$ ,  $C_P^M$  and  $C_P^L$  groups of days during the period August 1957 to April 1960.

It is seen from the figure that through the peak-to-peak amplitude of the average daily variation is slightly more in  $10^\circ \times 20^\circ$  than in  $23^\circ \times 37^\circ$  for  $C_P^H$  group of days, the overall time mean daily variation curves for both  $C_P$  group of days are similar in form and almost of comparable peak-to-peak amplitude for the two teleneopps studied. There is no significant difference in the peak-to-peak amplitudes of the average daily variation for geomagnetically quiet and disturbed days ( $\phi_1$  and  $\phi_2$  groups). The only morning maximum in the average daily variation appears only in the case of  $C_P^H$  and the  $10^\circ \times 20^\circ$  teleneopp.

Table 8 gives the parameters  $\bar{E}_1$ ,  $\bar{Z}_1$ ,  $\bar{E}_2$  and  $\bar{Z}_2$  of the diurnal and semi-diurnal components of the average daily variation for the three groups of days  $C_p^1$ ,  $C_p^M$  and  $C_p^H$  for the two telescopes during the period August 1957 - April 1958.

Table 8

Average parameters  $\bar{E}_1$ ,  $\bar{Z}_1$ ,  $\bar{E}_2$  and  $\bar{Z}_2$  of the diurnal and semi-diurnal components of daily daily variations recorded at zenith for days grouped according to the degree of geomagnetic disturbance. Period August 1957 - April 1958.

| Group   | Telescope $10^\circ \times 30^\circ$ |             |                       | Telescope $10^\circ \times 20^\circ$ |                       |             |                       |             |
|---------|--------------------------------------|-------------|-----------------------|--------------------------------------|-----------------------|-------------|-----------------------|-------------|
|         | $\bar{E}_1(\text{S})$                | $\bar{Z}_1$ | $\bar{E}_2(\text{S})$ | $\bar{Z}_2$                          | $\bar{E}_1(\text{D})$ | $\bar{Z}_1$ | $\bar{E}_2(\text{D})$ | $\bar{Z}_2$ |
| $C_p^1$ | 0.3120.02                            | $170^\circ$ | 0.1920.02             | $21^\circ$                           | 0.3220.03             | $143^\circ$ | 0.3220.03             | $31^\circ$  |
| $C_p^M$ | 0.3120.02                            | $125^\circ$ | 0.1920.02             | $45^\circ$                           | 0.3120.03             | $120^\circ$ | 0.3220.03             | $27^\circ$  |
| $C_p^H$ | 0.3120.02                            | $145^\circ$ | 0.0820.02             | $35^\circ$                           | 0.3220.03             | $105^\circ$ | 0.3220.03             | $15^\circ$  |

It is seen from the table that there is no significant change in the diurnal variation in the three groups of days. There is a significant change in the time of maximum of the diurnal component for  $C_p^1$  group of days; it being earlier by about 3 hours compared to  $C_p^H$  group of days in the  $10^\circ \times 20^\circ$  telescope. This shift to earlier

hours of  $A_1$  in  $C_1^1$  group is not statistically significant in the case of the  $25^\circ \times 37^\circ$  balance.

Figure 23 shows the histograms of the amplitudes of diurnal and annual components of daily variation  $x_1$  and  $x_2$  for the time  $C_1^1$  group of days studied for the two units  $25^\circ \times 37^\circ$  and  $20^\circ \times 30^\circ$ . The histograms for  $A_1$  and  $A_2$  for time of minimum of the diurnal and annual components of daily variation have been drawn for days on which the corresponding amplitudes  $x_1$  and  $x_2$  are significant at the 0.6 level. It is seen from the figure that:

- (1) The  $x_1$  histograms for the time  $C_1^1$  group of days do not differ significantly from each other;
- (2)  $x_2$  histograms do not differ significantly for the  $C_1^1$ ,  $C_1^2$  and  $C_2^1$  groups;
- (3)  $A_1$  histograms for  $C_1^1$  and  $C_1^2$  groups for  $25^\circ \times 37^\circ$  balance differ significantly. Statistical tests made to see whether the two distributions are different have confirmed this difference. Thus the diurnal variation of mean intensity exhibits on a day to day basis for a large number of days show that there is a tendency of the occurrence of the time of minimum  $A_1$  of the diurnal component earlier on geographically disturbed days ( $C_1^1$  group) than on quiet days ( $C_1^2$  group). The earlier time of maximum of the diurnal variation on geographically disturbed days is shown more clearly in the narrow angle telescope of  $30^\circ \times 30^\circ$  than in  $20^\circ \times 30^\circ$ . In the time averaged daily variation

station for geomagnetically quiet and disturbed days, we have seen that the time of maximum of diurnal variation is  $\approx$  3 hours earlier for  $C_p^S$  than for  $C_p^L$  group in  $10^\circ \times 20^\circ$  telescopes.

(4) % Histograms showing the distribution

of the time of maximum of the diurnal component of daily variation reveal statistically significant preference for early morning (0 - 3 hours) in all the cases.

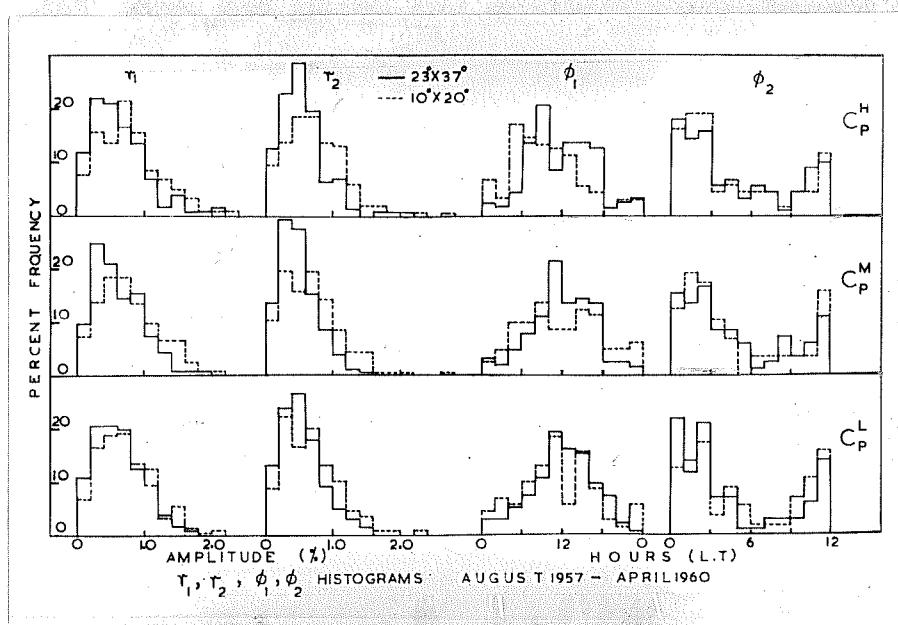


Fig. 23 - Frequency distribution of the amplitudes and times of maxima of the diurnal and semi-diurnal components of daily variation for  $C_p^S$ ,  $C_p^M$  and  $C_p^L$  groups of days recorded by  $25^\circ \times 37^\circ$  and  $10^\circ \times 20^\circ$  telescopes at Ulyanovsk during the period August 1957 to April 1960.

The relation between daily variation of cosmic ray intensity and geomagnetic disturbance as characterized by  $K_p$  or  $C_p$  has been studied by many workers. Sasaki and Kodama<sup>210</sup> and many others have found that the amplitude of solar diurnal variation becomes larger and the time of maximum shifts to earlier hours on geomagnetically disturbed days compared to geomagnetically quiet days. The dependence of the amplitude of diurnal variation and its time of maximum on  $C_p$  has been studied by many workers during the period 1935 - 1950 both for proton and nucleonic components at the low latitude stations of Kodama<sup>211</sup>, Tsurumi<sup>212</sup> and Apia<sup>213</sup>. Satyapal<sup>214</sup> from a study of the nucleonic component during 1957, recorded at Kodama, finds that there does not exist any distinct relationship between daily variation parameters and geomagnetic activity represented by  $K_p$  or  $C_p$ . Studying mean data from Kodama during 1936, Gandy<sup>215</sup> has also got similar results. He finds that for geomagnetically quiet days, the amplitude of diurnal variation is more than on geomagnetically disturbed days. This is not in agreement with the results of Sanderson<sup>216</sup> from a study of proton component at a high latitude station during 1947 - 1950. He has reported an increase in  $R_1$  and a shift of  $K_1$  to earlier hours with increasing  $C_p$ .

The difference between the results relating to 1947 - 1950 and 1936 - 1957 is understandable on the observations of Kodama<sup>217</sup> that the disturbance vector, which defines the vector representing the quiet period

diurnal variation, has itself an 11 year period of change in amplitude as well as time of maximum. Thus the amplitude and phase changes observed on  $C_p$  days will depend upon the position of solar activity.

The results presented in this section are in agreement with the ~~the~~ time of maximum of the diurnal component of solar daily variation occurring earlier on geomagnetically disturbed days, as suggested by Saito and Kondo<sup>210</sup>, Sanderson<sup>211</sup> and others.

However, during August 1957 - April 1960, the amplitude of the averaged solar diurnal variation does not appear to depend very much on the value of  $C_p$ .

## CHAPTER V

## DAY TO DAY CHANGES OF SUNSPOTS

5.1 Introduction.

A startling series of day to day changes of intensity of sunspot groups takes place during cosmic ray storms. Many cosmic ray storms have been associated with type IV bursts and sudden commencement magnetic storms, but often the solar relationship is not precisely known. Dorrus<sup>6</sup> has classified the solar streams into two kinds. According to him, streams of the first kind are fuzzified and are probably connected with faculae or with high latitude formations on the sun, which do not cause an appreciable amount of decrease in cosmic ray intensity. Such streams are likely to cause cosmic ray storms with gradual onset and of long duration with a marked recurrence tendency. Streams of the second kind which are very dense probably originate within the sun's equatorial region and are most probably intimately connected with sunspots. They are likely to cause large sudden commencement magnetic storms of short duration which show a very pronounced tendency. Török believes that storms of great density causing large cosmic ray storms carry strong magnetic fields ( $H \sim 10^{-6}$  gauss) whereas the streams of the first kind carry weak fields ( $H < 10^{-5}$  gauss).

The discovery of polar cap disruption events (PCD) provided a markedly unambiguous interpretation of a sequence

of events starting from an event on the sun followed by its effects in the solar system, represented here in cosmic rays. This study, therefore, starts by considering solar and planetary events and their relationship with cosmic ray changes. The three events which are important in producing cosmic ray change, the cosmic ray intensity changes are studied at various latitudes as related with variation both over as well as upon solar cycles in order to get an idea of the mutual relationship of the oscillating periods and the change of time of approach is presented.

The study of the day to day changes has also been continued.

(1) Looking upon cosmic ray changes as a clear indication of their association with solar events, the July 1929 event which is prominent for the intensity and duration is studied separately.

(2) Considering events where the maximum intensities of cosmic rays are attributed to magnetically triggered as reported by the C.R.C. This is indicative of an indication of the possibly early detection.

#### 4.2. The Data

The cosmic ray data is derived from stations in two different latitude belts. In Table I are listed the stations, the data of which are used in the analysis.

|   | 1 | 2 | 3 | 4 | 5 | 6 | 7 | 8 | 9 | 10 | 11 | 12 | 13 | 14 | 15 | 16 | 17 | 18 | 19 | 20 | 21 | 22 | 23 | 24 | 25 | 26 | 27 | 28 | 29 | 30 | 31 | 32 | 33 | 34 | 35 | 36 | 37 | 38 | 39 | 40 | 41 | 42 | 43 | 44 | 45 | 46 | 47 | 48 | 49 | 50 | 51 | 52 | 53 | 54 | 55 | 56 | 57 | 58 | 59 | 60 | 61 | 62 | 63 | 64 | 65 | 66 | 67 | 68 | 69 | 70 | 71 | 72 | 73 | 74 | 75 | 76 | 77 | 78 | 79 | 80 | 81 | 82 | 83 | 84 | 85 | 86 | 87 | 88 | 89 | 90 | 91 | 92 | 93 | 94 | 95 | 96 | 97 | 98 | 99 | 100 |
|---|---|---|---|---|---|---|---|---|---|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|----|-----|
| 1 | 1 | 2 | 3 | 4 | 5 | 6 | 7 | 8 | 9 | 10 | 11 | 12 | 13 | 14 | 15 | 16 | 17 | 18 | 19 | 20 | 21 | 22 | 23 | 24 | 25 | 26 | 27 | 28 | 29 | 30 | 31 | 32 | 33 | 34 | 35 | 36 | 37 | 38 | 39 | 40 | 41 | 42 | 43 | 44 | 45 | 46 | 47 | 48 | 49 | 50 | 51 | 52 | 53 | 54 | 55 | 56 | 57 | 58 | 59 | 60 | 61 | 62 | 63 | 64 | 65 | 66 | 67 | 68 | 69 | 70 | 71 | 72 | 73 | 74 | 75 | 76 | 77 | 78 | 79 | 80 | 81 | 82 | 83 | 84 | 85 | 86 | 87 | 88 | 89 | 90 | 91 | 92 | 93 | 94 | 95 | 96 | 97 | 98 | 99 | 100 |

The particulars of the location and the name of the principal investigator at each station from which transients the data have been made available to us are also indicated in the table. The count rates and intensities according to Gurney and Hobson's method and the mean energy of emission for the component calculated by using the coupling coefficients given by Fermi and adopting the method suggested by Fuchs are also indicated in the same table.

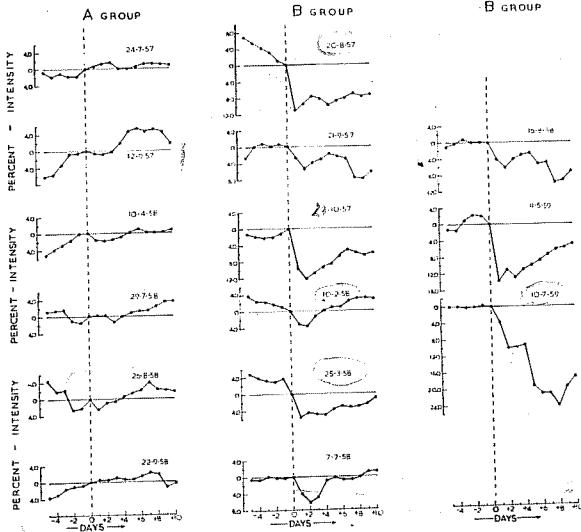
### 6.3 Polar cap absorption events.

Reid and Leinweber<sup>174</sup> have listed 24 intense polar cap absorption events recorded during the period May 1957 to July 1958. Of these, counts and intensity data at Oulu are available for 21 events. At three events on August 20, 1957; September 6, 1957; September 20, 1957; August 21, 1958; July 14, 1958 and July 16, 1958, which are preceded in each case by other PCA events within about 10 days have been eliminated from further consideration since we are interested in studying the day to day changes upto 10 days following the onset of the comic ray storms associated with some of these events.

Figure 38 illustrates the changes in the intensity as observed by the neutron monitor at Olmax for fifteen of these events. Olmax has been chosen for the initial classification of PCA events in terms of their association

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**POLAR CAP ABSORPTION EVENTS**  
NEUTRON INTENSITY CHANGES AT CLIMAX



**Fig. 26 - DAILY MEAN NEUTRON INTENSITY CHANGES FROM -5 TO +20 DAYS AT CLIMAX CORRESPONDING TO EACH POLAR CAP ABSORPTION EVENT. ALL INTENSITIES ARE SHOWN AS DEVIATIONS FROM THE VALUES ON CRASH DAY.**

with cosmic ray changes since all stations considered, Climax neutron monitor has the lowest mean energy of response and is, therefore, expected to respond most effectively to changes of particle intensity.

six events are found not to be associated with geomagnetic decreases exceeding 12.5 % at Elsinor following them within one or two days. These are referred to as 'group A' events. The other 9 events are followed by geomagnetic intensity decreases of the Faraday type exceeding 3 % at Elsinor within one or two days. These are referred to as 'group B' events in this section.

The polarimetric data of the FGA events, the probable dates and the magnetic storm are shown in table 10 for group A and group B events separately. Once analysis has been done for the following intensities from -5 to +30 days for A and B groups separately,

(1) Daily mean intensity.

(a) Mean intensity at Elsinor.

corrected for instrumental response  
and atmospheric temperature changes.

(b) Mean intensity at Churchill and  
Hermitsund corrected for instrumental  
response changes.

(c) Intensity intensity at Huxley, Clinch,  
Chicago, Ibadan, M. Wellington, Honolulu,  
Lisbon and Durban corrected for  
instrumental changes.

(d) If the mean annual intensity of the earth's  
magnetic field at Robinson.

中華人民共和國農業部令 第二號

10. The following table shows the number of hours worked by each employee in a company.

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• 93 • 101 • 102 • 103 • 104 • 105 • 106 • 107 • 108 •

1 2 3 4 5 6 7 8 9 10 11 12 13

卷之三十一

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Journal compilation © 2006 British Ecological Society,  
*Journal of Animal Ecology*, 75, 103–112

10. *Leucosia* (L.) *leucostoma* (L.) *leucostoma* (L.) *leucostoma* (L.) *leucostoma* (L.)

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2022-2023  
2023-2024

10. The following table shows the number of hours worked by 1000 employees in a company.

1 2 3 4 5 6 7 8 9 10 11 12 13 14 15 16 17 18 19 20

0 1 2 3 4 5 6 7 8 9 0 1 2 3 4 5 6 7 8 9 0 1 2 3 4 5 6 7 8 9

the same time, the author has been able to make a number of observations which will be of interest.

卷之三

(d) Ionospheric absorption at 25 Mc/s at

Guayaquil.

(e) Solar flux  $\eta$ , indices (only from +5 to +5 days).

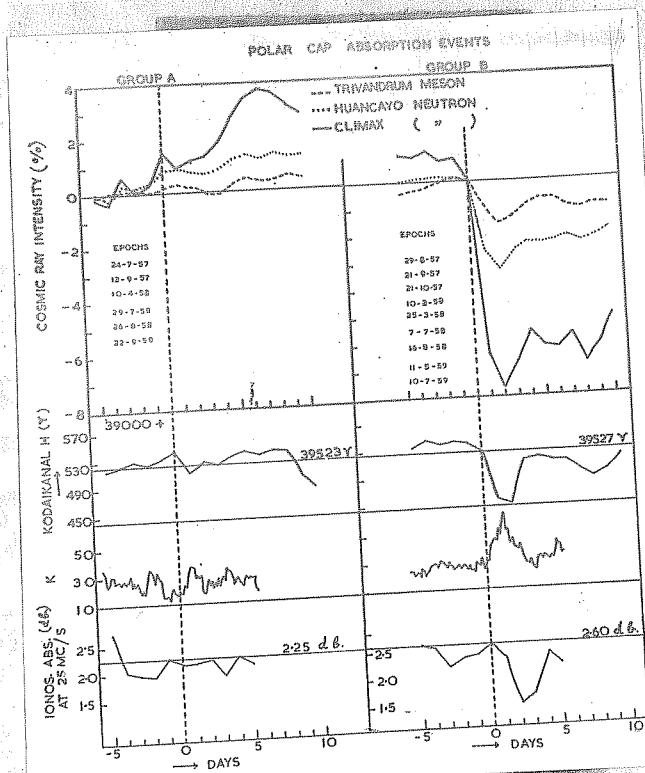
(f) Average daily variation for tritium meson and for neutron intensity at Huancayo, Cienza and Mt. Wellington for taken day groups comprising (-3, -2, -1), (0, +1, +2) and (+3, +4, +5).

In Figure 27, for group A events, the latitude, time, II, ionospheric absorption, etc. have been plotted as deviations from the mean value for the three days -5, -1 and +3 with respect to the epoch. For group B events, the reference level of intensities is taken as the value on each day. The cosmic ray intensities are shown for meson intensity at Trivandrum and neutron intensity at Huancayo, both stations at the equator, and the neutron intensity at Cienza, a high latitude equatorial station. Table 11 gives the results of chi-square analysis of daily mean intensity of meson and neutrons for group B events.

The average daily variations needed for A and B groups separately are shown in figure 28 and the parameters of the harmonic components are given in table 12.

For the harmonic analysis, the following

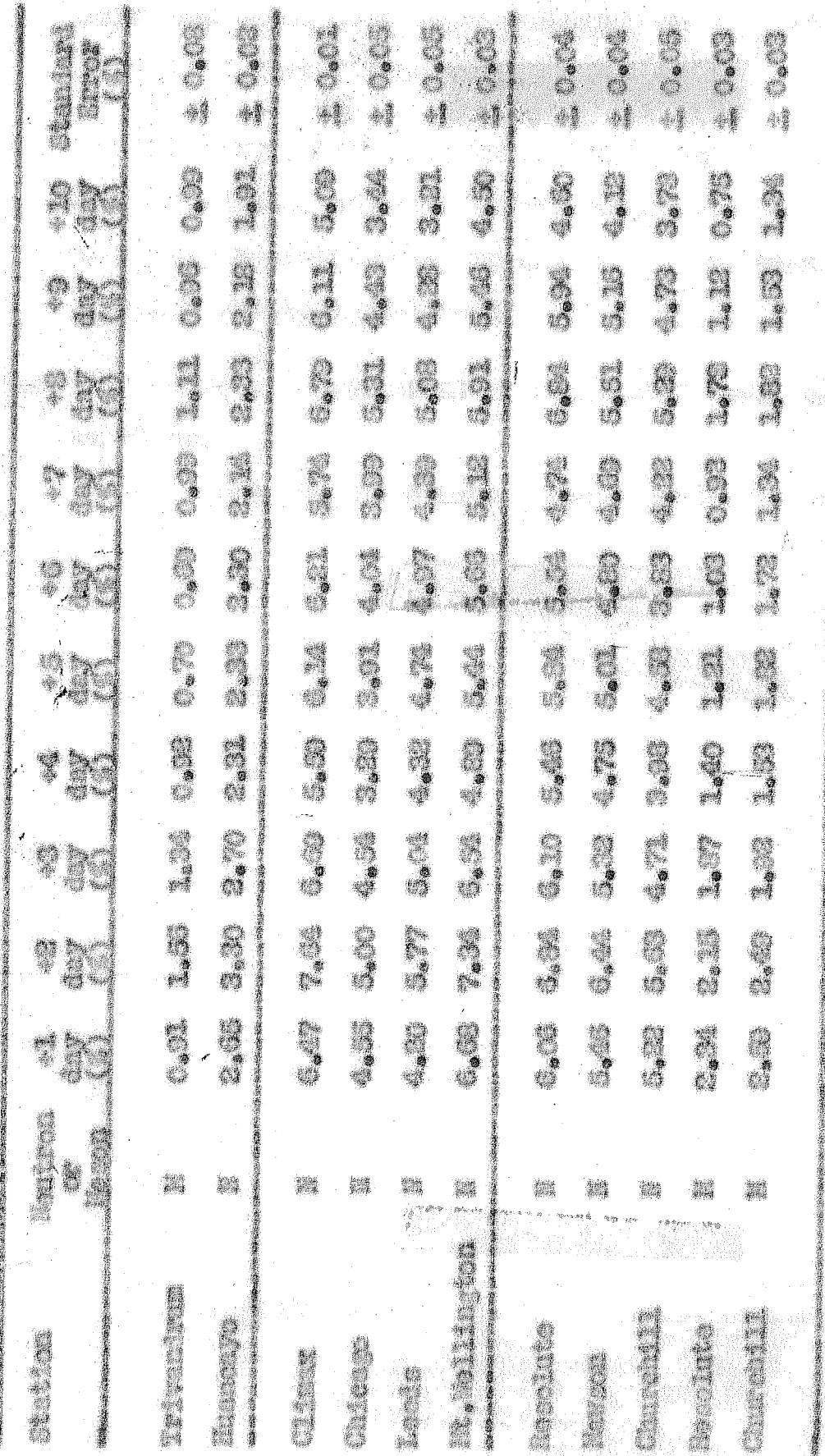
and much simpler form is used:



**Fig. 37 - Results of direct analysis of daily mean intensity of groups of absorption events. Results of analysis of Huancayo and Climax formant absorption events at 25 Mc/s. Comparison of results from ionos. absorption and direct analysis of absorption events for groups A and B.**

(2) **Huancayo and Climax Ionos. Absorption**  
stations were above normal on each day, from +6 to +8 day the transition interval. On 7th day, intensity is ~3.0% above mean level for Climax and for one of the

do que o do topo da estrada de pedestres



polar evolutions. The mean intensity at 1000 miles is also shown below from +6 to -2 days.

(3) Peak to peak amplitudes of daily variation and the diurnal and semi-diurnal components are high for the group comprising +3, +4 and +5 days as expected and high latitude stations.

(4) Rotational drift slightly lower by about 30° or +2 days.

(4)  $C_s$  is slightly above normal value on +1 day.

Fig. 123 - Average daily variation for group A and group B for groups of three day groups comprising (+3, +4, +5) days, (+6, +7, +8) days and (+9, +10, +11) days respectively.

Table showing the parameters  $\Sigma_1$ ,  $\Sigma_2$ ,  $\Sigma_3$  of the density distributions of the

stars in the cluster at different stages of evolution. The values of  $\Sigma_1$ ,  $\Sigma_2$ ,  $\Sigma_3$  are given in units of  $10^{-10} \text{ cm}^{-3}$ .

|         | $\Sigma_1$ | $\Sigma_2$ | $\Sigma_3$ |
|---------|------------|------------|------------|
| Initial | 100.0      | 100.0      | 100.0      |
| 1       | 100.0      | 100.0      | 100.0      |
| 2       | 100.0      | 100.0      | 100.0      |
| 3       | 100.0      | 100.0      | 100.0      |
| 4       | 100.0      | 100.0      | 100.0      |
| 5       | 100.0      | 100.0      | 100.0      |
| 6       | 100.0      | 100.0      | 100.0      |
| 7       | 100.0      | 100.0      | 100.0      |
| 8       | 100.0      | 100.0      | 100.0      |
| 9       | 100.0      | 100.0      | 100.0      |
| 10      | 100.0      | 100.0      | 100.0      |
| 11      | 100.0      | 100.0      | 100.0      |
| 12      | 100.0      | 100.0      | 100.0      |
| 13      | 100.0      | 100.0      | 100.0      |
| 14      | 100.0      | 100.0      | 100.0      |
| 15      | 100.0      | 100.0      | 100.0      |
| 16      | 100.0      | 100.0      | 100.0      |
| 17      | 100.0      | 100.0      | 100.0      |
| 18      | 100.0      | 100.0      | 100.0      |
| 19      | 100.0      | 100.0      | 100.0      |
| 20      | 100.0      | 100.0      | 100.0      |
| 21      | 100.0      | 100.0      | 100.0      |
| 22      | 100.0      | 100.0      | 100.0      |
| 23      | 100.0      | 100.0      | 100.0      |
| 24      | 100.0      | 100.0      | 100.0      |
| 25      | 100.0      | 100.0      | 100.0      |
| 26      | 100.0      | 100.0      | 100.0      |
| 27      | 100.0      | 100.0      | 100.0      |
| 28      | 100.0      | 100.0      | 100.0      |
| 29      | 100.0      | 100.0      | 100.0      |
| 30      | 100.0      | 100.0      | 100.0      |
| 31      | 100.0      | 100.0      | 100.0      |
| 32      | 100.0      | 100.0      | 100.0      |
| 33      | 100.0      | 100.0      | 100.0      |
| 34      | 100.0      | 100.0      | 100.0      |
| 35      | 100.0      | 100.0      | 100.0      |
| 36      | 100.0      | 100.0      | 100.0      |
| 37      | 100.0      | 100.0      | 100.0      |
| 38      | 100.0      | 100.0      | 100.0      |
| 39      | 100.0      | 100.0      | 100.0      |
| 40      | 100.0      | 100.0      | 100.0      |
| 41      | 100.0      | 100.0      | 100.0      |
| 42      | 100.0      | 100.0      | 100.0      |
| 43      | 100.0      | 100.0      | 100.0      |
| 44      | 100.0      | 100.0      | 100.0      |
| 45      | 100.0      | 100.0      | 100.0      |
| 46      | 100.0      | 100.0      | 100.0      |
| 47      | 100.0      | 100.0      | 100.0      |
| 48      | 100.0      | 100.0      | 100.0      |
| 49      | 100.0      | 100.0      | 100.0      |
| 50      | 100.0      | 100.0      | 100.0      |
| 51      | 100.0      | 100.0      | 100.0      |
| 52      | 100.0      | 100.0      | 100.0      |
| 53      | 100.0      | 100.0      | 100.0      |
| 54      | 100.0      | 100.0      | 100.0      |
| 55      | 100.0      | 100.0      | 100.0      |
| 56      | 100.0      | 100.0      | 100.0      |
| 57      | 100.0      | 100.0      | 100.0      |
| 58      | 100.0      | 100.0      | 100.0      |
| 59      | 100.0      | 100.0      | 100.0      |
| 60      | 100.0      | 100.0      | 100.0      |
| 61      | 100.0      | 100.0      | 100.0      |
| 62      | 100.0      | 100.0      | 100.0      |
| 63      | 100.0      | 100.0      | 100.0      |
| 64      | 100.0      | 100.0      | 100.0      |
| 65      | 100.0      | 100.0      | 100.0      |
| 66      | 100.0      | 100.0      | 100.0      |
| 67      | 100.0      | 100.0      | 100.0      |
| 68      | 100.0      | 100.0      | 100.0      |
| 69      | 100.0      | 100.0      | 100.0      |
| 70      | 100.0      | 100.0      | 100.0      |
| 71      | 100.0      | 100.0      | 100.0      |
| 72      | 100.0      | 100.0      | 100.0      |
| 73      | 100.0      | 100.0      | 100.0      |
| 74      | 100.0      | 100.0      | 100.0      |
| 75      | 100.0      | 100.0      | 100.0      |
| 76      | 100.0      | 100.0      | 100.0      |
| 77      | 100.0      | 100.0      | 100.0      |
| 78      | 100.0      | 100.0      | 100.0      |
| 79      | 100.0      | 100.0      | 100.0      |
| 80      | 100.0      | 100.0      | 100.0      |
| 81      | 100.0      | 100.0      | 100.0      |
| 82      | 100.0      | 100.0      | 100.0      |
| 83      | 100.0      | 100.0      | 100.0      |
| 84      | 100.0      | 100.0      | 100.0      |
| 85      | 100.0      | 100.0      | 100.0      |
| 86      | 100.0      | 100.0      | 100.0      |
| 87      | 100.0      | 100.0      | 100.0      |
| 88      | 100.0      | 100.0      | 100.0      |
| 89      | 100.0      | 100.0      | 100.0      |
| 90      | 100.0      | 100.0      | 100.0      |
| 91      | 100.0      | 100.0      | 100.0      |
| 92      | 100.0      | 100.0      | 100.0      |
| 93      | 100.0      | 100.0      | 100.0      |
| 94      | 100.0      | 100.0      | 100.0      |
| 95      | 100.0      | 100.0      | 100.0      |
| 96      | 100.0      | 100.0      | 100.0      |
| 97      | 100.0      | 100.0      | 100.0      |
| 98      | 100.0      | 100.0      | 100.0      |
| 99      | 100.0      | 100.0      | 100.0      |
| 100     | 100.0      | 100.0      | 100.0      |

(5) There are no clear changes in longitudinal absorption at absorption.

(6) Only three events are having SC change following them within one or two days, out of which only one is positive.

(7) All probable flares corresponding to these events occurred on the western half of the solar disc.

(8) Absorption at 27.6 Mc/s is comparatively less, mean value for 6 events being  $\sim 3$  dB.

(9) The average delay between the occurrence of probable flares and PEA is small  $\sim 2$  hours for the six events.

#### Characteristics of group II events

(1) Ground ray intensities show abrupt fall from the epoch day to +1 day and reach minimum value on +2 day. The intensity recovery is slow.

(2) Peak to peak amplitude of daily variation is maximum for the group of days comprising epoch, +1 and +2 days.

(3) Total annual E is lower by about 70% on +1 day, the depreciation being much larger than in group I events.

(4) C<sub>p</sub> reaches minimum value on +1 day, the changes being considerably larger than for a group of

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events. Three hourly indices of  $K_p$  are also simultaneously available on +1 and +2 days.

(6) Daily variation and the daily mean  
geomagnetic absorption at Ahmedabad are low on +2 and  
+3 days.

(6) 6 out of 9 events are followed by 3  
storms. Out of them one is either moderately severe  
or severe.

(7) The time lag is preference for either  
the eastern or western half on the solar disc.

(8) Geomagnetic absorption at 37.5 Mc/s  
is more in this case, average for 9 events being >12 db.

(9) The delay in the PEA after the occurrence  
of the flare is comparatively high. On one occasion,  
it is ~ 60 hours. The average time delay for the 9 events  
is ~ 18 hours and excluding the one event, the average  
delay is ~ 8 hours.

It is of interest to compare our observation  
with the study of Ohyaichi and Iwama<sup>177</sup>, which fully  
concerns with the association of geomagnetic storms,  
PEA and type IV solar outbursts. This is summarized  
in table 3 of their paper which is reproduced below:-

## Solar Flares and Ionization in the Earth's Upper Atmosphere.

### Cloud

Electromagnetic soft X rays radio noise  
radiation and ionospheric and major radio  
outbursts, aurorae, geomagnetic outbursts,

### Coronal cloud

|                                        |                                        |                                                  |
|----------------------------------------|----------------------------------------|--------------------------------------------------|
| <u>radioactive</u><br><u>electrons</u> | <u>synchronous</u><br><u>radiation</u> | <u>Type IV out-</u><br><u>bursts,</u>            |
| <u>radioactive</u><br><u>protons</u>   | <u>solar cosmic ray</u>                | <u>Original increase</u><br><u>of cosmic ray</u> |
| <u>radioactive</u><br><u>neutrons</u>  | <u>solar cosmic ray</u>                | <u>Radio-ray unde-</u><br><u>reco</u>            |

Auroral youself  
(trapped in the  
cloud)

Auroral zone  
blackouts

Auroras, solar  
magnetic storms

Coronal disturbance Coronal disturbance Dst of magnetic  
disturbance disturbance disturbance storms.

These comments on the aurora may suffice and  
on "aurora".

It is worth while to note here that the Forbush  
type decrease of cosmic ray intensity are usually correlated  
with those of the high-energy type. An intimate relation  
between Type XI outbursts and cosmic-ray storms (Forbush  
decreases) has been noted by Tsuruta and Iida (1960). Since  
it is generally believed that cosmic-ray storms are caused

by the shielding of incident ionizing cosmic-ray particles due to clouds carrying charged magnetic fields, the present results may support the idea that the cometary clouds of high-energy type are magnetic. As shown in Figure 8, a large Forbush-type decrease was observed simultaneously all over the world on October 21-22, 1957, and no appreciable decrease was noticed on November 6-7, 1957. In the case of a large Det. 210A of the geomagnetic storm such as on February 13, 1958, the geomagnetic intensities at middle and low latitudes show considerable short-time increases (Kondo, Nagashima, Tomida, and Ueda, 1960), though they are usually superposed on the Forbush-type decreases. Clouds of low-energy type, on the contrary, will tend to enhance the cosmic-ray intensity observed on the earth, owing to the distortion of the outer geomagnetic field.<sup>27</sup>

Since we have three types of clouds, we have, according to Chapman and Ferraro<sup>27</sup>, either high-energy cometary clouds or a combination of high and low-energy clouds. We find that magnetized those with the highest energy, i.e. which generally have the shortest time as in group A, produce strong geomagnetic effects and almost no cosmic-ray effects. Thus Chapman and Ferraro's classification can be extended to a third category, consisting of events where relativistic protons are emitted without any major emission of either low or high-energy plasma clouds. The originally plasma clouds may be

In group A events from -3 to +7 days represented in many cases a recovery from earlier events 1/24 hours apart and this indicated that the conditions for the removal of relativistic protons by themselves had not improved with the interplanetary magnetic fields have been stretched out toward the earth through earlier high or low energy plasma clouds.

#### 5.4. Events Selected

For a study of the intensity changes during several short storms, the following selection has been used for selecting the events taken. The daily mean intensity of the measured component at the 1000 hr standard for the period September 1957 to December 1960 and days are selected on which the decrease in intensity during 24 hours is more than 8 %. The twelve events may correspond to this condition observed on the following dates:-

|           |           |           |           |
|-----------|-----------|-----------|-----------|
| September | 20, 1957. | January   | 12, 1958. |
| October   | 22, 1957. | May       | 12, 1958. |
| November  | 20, 1957. | July      | 12, 1958. |
| December  | 20, 1957. | July      | 15, 1958. |
| March     | 20, 1958. | September | 4, 1958.  |
| August    | 27, 1958. | December  | 9, 1958.  |

None of these individual storms have been studied by various workers, following the intensity changes in short intervals, continuing even upto a few

stitution. The comparatively low counting rate of the Geiger counter at Trivandrum does not permit this type of study. Hence, an attempt is made here to examine only the gross effects of the 12 cosmic ray storms. Results of the gross analysis of the daily mean intensity of muons at Trivandrum and neutron intensity at Huancayo, Climax and Mt. Wellington are given in Figure 30. The intensities are represented as deviations from mean of -3, -6 and -1 day from the epoch. The gross analysis of  $C_p$  has also been done for three days and the changes are represented in the same manner.

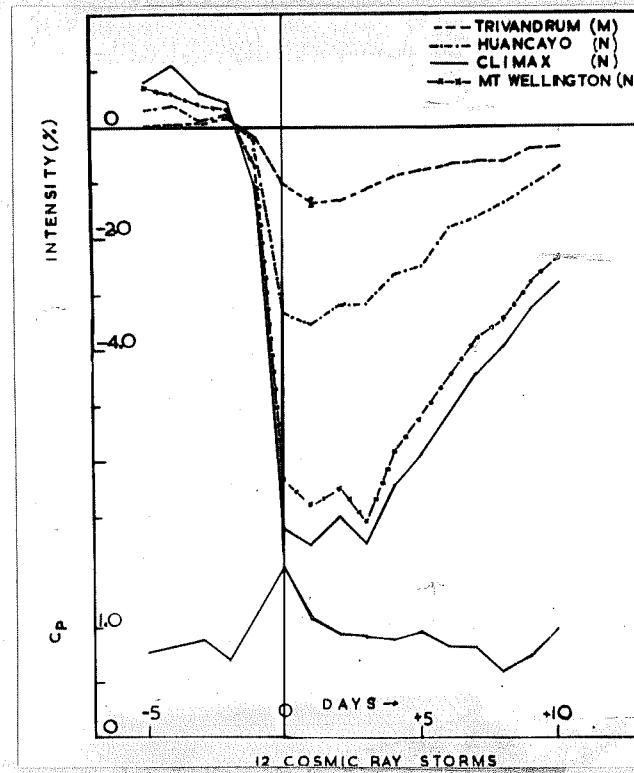


Fig. 30 - Results of gross analysis of daily mean intensity of muons at Trivandrum, neutron intensity at Huancayo, Climax and Mount Wellington and of  $C_p$  for 12 cosmic ray storms.

It is seen from the figure that the maximum depression in the intensity occurs one day after the epoch, even though the fall in intensity is very steep from -1 day to epoch day. The decreases in intensity from the reference level on different days are shown in table 13. It is seen from the figure that the recovery of the intensity to the preepoch level takes more than 10 days (from all the curves are added together). In the case of individual atoms, it has been noted that the recovery time is much shorter than this in some cases.

The ratios of decreases of neutron to muon intensity at equatorial and polar latitudes on days following the cosmic ray storm are also shown in table 13. The equatorial changes of the ratio at equatorial and polar latitudes show an almost opposite trend to each other on the epoch and the first two days after the epoch. The ratio of neutron to muon intensity at equatorial latitudes is high on epoch day and decreases on +1 and +2 days after the epoch, indicative of the spectrum being flat on the epoch day and becoming flatter thereafter. The ratio again increases on +3 and +4 days after the epoch. On the other hand, the corresponding ratio at polar latitudes is low on the epoch day and increases on +1 day. It is low again on +3 day and thereafter it increases progressively till the 6th day following the epoch. No changes are indicative of the spectrum being flat on the epoch day. An peculiar behavior of the

The following describes Intensity on open day and on the days 1000  
for the epoch for 10 cosmic ray storms during the period September  
1957 to December 1960. Reference level of Intensity has been taken  
as the mean value of -3, -2 and -1 day from the epoch.

| Station            | of   | Dyadic | +1   | -2   | -3   | -4   | -5   | -6    | Standard |
|--------------------|------|--------|------|------|------|------|------|-------|----------|
|                    |      |        | 0    | 1    | 2    | 3    | 4    | 5     | error    |
| Edmonton (B)       | 1.00 | 1.80   | 1.31 | 1.23 | 0.97 | 0.60 | 0.71 | 20.06 |          |
| Hanover (B)        | 0.97 | 0.62   | 0.28 | 0.25 | 0.00 | 0.48 | 1.70 | 20.01 |          |
| Ottawa - (B)       | 7.81 | 7.01   | 7.01 | 7.03 | 6.46 | 6.91 | 6.10 | 20.01 |          |
| Chicago - (B)      | 5.57 | 5.43   | 4.92 | 5.62 | 4.93 | 4.60 | 4.37 | 20.03 |          |
| Largo - (B)        | 6.40 | 7.01   | 7.11 | 7.00 | 6.40 | 6.20 | 6.88 | 20.03 |          |
| Mt. Washington (B) | 6.30 | 6.70   | 6.32 | 7.13 | 6.70 | 6.15 | 6.47 | 20.03 |          |
| Rancho (B)         | 6.20 | 6.70   | 6.94 | 7.00 | 6.54 | 6.32 | 4.70 | 20.04 |          |
| Winnipeg (B)       | 5.01 | 5.01   | 5.07 | 5.07 | 5.32 | 4.70 | 4.91 | 20.03 |          |
| Churchill (B)      | 5.01 | 5.38   | 5.00 | 5.00 | 4.30 | 5.00 | 3.70 | 20.03 |          |
| Rancho (B)         | 2.77 | 2.40   | 2.40 | 2.05 | 1.70 | 2.00 | 1.00 | 20.03 |          |
| Charon (B)         | 3.75 | 2.50   | 2.50 | 2.50 | 1.00 | 1.00 | 0.15 | 20.03 |          |
| <hr/>              |      |        |      |      |      |      |      |       |          |
| <b>Region</b>      |      |        |      |      |      |      |      |       |          |
| Banff (B)          | 2.40 | 2.70   | 2.31 | 2.31 | 2.10 | 2.02 | 2.00 |       |          |
| Edmonton (B)       | 2.40 | 2.05   | 2.43 | 2.50 | 2.00 | 2.10 | 2.00 |       |          |
| Prov. (B)          | 2.40 | 2.05   | 2.43 | 2.50 | 2.00 | 2.10 | 2.00 |       |          |
| Ottawa (B)         | 2.47 | 2.13   | 2.20 | 2.30 | 2.05 | 2.00 | 2.00 |       |          |
| Hanover (B)        | 2.47 | 2.13   | 2.20 | 2.30 | 2.05 | 2.00 | 2.00 |       |          |

ratio of neutron to proton at different latitudes sufficient that a similar spectrum could be fitted for all the experiments at different latitudes.

Maximum value of  $E_1$  occurs on the epoch day and the value is high on -1 and +1 days.

The amplitudes of the diurnal components of the average daily variation for the 10 storms studied together on -1, 0 and +1 days for the mean intensity as determined corrected for protons and temperature changes are given below:-

| Day   | $E_1(0)$        | $\theta = E_1(0)$ | $E_2(0)$        | $\theta$   |
|-------|-----------------|-------------------|-----------------|------------|
| -1    | $0.35 \pm 0.00$ | $91^\circ$        | $0.15 \pm 0.00$ | $97^\circ$ |
| Epoch | $0.55 \pm 0.00$ | $31^\circ$        | $0.41 \pm 0.00$ | $61^\circ$ |
| +1    | $0.10 \pm 0.00$ | $15^\circ$        | $0.11 \pm 0.00$ | $25^\circ$ |

It is seen that the  $E_1$  and  $E_2$  amplitudes are high and the time of maximum of the diurnal component of average daily variation is significantly earlier on epoch day.

#### 6.6

#### Second part, month of July 1960.

An interesting group of events of great importance occurred during the month of July 1960. There were three solar flares of importance A<sup>+</sup> which occurred at 0910 UT on

20th July, 2342 UT on 16th July and 2115 UT on 15th July respectively. The subsequent solar and terrestrial effects have been described in detail in section 2.9.

In figure 30 we have presented the changes occurring in the daily mean moon intensity at Beltsville corrected for atmospheric changes of pressure and temperature. For comparison the intensity of the nucleonic component recorded at Hanover, Glens and New Millington corrected for barometric changes is also given. The value of  $C_p$  for each day has also been plotted. The intensities have been expressed as deviations from the value on 10th July 1950.

The Intensity falls in three steps from 10th July to reach the minimum value on 16th. At each step, there is a temporary fair recovery after the first decrease but the next decrease starts before the recovery is complete. The recovery from the third (16th) decrease is very slow and even by the end of July, the neutron intensity does not recover to the level on the 10th. The maximum departure in intensity on 16th which returned to normality on 10th July is given below:

|            |        |                   |
|------------|--------|-------------------|
| Beltsville | (mean) | $0.7 \pm 0.3 \%$  |
| Hanover    | (mean) | $0.7 \pm 0.1 \%$  |
| Glens      | (mean) | $24.1 \pm 0.1 \%$ |
| Glengro    | (mean) | $25.7 \pm 0.1 \%$ |

|                |           |                   |
|----------------|-----------|-------------------|
| Zoos           | (neutron) | $19.7 \pm 0.1 \%$ |
| Mt. Wellington | (neutron) | $20.3 \pm 0.1 \%$ |
| Ranisito       | (neutron) | $20.9 \pm 0.1 \%$ |
| Hawson         | (neutron) | $21.1 \pm 0.1 \%$ |
| Churchill I    | (neutron) | $20.4 \pm 0.1 \%$ |
| Ranisito       | (neutron) | $20.8 \pm 0.1 \%$ |
| Churchill II   | (neutron) | $20.0 \pm 0.1 \%$ |

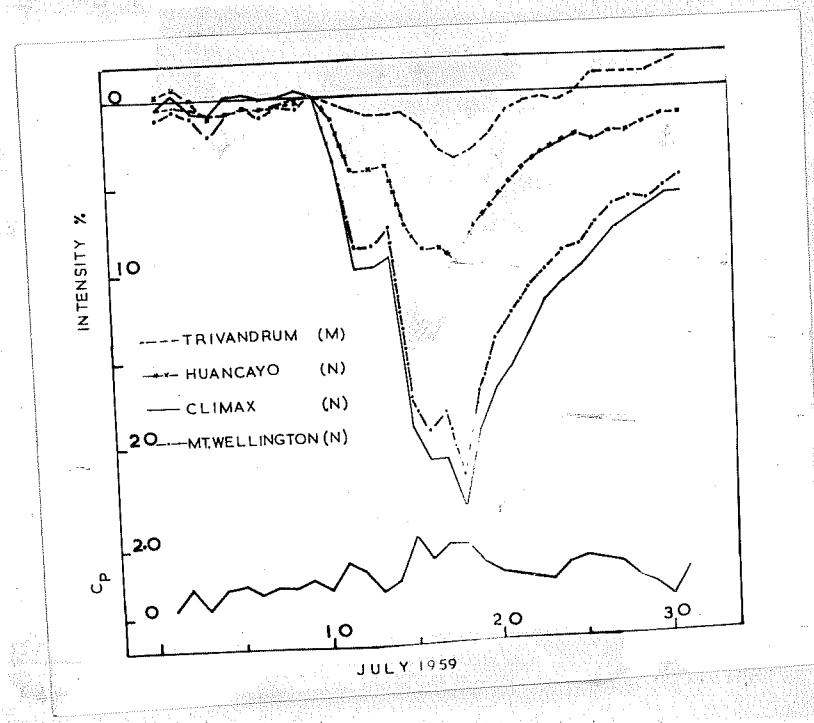


Fig. 30 - Daily neutron intensity changes of normal and enriched components during the month of July 1959.

It is seen from Figure 30 that Huancayo neutron intensity and Trivandrum neutron intensity would drop more rapidly than the intensity recorded at high latitudes after the third Torbush decrease. The three Torbush decreases are seen very clearly in neutron intensity at high latitudes and at Huancayo. At Trivandrum, the second and the third Torbush decreases almost merge together, indicative of the small amplitude of the further changes.

The ratios of the percent changes of the daily mean intensity for detectors of different mean energies of protons are indicated in Table 7 for the period July 12, 1960 to July 30, 1960. The following intensities are used:

(2) Mean value of the change in daily mean neutron intensity at Vernon, Churchill and Resolute to 2000 m.e.t. The decrease in neutron intensity at polar region [Go (ii)].

(3) Mean value of mean intensity decrease at Resolute and Churchill to represent the decrease in neutron intensity at polar region [Go (ii)].

(3) Mean value of the neutron intensity decreases at Iqaluit and change to represent the neutron intensity changes at sea level at  $\sim 50^\circ$  [Go sea level (ii)].

(4) The decrease in the neutron intensity at Olmex representing the radiogenic neutron intensity at mountain elevation at  $50^\circ$ .

(5) The decrease in neutron intensity is  
approximately proportional with increasing  
mountain elevation at equator.

(6) Neutron intensity at equator,

Table 24

Table showing the ratio of the decrease in intensity  
from the level on July 10, 1953, at different latitude  
regions and for different components.

| Ratio                    | 11/7 | 12/7 | 13/7 | 14/7 | 15/7 | 16/7 | 22/7 | 23/7 | 24/7 | 25/7 |
|--------------------------|------|------|------|------|------|------|------|------|------|------|
| Hand<br>cap (H)          | 2.50 | 4.50 | 3.50 | 3.40 | 4.00 | 3.00 | 3.78 | 3.62 | 3.32 | 3.04 |
| drum (D)                 | 1.00 | 0.90 | 0.80 | 0.70 | 0.60 | 0.50 | 0.48 | 0.45 | 0.40 | 0.37 |
| Climax (C)               | 2.71 | 2.87 | 2.50 | 2.50 | 2.00 | 2.04 | 2.37 | 2.45 | 2.50 | 2.04 |
| Hump<br>cap (H)          | 1.00 | 0.90 | 0.80 | 0.70 | 0.60 | 0.50 | 0.48 | 0.45 | 0.40 | 0.37 |
| $\lambda = 60^\circ$ (H) | 1.71 | 2.61 | 3.21 | 3.41 | 2.42 | 2.32 | 2.31 | 2.29 | 2.27 | 2.32 |
| $\lambda = 60^\circ$ (D) | 1.00 | 0.90 | 0.80 | 0.70 | 0.60 | 0.50 | 0.48 | 0.45 | 0.40 | 0.37 |
| Climax (C)               | 1.05 | 1.04 | 1.05 | 1.03 | 1.01 | 1.00 | 1.01 | 1.02 | 1.00 | 1.00 |
| $\lambda = 50^\circ$ (H) | 1.00 | 0.90 | 0.80 | 0.70 | 0.60 | 0.50 | 0.48 | 0.45 | 0.40 | 0.37 |
| level (L)                | 1.00 | 0.90 | 0.80 | 0.70 | 0.60 | 0.50 | 0.48 | 0.45 | 0.40 | 0.37 |

At the equator, the ratio of neutron to moon is  
high on the 12th and again on the 13th and 16th in the  
first and second Boronite latitudes respectively. Similarly  
the ratio of neutron to moon at  $\lambda = 30^\circ$  increases

significantly from the 11th to the 12th, but unlike at the equator it continued to increase till the 20th. After the 10th of July, 1937, the ratio increased in a similar manner as at the equator. It is interesting to observe that the ratios of Clinton neutron to Rutherford neutron and Clinton neutron to  $\lambda_0$ , zero level neutron do not exhibit the type of change observed in the ratio for neutrons to protons.

The three successive Fermat-type events are significant in relation to the steady spectrum of variation that is observed in such cases, and to discussed in detail in chapter IV.

### 5.6

#### Ion content density and associated change in daily mean intensity.

During 1937 the variation studies have been made with approximately telephones at intervals during 1937-1938 by Dr. C. E. Johnson of this report. Component has been found with colatitude inclined at  $45^\circ$  to the vertical and pointing in the east and west directions. If  $W$  and  $E$  represent the intensities measured by the west and east pointing telephones, then east-west asymmetry is defined as

$$\frac{W-E}{W+E} \times 100.$$

This quantity is represented in figure 10. The value of the east-west asymmetry is plotted against the time of day for each month. The values are plotted in pairs, one pair for each month, and the two pairs are plotted in pairs for each year. The values are plotted in pairs, one pair for each month, and the two pairs are plotted in pairs for each year.

The average value of east-west asymmetry at Ahmedabad is  $14.37 \pm 0.65$  & on days on which the asymmetry is above ( $> 15.2$  %) or below ( $< 13.0$  %), the  $S^6$  level have been classified by him as days of high and low asymmetry respectively. There are 13 low east-west asymmetry days during the period November 1957 to December 1958.

The changes of daily east asymmetry and geomagnetic associated with these 13 epochs have been studied through three analysis of

- (a) Mean and neutron intensities at different stations and at Ahmedabad.
- (b) In the horizontal component of the earth's magnetic field measured at Kodaikanal.
- (c)  $C_p$ , the interplanetary character figure.

The Geomagnetic index has been retained to ±20 days. All the intensities are expressed as deviations from the value on the -1 day, with reference to initial intensity. The results of Geomagnetic analysis for four representative intensities,  $C_p$  and  $I$  are shown in Figure 31.

An examination of the interplanetary charged particle events

- (2) The proton intensity at Ahmedabad and the neutron intensity at stations in the equatorial region show minimum value on the epoch day. Maximum neutron intensity is generally the same on epoch day and -1 day.

- (B) All stations indicated at high latitude and lower values reach minimum value on -1 day.
- (C) H, the horizontal component of the earth's magnetic field at the equatorial station Kodaikanal reaches minimum value on -3 day of the epoch and remains comparatively low even on epoch day.
- (D)  $C_p$  shows maximum value on -2 day.
- (E) The decreased asymmetry returns to the usual value (on -6 day) in about 4 to 6 days after epoch. It is seen that at lower latitudes, the recovery is faster than at higher latitudes.

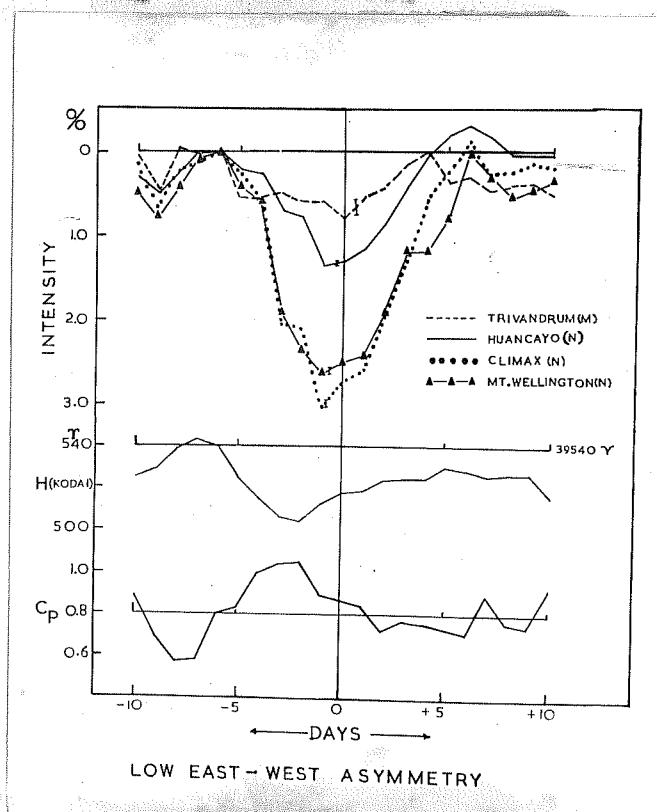


Fig. 31 - Associated changes of daily mean intensity and of polarization for the 10 October 1969 geomagnetic disturbance.

Most of these events are found to be associated with sudden commencement magnetic storms within 3 days prior to each.

These results are in very good agreement with those of Sazanov and Rao<sup>10</sup>, who have studied the associated changes for 25 low east-west asymmetry events during the period January 1957 to December 1963.

## WILSON SPECIFIC AND INVERSE OF CLASSICAL PAY OFF MATRIX

### 6.1. SPECIFIC PAY-OFFS

In chapter 3, section 3.3, we analyzed a situation as depicted by figure 3.3, here from which the players are the most popular form the possible valuation of goods by firmality. In developing a mathematical mechanism to explain the third dimension of strategy during a battle of wits, it is important to know the exact specific of valuation  $s_{111}$  and whether the valuation is discrete. In chapter 3, Wilson's basic continuity assumption at  $\lambda = 0$  and  $\lambda = \infty$  provides a large space of possible continuities. Wilson has illustrated that the primary assumption of continuity can be replaced by

$$\begin{array}{c} s_{111} \\ \text{and} \\ s_{111} \end{array} \quad \boxed{s_{111}} \quad \begin{array}{l} s_{111} < 40 \text{ pay} \\ s_{111} > 40 \text{ pay} \end{array}$$

which has replaced a version of the VPO (3.3) of the form

$s_{111} \in [0, 100]$

For the most part, Wilson's analysis has suggested that this strict upper bound of valuation for Fortune-type decisions can be replaced by

$$\begin{array}{c} s_{111} \\ \text{and} \\ s_{111} \end{array} \quad \alpha \in [0.7 \text{ to } 1.0]$$

Hoppsall and Wilson (1997) extended his analysis to the case where individual players have different pay-off functions and showed that the mixed strategy Nash

obtained on experiment 110 for the calculated energy  
dependence

In this chapter we shall discuss the method of  
calculating the variable dipole force of  $S_{(E)}$  the trans-  
formation of intensity that could be observed in trans-  
mitted light at different light source and calculated with  
different methods and criteria from transmittance. We shall  
analyze what can be deduced to calculate the observed  
distribution of changes of intensity on the materials  
by levitating only a single microgram of material.

However the one between the expected and  
observed changes is not good as found of a spectrum of  
the form  $S_{(E)} = \alpha$ , we expect to observe the change  
depending in one of the following ways:

(1) By putting a low number coefficient  $\alpha$  in  
 $S_{(E)} = \alpha$   $\approx 0.02$

(2) By putting a high number coefficient  $\alpha$  above  
which  $S_{(E)} = 0$ , [such a spectrum had been suggested by  
Froome with  $\alpha = 0$ , and  $E \approx 10$  Rev, during later work  
he found only linear intensity changes at  $\lambda = 0$  and  $\lambda = 50$ .  
He puts the value of ' $\alpha$ ' to be  $\sim 0.3$  if the reason is caused by  
decrease at  $\lambda = 50$  is  $\sim 10\%$ ]

(3) By normalizing the spectrum so that it  
comes to have a latitude dependence, and

(4) Finally, allowing a variation produced by  
change of normalized constant  $\alpha$  with its relation to wavelength  
 $S_{(E)} = \alpha_1$

We thus derive from a phenomenological model part  
of the basic characteristics of the scattering process. We  
thus assume that over two years to measure and study the  
behavior of which it is particularly likely after the onset of  
a cosmic ray shower.

#### 6.22. MODEL OF PARTIAL AND TOTAL SCATTERING

##### OF NEUTRONS

Assume that enough raw data to approximately  
correct for hadronological effects that a reasonable  
correction to applied to the neutron energy and a  
correction for magnetic and magnetic susceptibility  
changes to applied to more precisely, this reflects  
changes of the energy will be indicated by the fact the  
form of equation (3) in section 4.

The resulting correction the incident model form  
will reduce to equation (3) for and the incident has  
one level and the magnetic moment of all substances  
doped at 100 cm/cm<sup>3</sup> have been given by Fermi and  
Volberg. This can be used directly for the various  
and densities illustrated in table 6 in chapter V. Another set

#### 6.23. MODEL OF PARTIAL AND TOTAL SCATTERING

##### OF ELECTRONS

The contribution due to the effects of the particle  
ratio effect may in the intensity changes can be calculated

## Figure 30. Summary

$$\delta_{\lambda_2}^{\text{min}} = \frac{1}{\lambda_2} (\lambda_1 + \lambda_3)$$

where  $\delta_{\lambda_2}^{\text{min}}$  is the change in good quality cutoff and  
 $\lambda_2$

$\lambda_1$  is the value of the quality coefficient  
 $\lambda_3$

as  $\lambda_2$  goes from zero to infinity,  $\delta_{\lambda_2}^{\text{min}}$  increases and reaches a maximum.

In Figure 30 are illustrated the values of the  
 proposed change of quality  $\delta_{\lambda_2}^{\text{min}}(\lambda_3)$  for which a certain  
 $\lambda_2$

and different for different values of the change of quality  
 cutoff  $\delta_{\lambda_2}^{\text{min}}$ . It is seen from the figure that

(1) For the same change in  $\delta_{\lambda_2}^{\text{min}}$  the change of quality for different values of  $\lambda_3$  is not the same but the  
 less  $\lambda_3$  the greater the value.

(2) For a given quality coefficient  $\lambda_3$  the change of quality is the ratio of the changes of  $\lambda_2$  is 2 for the same change in  $\delta_{\lambda_2}^{\text{min}}$ .

(3) Comparing changes for different qualities, it  
 is seen that the coupling coefficients at  $\lambda_2^{\text{min}}$  are 7.00  
 and 1.70 and the values of  $\lambda_2^{\text{min}}$  3.77 and 3.04 for two  
 respectively. Hence for the same change in quality, the  
 ratio of the changes in fluctuation of quality and quality  
 is 4.12.

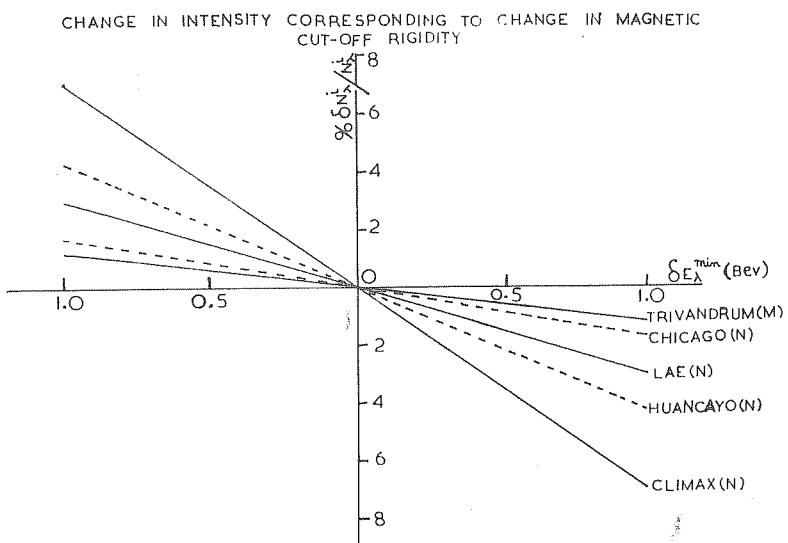


FIG. 203 - Shows that the percent change of intensity for each station and detector for change of protonotic cutoff

$\delta E_{\Delta}^{\min}$

(a) Since the protonotic and magnetotactic field in the polar region where the magnetotactic cutoff is lower than the protonotic cutoff, the magnetotactic cutoff is almost zero. Hence the change in cutoff will have an effect only at the Chicago because the magnetotactic cutoff

(a) For  $\lambda_0$  values (Hamblett and Campbell)

The results of the coupling constants in Table 20  
below 6.0 bar, where the dissociation cutoff is taken  
lower than 6.0 bar, where the dissociation cutoff is taken  
higher than 6.0 bar. Then a change in dissociation  
cutoff up to  $\sim 6.0$  bar will not affect the polarizability  
significantly.

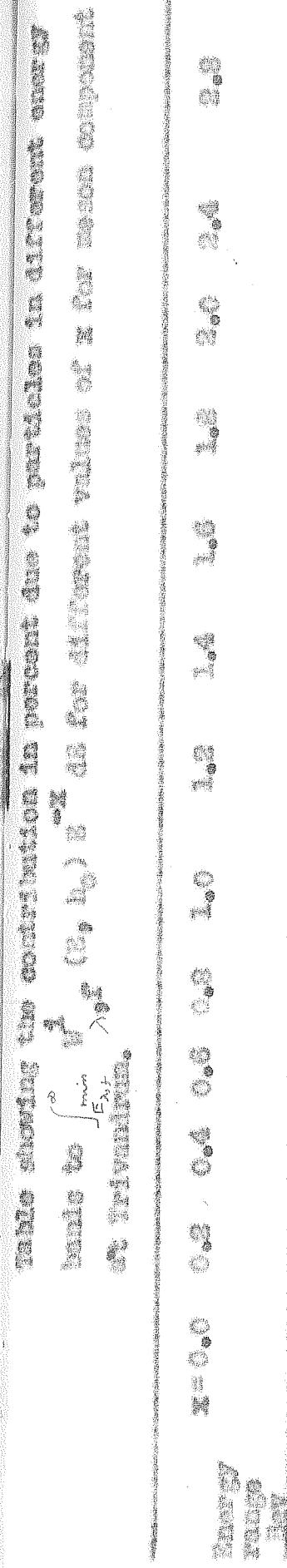
## 6.22 Effect of change of primary update.

If the most general form of the primary update  
 $S_{P(i)} = \alpha^{\frac{1}{2}}$  is chosen with values of  $\alpha = 0, 0.8, 0.4,$   
or  $-0.8$ , the percent change in polarizability for  $\alpha = 1$   
is given by

$$\frac{\delta P}{P} = \frac{1}{2} \left( \frac{\alpha}{\alpha^2 + 1} \right)^{\frac{1}{2}} \int_{-\infty}^{\infty} \frac{1}{\sqrt{1 + (\alpha/\alpha_0)^2}} d\alpha = \frac{1}{2} \dots (27)$$

To calculate the order number of this product  
using the rules for coupling constants, it is necessary  
to consider the summation over successive terms of order  $\alpha$ .  
Table 25 indicates the percent change in polarizability  
bonds for the various bond components for values of values  
of the coupling parameter  $\alpha$  listed below have been computed  
for update and reupdate components at other values also.  
When this is repeated with a parameter value of  $\alpha = 0$  it is  
to be considered, the contribution of the update values to the  
final total has to be regulated.

values of  $\sigma$  at  $10^{\circ}$ .



In Table A6, the total contribution due to  
the value of decay about the present sea level  
for the cities is estimated for the different values of  
the constant coefficient for the oceans and for the  
constant continents. It is seen that the contribution  
of Table A5 and its like

(b) For the one solution where the effective  
surface slope and the depth of the surface is  
represented by the simple linear law,

(c) In the specific problem given, the  
marked difference between Illinois and Indiana is  
that the relative depression between low land and  
sea level is high and that between mountain and sea level  
depressions is moderate. The Illinois both climate and economy are  
more  $\rightarrow$   $\infty$  in a single direction unless included others  
as building than Chicago in elevation, but has  $\lambda$  higher  
than that for Canada and Siberia. In consequence, water  
leaves the continent of the Americas a great distance  
while even though the total extent of Chicago is less  
in climate terms than in Chicago or London, the present  
distance is larger at Chicago and larger still at Chicago  
and especially at a low value of  $\lambda$ . It is expected  
that precipitation will be very cold because the supply  
is that the change at Chicago and London are smaller  
than that found at Chicago.



### 6.2 PREDICTION OF THE NUMBER OF STORMS.

In making this calculation, we have taken the distribution of events by intensity under the assumption

(1) Cosmic ray storms associated with polar cap absorption events (B group). Of the 9 events shown in figure 20, there are 4, events of a special nature. For 3 of these, viz., on August 20, 1957, on February 20, 1958 and on March 25, 1958, the intensity starts falling from prior to the first event. The fourth is the July 1958 event which is well known for the very remarkable character. A study has been made of the magnitude of all 9 events as well as the 6 events after excluding the above mentioned 4 events.

#### (2) July 1958 storm.

We estimate that cosmic rays entering at the pole absorb, the equatorial drifts and the winds at high latitude.

### 6.3 GROUP B POLAR CAP ABSORPTION EVENTS.

#### (a) Polar region.

In table 27 we present the intensity changes observed on 20 days following onset of the polar cap absorption events studied together. The details of these 9 events have been given in section 6.2. It is seen from the table that there are several significant changes in the overall mean intensity at the two stations as well as the relative intensities at the three stations. The change of opacity is used for further information than for the

distortion. This is particularly important because the second derivative of the spectrum has a number of effects which are discussed below.

In calculating dose representations of the particle-ray latitude profile, we imposed an arbitrary limit (0.6 and 0.9 % for radiation and ionization respectively), for convenience, between the energies at the boundaries which can be used the data from any one of the models if it differs by more than the permissible limits. The statistics which have been collected on a particular day are indicated in table 17. An examination of table 17 shows that even +1 to +7 day, the ratio of  $\frac{\lambda_{0.9\%}}{\lambda_{0.6\%}}$  varies steadily from a value 0.26 to a value 3.00. This indicates an increase in the value of the exponent from +1 to +7 day, or if we use an exponent 0, it requires a high energy cutoff which decreases from 26 MeV on +1 day to 25 MeV on +7 day.

We can consider the implications of a change of cumulative effect taking place on a day when the spectrum is also disturbed. A low energy cutoff in the particle spectrum occurs at 1.9 MeV corresponding to a latitude of  $63.4^\circ$  in the dipole coordinate system. This is indicated by the 'loss' of the latitude effect during the 1.6-2.2 period as reported by Baker and Anderson. If we take the effects for 6 days, this cutoff is taken to a value 2, the strength of the reduced and the exponent of the energy function applied to the 2.0 relative change of ionization and ionizing energy are shown in table 18 for +1 day of real events.

Table 32

Table showing the current and the strength of the magnet required to explain the detection of muon and neutrino interactions at polar stations on the first day of RKA events for different values of  $B'$ .

(1) without temperature correction, (2) with temperature correction of +0.5 ° and (3a) with temperature correction of -0.5 °,

for muons.

| Required current   | $\lambda_{\mu}(B)$ | $B'$                        | Gas                | Magnetic field     | Strength           |
|--------------------|--------------------|-----------------------------|--------------------|--------------------|--------------------|
| $\lambda_{\mu}(B)$ | $\lambda_{\mu}(B)$ | adjusted $\lambda_{\mu}(B)$ | $\lambda_{\mu}(B)$ | $\lambda_{\mu}(B)$ | $\lambda_{\mu}(B)$ |
|                    |                    |                             |                    |                    |                    |

(1) Without temperature correction for muons

|      |      |      |     |      |     |       |
|------|------|------|-----|------|-----|-------|
| 0.00 | 0.07 | 0.03 | 1.0 | 1.00 | 0.1 | 0.015 |
|      |      |      |     |      |     |       |
| 2.5  | 0.31 | 0.16 | 0.5 | 0.50 | 0.6 | 0.030 |
| 3.0  | 0.35 | 0.18 | 0.7 | 0.70 | 0.7 | 0.030 |
| 5.0  | 0.31 | 0.20 | 1.0 | 1.00 | 1.0 | 1.000 |
| 7.0  | 0.31 | 0.20 | 1.0 | 1.00 | 1.0 | 0.700 |
| 9.0  | 0.34 | 0.22 | 1.0 | 1.00 | 1.0 | 0.700 |

(2) With temperature correction of +0.5 ° for muons

|      |      |      |     |      |     |       |
|------|------|------|-----|------|-----|-------|
| 0.00 | 0.07 | 0.03 | 1.0 | 1.00 | 0.1 | 0.015 |
|      |      |      |     |      |     |       |
| 2.5  | 0.30 | 0.16 | 0.5 | 0.50 | 0.6 | 0.030 |
| 3.0  | 0.30 | 0.18 | 0.7 | 0.70 | 0.7 | 0.030 |
| 5.0  | 0.30 | 0.20 | 1.0 | 1.00 | 1.0 | 0.030 |
| 7.0  | 0.30 | 0.20 | 1.0 | 1.00 | 1.0 | 0.030 |
| 9.0  | 0.37 | 0.20 | 1.0 | 1.00 | 1.0 | 0.070 |

(3a) With temperature correction of -0.5 ° for muons

|      |      |      |     |      |      |       |
|------|------|------|-----|------|------|-------|
| 0.00 | 0.07 | 0.03 | 1.0 | 1.00 | 0.75 | 0.015 |
|      |      |      |     |      |      |       |
| 2.5  | 0.34 | 0.16 | 0.5 | 0.50 | 0.25 | 0.030 |
| 3.0  | 0.34 | 0.18 | 0.7 | 0.70 | 0.30 | 0.030 |
| 5.0  | 0.34 | 0.20 | 1.0 | 1.00 | 0.30 | 0.030 |
| 7.0  | 0.34 | 0.20 | 1.0 | 1.00 | 0.30 | 0.030 |
| 9.0  | 0.34 | 0.20 | 1.0 | 1.00 | 0.30 | 0.030 |

It may be argued that the higher intensity of  
radiation generally does not increase the duration of  
recovery after irradiation. This might easily result from the fact  
that with the increased dose there are more cells killed by  
the action of oxygen than by the ionizing energy  
*colder*  
in effect of heat generated during the oxidation  
process which increases with increasing dose.  
However it is also possible that the greater intensity of  
radiation intensity of dose we had minimal protection for  
the two curves are given in table II on 11 day of rea  
growth. The principal conclusion was not altered by  
changes of the atmospheric temperature of the animal  
incubator.

An explanation of why the curve that the intensity  
and mean intensity decreases at  $\lambda = 0^\circ$  can be explained  
as follows:

(a) If oxygen is first present in small  
the products of this reaction will be small and it will have  
a very low value.

(b) If oxygen is very strong exposure to which  
will be small in the reaction will have a high value and  
it will also have a high value.

These products already mentioned earlier as well  
as the dose of oxygen present in the reaction will change the very  
little.

## (b) Estimated values.

The coordinates of the stations considered in this section, namely, Gravitation and Ibarra, have been already given in table 9. The intensity decrease in 10 days following the Tia events are shown in table 10.

Table 10

Table showing the intensity decrease in 10 days following the Tia events and the value and direction of drift in the gravimeter and barograph.

| Day. | Santiago<br>( $\mu$ ) | Ibarra<br>( $\mu$ ) | Grav.( $\frac{1}{10}$ millim.) | Baro-<br>graph     |
|------|-----------------------|---------------------|--------------------------------|--------------------|
| +1   | 0.91                  | 2.60                | 8.01                           | 2.4                |
| +2   | 1.50                  | 3.20                | 2.10                           | 1.0 10.00 ± 0.50   |
| +3   | 1.90                  | 2.70                | 2.01                           | 1.6 6.20 ± 0.30    |
| +4   | 0.90                  | 2.31                | 2.02                           | 2.0 12.80 ± 0.50   |
| +5   | 0.70                  | 2.33                | 2.33                           | 2.0 200.00 ± 10.00 |
| +6   | 0.60                  | 2.30                | 2.33                           | 2.0 200.00 ± 10.00 |
| +7   | 0.50                  | 2.14                | 2.10                           | 2.0 5.20 ± 0.30    |
| +8   | 1.11                  | 2.33                | 2.10                           | 1.6 5.40 ± 0.30    |
| +9   | 0.60                  | 2.11                | 2.17                           | 1.6 4.00 ± 0.30    |
| +10  | 0.80                  | 2.01                | 2.00                           | 1.4 2.10 ± 0.15    |
| Mean | ± 0.03                | ± 0.02              |                                |                    |

If we come from this table that the relative intensity between 1000 and 10000 hours is high on +1 day, decreases on +2 and +3 days and increases again up to +6 days. On the other hand, the ratio at the polymer region is low on the first day and increases progressively until the 7th day.

If we consider that the change in intensity is only due to a variational spectrum of the form  $\delta(\omega) = \omega^{\alpha}$  the value of the exponent  $\alpha$  and the strength of the source on +1 to +10 days are indicated in table 20. We find that to explain the variation region and intensity variation intensity changes on different days, the exponent of the spectrum should be high and the corresponding values of  $\lambda_{\alpha}$  should also be high particularly on +1, +2, +3, +5 and +6 days.

If a change in the relative output intensity is associated with the onset of the source very soon and the effect of this change is superposed on the effect due to the variational spectrum, we can calculate alternatively values of the change in relative output intensity and the strength of the source  $\lambda_{\alpha}$  for different assumed values of the exponent  $\alpha$ . Table 20 indicates the values of  $\lambda_{\alpha}$  and  $\lambda_{\alpha}^{(1)}$  required for the different values of the exponent to explain the change in intensity at polymer and absorption on +1 day.

Thus, possible intensity changes at polymer and absorption intensity changes at absorption can be explained

source with a strength equivalent with a total value of  $10^{-2}$  and  
a flat spectrum ( $\alpha = 0.0$ ) with a total value of  $10^{-2}$   
is the source of the corresponding entries of the order of  
0.0 Bary.

Table 20

Table showing the value of the distance of the source in  
the unit of the corresponding value of in the strength  
of the source of equivalent strengths for all values of  
+1 day.

| Observed<br>Exponent (n)<br>Mars (n) | Calculated<br>Exponent (n)<br>Mercury (n) | Distance<br>in<br>days | Strength<br>in<br>$10^{-2}$<br>(Bary) | Strength<br>of the<br>source<br>$n \lambda_0$ |
|--------------------------------------|-------------------------------------------|------------------------|---------------------------------------|-----------------------------------------------|
| 0.34                                 | 0.31                                      | 0.4                    | - 0.03                                | 0.018                                         |
| 0.34                                 | 0.75                                      | 0.8                    | 0.02                                  | 0.062                                         |
| 0.34                                 | 0.69                                      | 0.9                    | 0.03                                  | 0.094                                         |
| 0.34                                 | 0.63                                      | 1.0                    | 0.03                                  | 0.214                                         |
| 0.34                                 | 0.59                                      | 1.2                    | 0.05                                  | 0.471                                         |
| 0.34                                 | 0.54                                      | 1.4                    | 0.04                                  | 1.061                                         |
| 0.34                                 | 0.50                                      | 1.6                    | 0.03                                  | 0.343                                         |
| 0.34                                 | 0.46                                      | 1.8                    | 0.037                                 | 0.032                                         |
| 0.34                                 | 0.43                                      | 2.0                    | 0.03                                  | 11.000                                        |
| 0.34                                 | 0.37                                      | 2.4                    | 0.10                                  | 50.000                                        |

## (a) ILLUMINANCE IN 150 RADIAN

Table 9 gives the coordinates of stations considered in this region. Column 1 is taken to represent a mountain altitude station, Lure and Chalons being more level stations. The value of  $\bar{A}_{150}$  at Lure and Chalons were the two best measurable values obtained from the total stations available. The values of  $\bar{A}_{150}$  are given in Table 9.

## Table 9

Table showing the mean of monthly overexposures in 150 days following the area around of point 1 in high latitude regions

| Days | Latitude, Longitude |       | 150 deg. Lat. 150 deg. Long. |       | 150 deg. Lat. 150 deg. Long. |     |
|------|---------------------|-------|------------------------------|-------|------------------------------|-----|
|      | (N)                 | (S)   | (N)                          | (S)   | (N)                          | (S) |
| +1   | 4.20                | 4.55  | 4.33 ± 0.13                  | 6.17  | 2.40                         |     |
| +2   | 3.77                | 5.00  | 5.33 ± 0.23                  | 7.04  | 3.42                         |     |
| +3   | 6.04                | 4.51  | 4.70 ± 0.25                  | 6.90  | 3.40                         |     |
| +4   | 4.22                | 3.93  | 4.33 ± 0.05                  | 6.30  | 3.20                         |     |
| +5   | 4.74                | 3.91  | 4.33 ± 0.42                  | 6.14  | 3.60                         |     |
| +6   | 4.97                | 4.04  | 4.30 ± 0.46                  | 6.21  | 3.30                         |     |
| +7   | 4.30                | 3.89  | 4.10 ± 0.40                  | 6.74  | 3.37                         |     |
| +8   | 5.03                | 5.01  | 5.30 ± 0.12                  | 6.70  | 3.21                         |     |
| +9   | 4.23                | 4.43  | 4.26 ± 0.07                  | 6.11  | 3.40                         |     |
| +10  | 3.21                | 3.44  | 3.33 ± 0.13                  | 6.09  | 3.20                         |     |
|      | 2220.00             | 20.06 |                              | 40.01 |                              |     |

It is about 1970.

(1) The first claim belongs to  $\lambda = 0^\circ$

LEWIS' model implies that initially one-half the value  
is available when one day has been lost or not used. This  
claim is true, however, until loss time is two days.  
Afterwards it fails.

(2) The discrepancy between the two loss

loss situations is just kept on increasing.

(3) To explain the apparent increase in

bulge and ledge with respect to the observed decrease  
at  $\lambda = 0^\circ$  requires a short excursion with a low order  
example at  $\lambda = 0^\circ$  for a loss study involving involve  
predictable discrete changes. If it is necessary approximate the  
no net values of  $l_{\text{av}}$  for  $\lambda = 0^\circ$  which are considerably  
less than those demand for an equally steep approach  
for  $\lambda = 0^\circ$ , using Raman's function and its derivative functions.  
Table 20 illustrates that  $l_{\text{av}}(\lambda)$  for  $\lambda > 0^\circ$  for the same  
relation of the approach can decrease very slightly or even  
increase depending on the strength of the slope is required.  
Even assuming that  $l_{\text{av}}$  does not decrease further as  
going to higher latitudes, the results of changes of  
values of  $\lambda = 20^\circ$  using the value of  $l_{\text{av}}$  at  $\lambda = 0^\circ$  and  
the slope function which is fixed to  $\lambda = 0^\circ$  and  $\lambda = 20^\circ$   
is too small as compared to the observed value. Thus we  
have to predict the possibility of maintaining the slopes  
without involving the additional mechanism through the  
changes of gravitational constants.

## TABLE 23

Table showing the amount of the uptake and the strength of the waves of variation for 10 days following the first exposure.

| Day. | Exposure. | $\lambda = 30^\circ$ | $\lambda = 60^\circ$ | $\lambda = 90^\circ$ |
|------|-----------|----------------------|----------------------|----------------------|
| +1   | 2.4       | 75.80 ± 3.80         | 5.00 ± 0.20          | 12.80 ± 0.70         |
| +2   | 1.6       | 15.20 ± 0.50         | 2.00 ± 0.07          | 8.30 ± 0.60          |
| +3   | 1.6       | 6.01 ± 0.60          | 1.00 ± 0.06          | 3.00 ± 0.20          |
| +4   | 0.0       | 12.30 ± 0.60         | 2.50 ± 0.08          | 7.37 ± 0.70          |
| +5   | 2.0       | 200.00 ± 25.00       | 20.00 ± 0.80         | 20.01 ± 1.00         |
| +6   | 2.0       | 200.00 ± 25.00       | 10.10 ± 0.80         | 20.80 ± 2.20         |
| +7   | 1.0       | 6.30 ± 0.30          | 1.40 ± 0.07          | 3.00 ± 0.20          |
| +8   | 1.0       | 6.05 ± 0.20          | 1.10 ± 0.06          | 3.00 ± 0.20          |
| +9   | 1.0       | 4.00 ± 0.30          | 1.00 ± 0.06          | 3.00 ± 0.20          |
| +10  | 2.4       | 3.96 ± 0.15          | 0.00 ± 0.01          | 2.07 ± 0.19          |

As at the equator, involving a change in the angular cut-off at  $\lambda = 30^\circ$  the changes of neutron absorption at Ceylon and Chicago can be explained with a low energy approach. In table 23 we show the change in the equilibrium capture required to explain the observed ratio of capture of fast neutrons at Ceylon and Chicago on the day of full exposure. The variation relation of this parameter to the equilibrium capture of the material is also given in that same table.

## TABLE 33

TABLE showing THE VALUE OF THE EXPRESSION FOR THE CHANGE IN ENERGY AND THE CORRESPONDING VALUE OF  $\lambda_{\text{DP}}$  THE LENGTH OF THE SOLVED AT  $\lambda = 60^\circ$  FOR THE OPTIMUM GROWTH.

| $\lambda_{\text{DP}}$ FOR LEVEL (60) | EXPERIMENTAL      | $\delta_{\text{DP}}$ | $\lambda_{\text{DP}}$ |
|--------------------------------------|-------------------|----------------------|-----------------------|
| OPTIMUM<br>RATIO.                    | OPTIMUM<br>RATIO. | (PPV)                | (PPV)                 |
| 0.60                                 | 0.08              | 0.4                  | 0.60                  |
| 0.62                                 | 0.02              | 0.6                  | 0.42                  |
| 0.64                                 | 0.01              | 0.8                  | 0.33                  |
| 0.66                                 | 0.02              | 1.0                  | 0.26                  |
| 0.68                                 | 0.03              | 1.2                  | 0.20                  |
| 0.70                                 | 0.04              | 1.4                  | 0.15                  |
| 0.72                                 | 0.07              | 1.6                  | 0.11                  |
| 0.74                                 | 1.00              | 1.8                  | 0.09                  |
| 0.76                                 | 1.00              | 2.0                  | 0.08                  |
| 0.78                                 | 1.10              | 2.4                  | 0.07                  |
| 0.80                                 | 1.30              | 2.8                  | 0.06                  |

## (d) REGULARITY OF GROWTH CURVES

Some evidence on the simple model for a step function model, assuming a latitudinal variation of the strength of the wind can alone explain the observed fluctuations at tropical places, we obtained the results given

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by multiplying all the factors between and dividing by a constant of proportionality introduced in figure 21 and plotted the results giving the relationship between the strength of the source and the exponent of the law of variation at  $\lambda = 0$  (given in table 20) and at  $\lambda = 80^\circ$  (given in table 23). The value of the exponent and the strength of the source corresponding to the point of intersection of the two curves with the respective changes in the predominant amplitude and frequency of motion and stated earlier are given for other days also in table 24 along with the value of the exponent and strength of the

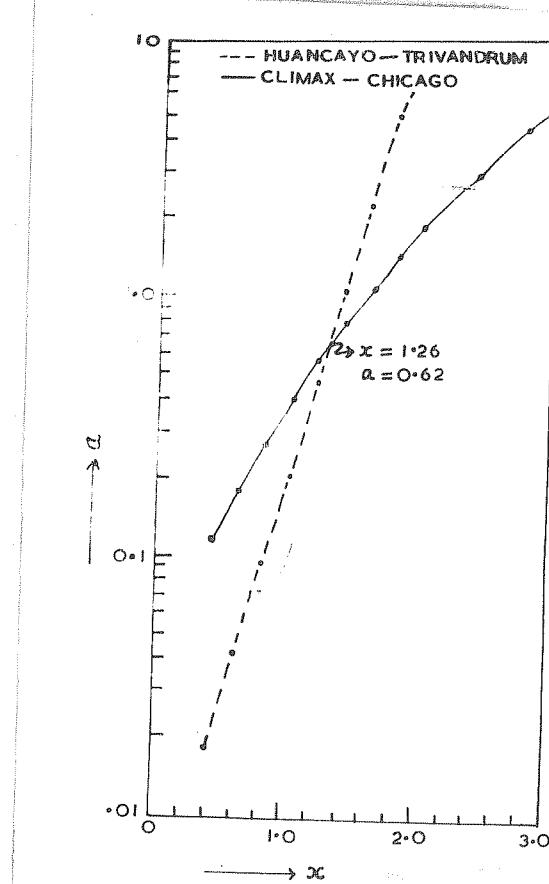


FIG. 23. - Graph showing the relation between the exponent  $x$  and the strength of the source & equivalent to point of maximum change at  $\lambda = 80^\circ$ .

source corresponding to the exposure point for various  
days following the test particle. The expected decrease  
of neutron skin mean free path at nuclear latitudes for  
such a source are also shown in the same table. It is  
clear that with such a source we cannot explain the  
 $\lambda = 80^\circ$  measurement of neutron at  $\lambda = 80^\circ$ .

Table 24

Table showing the value of the exposure, the strengths of  
the source and the required range of source which can  
explain the decrease of exposure at equatorial and high latitudes  
for different type of test particle.

| DAY | $\lambda$ | STRENGTH<br>OF THE<br>SOURCE | $\lambda = 80^\circ$ |      | $\lambda = 80^\circ$ |      | $\lambda = 80^\circ$ |      |
|-----|-----------|------------------------------|----------------------|------|----------------------|------|----------------------|------|
|     |           |                              | (km)                 | (km) | (km)                 | (km) | (km)                 | (km) |
| 1   | 1.00      | 0.40                         | 0.33                 | 0.40 | 0.65                 | 2.50 | 4.2                  | 5.5  |
| 2   | 0.85      | 0.30                         | 0.30                 | 0.38 | 1.44                 | 2.20 | 3.2                  | 4.6  |
| 3   | 0.80      | 0.30                         | 0.29                 | 0.38 | 1.00                 | 1.00 | 3.3                  | 4.7  |
| 4   | 1.71      | 1.00                         | 0.31                 | 0.40 | 0.30                 | 1.00 | 4.7                  | 6.0  |

We finally attempt to explain the values obtained  
by assuming a single source with model of the test source separation  
as well as involving a range in geographic cut-off. We  
first determine the strength and the strength of the source  
at  $\lambda = 80^\circ$ , assuming it to be at 1.6 km. (low energy cut-  
off), which can explain the observed width of neutron to

changes at these latitudes. For this purpose we have calculated the change in the geodeticetic cut-off required and the strength of the source at  $\lambda = 30^\circ$  and  $\lambda = 0^\circ$ . The results are summarized in Table 85. The strength of the source at  $\lambda = 0$  is much smaller than the strength at  $\lambda = 30$  for all the days considered. The change in cut-off required at  $\lambda = 0$  is also greater than that required at  $\lambda = 30^\circ$ . Indeed, if we assume the strength at all latitudes to be the same as at  $\lambda = 30^\circ$ , then the changes in exponent required at various latitudes are also indicated in the same table.

We thus finally conclude that no hypothesis covered with or without a change in cut-off can explain the observed fluctuations of decays. We have to invoke the change in geodeticetic cut-off as well as a latitude dependence of the strength of the source or the exponent.

### 3.32 JULY INFLUENCES

We have attempted to find the cause of dependence of the changes of intensity for the July event. The ratios of the successive observed decays at different latitudes and their changes from day to day have already been discussed in section 3.5. To determine the cause by comparison with the nature of the source of modulation, we follow the same steps as for the May events.

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The image consists of a 10x10 grid of small circular icons. Each icon contains a stylized letter, either 'S' or 'F'. When viewed as a whole, the icons are arranged to form larger letters 'S' and 'F' across the grid. The icons are white with black outlines, set against a light gray background.

Using the values calculated by the two methods we can show that neither an isotropic source with a change in cut-off nor a hyperbolic source by itself can explain the observed decreases at all places. Using the same method as given in the previous section, we have calculated the apparent, the changes in cut-off and cut-off required at  $\lambda = 30^\circ$  and  $\lambda = 0^\circ$  and the strength of the source required to explain the decreases at  $\lambda = 30^\circ$ ,  $60^\circ$  and  $0^\circ$  for some representative days. The two successive Poincaré decreases have their effect on 12th, 16th and 17th respectively.

In Table 26, the change in cut-off required and the strength of the source at  $\lambda = 0$ ,  $60$  and  $30$  for these days are given. It will be seen that on 12th, 17th and 16th we find the decreases for  $\lambda = 0^\circ$  and  $30^\circ$  analogous to those fitted for the MA events. The main difference is in the magnitude of these individual parameters. Unlike the MA events, the values of the changes in cut-off required as well as the strength of the source at all latitudes are much higher than for MA events. However, just as for MA events, the value of ' $a$ ' at  $\lambda = 0$  is smaller than at  $\lambda = 30^\circ$  and the change in cut-off at  $\lambda = 0$  is higher than at  $\lambda = 30^\circ$ .

An important point that arises, however, is that on 12th, 16th and 17th over the various parameters which we have considered above fail to explain the observed decreases at  $\lambda = 0^\circ$ . The decrease of mag-

## Table 2

Table shows the exposure, starting in of the source at different latitudes and the resulting dose rate which was calculated to explain the decreases in intensity on those days over the days during July 1962 months.

| Date    | Lat<br>(deg) | 8 hr                      | 5 hr                      | Strength of the source |                  |                  |
|---------|--------------|---------------------------|---------------------------|------------------------|------------------|------------------|
|         |              | at $\lambda = 0$<br>(deg) | at $\lambda = 0$<br>(deg) | at $\lambda = 0$       | at $\lambda = 0$ | at $\lambda = 0$ |
| 11-7-62 | 0.4          | 0.326                     | 0.301                     | 0.306                  | 0.057            | 0.0035           |
| 12-7-62 | 0.6          | 0.300                     | -                         | 0.300                  | 0.564            | -                |
| 13-7-62 | 0.6          | 0.300                     | -                         | 0.780                  | 0.652            | -                |
| 16-7-62 | 0.6          | 0.712                     | -                         | 0.742                  | 0.672            | -                |
| 17-7-62 | 0.6          | 0.98                      | 1.00                      | 0.720                  | 0.712            | 0.1670           |
| 18-7-62 | 0.6          | 0.87                      | 1.17                      | 1.016                  | 0.999            | 0.3350           |

Assuming the effect about 1/7 the decrease of radiation intensity at 10000/0 will not be a reasonable value for this situation as a reasonable source of magnetized particles for the atmosphere considered. On other days, however, we can explain the global distribution of intensity changes by a magnetosphere source and involving a change in precipitating sources.

## EFFECTS OF THE SOLAR CYCLE

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THE EFFECTS OF THE SOLAR CYCLE

THE EFFECTS OF THE SOLAR CYCLE

CONTINUATION

As we have shown that during the period 1907 to 1930, at Ahmadabad as well as at Hyderabad, there is no significant difference in the daily variation recorded by telescopes of field-angle of opening  $6^\circ \times 30^\circ$ ,  $30^\circ \times 30^\circ$  and  $30^\circ \times 37^\circ$ . This holds not only for the 12 month mean time averaged daily variation, but for the distribution of the daily variation on a day-to-day basis. Considered for the period of 1921 to 1930, during 1923 and 1924, when the former used telescope situated lower latitudes, there is a noticeable difference between Ahmadabad and Hyderabad giving a period of low solar activity and of high solar activity. This, however, requires further confirmation during the ensuing solar minimum. If systematically observed two are compared to each other the general trend of solar variation will be found more noticeable in larger angle telescope than in wider angle telescope as was found by Galileo et al.

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THE EFFECTS OF THE SOLAR CYCLE

Comparative studies of the 12 month mean daily variation at Hyderabad, Kodaikanal and Ahmadabad with

telenocopes at similar elevations show that during 1937, 1948 and 1949, the daily variations at Andean and Tropicana are comparable while the daily variation at Kodialpore has a somewhat smaller amplitude in all three years. The time of minimum of the diurnal component at all three places is comparable. Thus, apart from the lower amplitude of the daily variation at Kodialpore, which appears to be related to the lower mean density of exposure of the noon detector at relatively elevation, the instruments at all three stations respond to the same degree of the daily variation. There appears to be a greater variability in the time of minimum of the maximum at Kodialpore responding to a lower mean density of exposure, and this agrees with the observations of <sup>220</sup>Rn comparing the variability of the diurnal time of minimum in west pointing telecopes with east pointing telecopes.

There is no significant change in the amplitude of the time of minimum of the daily variation during 1937 and 1948, but during 1949 the amplitude of the 12 month mean daily variation increased on account of the decreased variability of the time of minimum from day to day.

At low latitudes, there is a significant semi-diurnal variation of mean intensity, which is not related to the seasonal variation of pressure. This is in agreement with the observation of <sup>220</sup>Rn from eastward

the general relation made by Alfvén<sup>1</sup> that a perpendicular component is associated with the daily variations produced by the primary anisotropy of cosmic rays.

### 7.3

### TYPE AND AMOUNT OF EARTHQUAKE

TYPE AND AMOUNT FOR EARTHQUAKES DURING 1937 AND 1938 WHICH ARE NOT ASSOCIATED WITH ANY SIGNIFICANT CHANGE OF COSMIC RAY INTENSITY AND PRODUCE ONLY SMALL EFFECTS ON THE ELECTROMAGNETIC FIELD. IF WE CONSIDER THESE AS TYPE I<sup>1</sup> EARTHQUAKES DIFFERENTIATING THEM FROM TYPE II<sup>1</sup> EARTHQUAKES BY MAJOR CHANGES IN COSMIC RAY INTENSITY AND IN ELECTROMAGNETIC FIELD, IT IS FOUND THAT TYPE I<sup>1</sup> EVENTS ARE ASSOCIATED WITH FLUXES ON THE WATER LINE AND, ON THE AVERAGE, WITH FLUXES WHICH ARE LARGER THAN TYPE II<sup>1</sup> EVENTS. THIS CAN BE INFERRED FROM THE EJECTION OF RELATIVISTIC PARTICLES FROM THE SUN WITHOUT ANY MAJOR CHANGES OF ORIGIN LOW OR HIGH ENERGY COSMIC RAYS AS INDICATED BY ODEBARI AND HANIS<sup>27</sup>. THE I<sup>1</sup> TYPE EARTHQUAKES WHICH DO NOT INDICATE ANY SIGNIFICANT CHANGES OF AND IN COSMIC RAY INTENSITY OR OF COSMIC RAY FIELD ARE THEREFORE DEFINITIVE IN PREDICTING AN OUTBREAK OF COSMIC RAYS 3 TO 6 DAYS AFTER THE EARTHQUAKE.

FOR THE TYPE EARTHQUAKES, TABLE 27 INDICATES THE VALUE OF CHANCE OF PAGES OF WOODENEN SAILING IN MORE THAN 50% OF EARTHQUAKES AND IN 25% OF THE EARTHQUAKES

### (a) EARTHQUAKES

TABLE 27  
EARTHQUAKES WHICH ARE ASSOCIATED WITH

(b) ~~DATA FOR THE PIA~~~~ANNUAL RATE OF CHANGE  
INSON INTENSITY AT GROUND LEVEL~~(c) ~~ANNUAL RATE OF CHANGE~~~~ANNUAL RATE OF CHANGE  
INSON INTENSITY AT GROUND LEVEL~~

Table 87

Table showing the prospective change of the ratio of inson intensity from +1 to +10 days following PIA events. 1 represents the ratio of average intensity for all 9 events, 2 for average intensity for 8 events excluding the events on August 20, 1957; February 20, 1958; March 26, 1958 and July 10, 1958, and 3 represents for the intensity following the July event on July 10, 1958.

| Day | January (1)<br>February (2) |       |      | July 10 (10)<br>July 20 (9) |      |      | August (1)<br>September (2) |      |      |
|-----|-----------------------------|-------|------|-----------------------------|------|------|-----------------------------|------|------|
|     | 1                           | 2     | 3    | 1                           | 2    | 3    | 1                           | 2    | 3    |
| 1   | 8.91                        | -2.60 | 3.50 | 8.26                        | 8.46 | 1.71 | 8.44                        | 8.61 | 8.71 |
| 2   | 8.12                        | 1.99  | 4.39 | 8.38                        | 8.71 | 8.51 | 8.02                        | 8.36 | 8.37 |
| 3   | 8.01                        | 1.92  | 3.98 | 8.01                        | 8.17 | 8.31 | 8.48                        | 8.57 | 8.30 |
| 4   | 8.01                        | 8.67  | 3.49 | 8.48                        | 8.56 | 8.41 | 8.42                        | 8.68 | 8.00 |
| 5   | 8.70                        | 8.67  | 6.90 | 8.67                        | 8.56 | 8.48 | 8.64                        | 8.91 | 8.33 |
| 6   | 8.23                        | 8.00  | 6.00 | 8.75                        | 8.90 | 8.51 | 8.70                        | 8.96 | 8.34 |
| 7   | 8.16                        | 1.60  | 2.76 | 8.06                        | 4.05 | 8.11 | 8.03                        | 8.24 | 8.37 |
| 8   | 8.10                        | 8.62  | 2.68 | 8.00                        | 4.14 | 8.23 | 8.91                        | 8.19 | 8.48 |
| 9   | 8.07                        | 8.67  | 2.92 | 8.72                        | 8.23 | 8.27 | 8.80                        | 8.20 | 8.63 |
| 10  | 1.00                        | 1.71  | 2.04 | 8.10                        | 8.00 | 8.01 | 8.33                        | 8.26 | 8.00 |

The table indicates the ratio for 10 successive days after the onset of the PIA events. The first column

In each case following to the S-B type via events (table 10), the second column relates to 6 of these after excluding 4 events of a special type as described in section 6.2. Column 3 refers to the July 1940 event where the first day corresponds to the 11th of July. Tables (a) and (b) clearly demonstrate that about 5 to 6 days after experiencing a disturbance the disturbance results in a change from predominantly at the ultraviolet maximum. Table (a) demonstrates that the initial appearance of the change always ultimately after the 6th day. The secondary disturbance also after 5 days in a general sense substantiates the relation to the possibility of any one such event followed by a cosmic ray effect.

#### 7.4 Cosmic Ray Effects

Perhaps the most important conclusion that can be drawn from the present work relating to the consideration of the type of protons which are responsible for cosmic ray effects, may perhaps have come from the fact that during several occasions, the intense intensity has a later type of effect and in addition recovery time increases increased. Many have also commented on the low intensities of charge observed at sea level stations at  $\lambda = 50^\circ$  compared to elevated stations at a comparable latitude. However, the implications of these observations will be subject to the models of the modulated air masses have been largely ignored. It is

dispersed over the present area, but there are enough  
isolated areas now, having a few inhabitants, where it  
is hard to determine the reality of these obser-  
vations and that if we attempt to take them into account,  
the apparently scattered nature and locality of such points  
expands the actual distribution of cities of size

between and within two countries and the  
allowing of the local plan of the bush villages  
will point to the scattered cities. The recent work  
demonstrates that there is a definite and very  
close to the place of the activity. This may very well  
be expected in view of the large number of  
the strength of the action of the weather or variations  
in physical conditions the former seems to be the  
more likely explanation. The intensity is most evident  
on the two days immediately following the onset of the  
fog and ends on the 2nd and the 3rd day.

The general impression obtained is that absence  
of the primary evidence by themselves cannot explain the  
global distribution of fog, even at a distance of some  
two type of secondary opinion which also requires that  
the development of suspended dust must occur. This dust  
is often found to be larger at the surface than at higher  
altitudes and larger on +1 and +2 days than on +3 and +6 days.  
The secondary effect on +3 and +6 days is predominantly  
seen at the surface rather than at higher altitudes.

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At the present moment the question is  
whether purely an international constitution and  
there may be alternative models. And it is important  
to place our interpretation in a theoretical model  
allowing the interaction of other plans in view of  
the specificities of each nation and people.  
Since the different cultures must necessarily  
follow one another, it is necessary to suppose that  
produced by the most varied social needs of the  
countries at different times they are necessarily  
conflicting. But it is evident that the different  
models are extracted from the most varied and the  
opposite. However, at the present stage of thought  
and political studies there would seem to be  
largely agreement that in the case of the conflict  
that all civil society cannot remain indifferent to  
overall human development and social justice of  
course may choose.

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